Attenuating the ice flexural wave on arctic seismic data

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Summary

- Introduction—What is the ice flexural wave, how is it excited?
- Characteristics of the flexural wave
- Noise attenuation methods
- R-T domain techniques—spectral clipping, Gabor deconvolution
- Example—Hansen Harbour
- Conclusions

What is the ice flexural wave?

- Flexural wave is often observed on floating ice in the Arctic
- Flexural wave motion is similar to that of a drum membrane
- Flexural wave can be described as P-SV internally reflected modes (Ewing et al)
- Flexural wave is excited by surface vertical source or internal compressional source in a hard layer bounded by fluids

Characteristics of the flexural wave

- Very powerful—usually the strongest wavefield on a shot record by orders of magnitude
- 2-D wave—attenuates as 1/r, not 1/r²
- Highly dispersed—high frequencies can move at ice P-wave velocity; *low* frequencies slower than air velocity
- Confined to the ice—wave does not propagate past edge of floating ice, but reflects efficiently from shore

Noise attenuation methods

- Model noise in R-T domain and subtract in X-T domain—*linear*
- Attenuate noise directly in R-T domain using spectral whitening—linear, or spectral clipping—nonlinear

Hansen Harbour CREWES 3-C

- Receiver spread—50 3-C single phones 15 metres apart
- Colinear shot line centred on receiver spread—203 shots 30 metres apart
- Dynamite and Vibroseis used as sources to record separate profiles
- Vertical component Vibroseis data used for ice wave attenuation study







Dispersion and aliasing

- Dispersion provides unique separation of ice wave frequency components in R-T domain
- Spatial aliasing compromises effectiveness of frequency separation by moving components up into seismic band
- Ideal acquisition would sample ice wave with no aliasing at any frequency

Aliasing of the ice flexural wave



No NMO—Ice wave aliased at all frequencies

Aliasing of the ice flexural wave



1000 m/s linear NMO—lower frequencies still aliased

Aliasing of the ice flexural wave



250 m/s linear NMO—higher frequencies aliased



trajectories encounter monochromatic noise, due to dispersion of ice wave



R-T fan transform of raw shot gather—ice wave on radial trace is monochromatic



Spectrum of radial trace with ice wave noise

Editing the spectrum

- Gabor deconvolution—linear, more effective on weaker, less monochromatic noise
- Spectral clipping—nonlinear, most effective on the strongest, most monochromatic noise



Seismic trace contaminated with monochromatic noise is transformed to frequency domain



Median spectrum computed from raw spectrum; threshold curves placed parallel to median; peaks in raw spectrum exceeding threshold are flagged



Flagged raw spectral amplitudes are replaced with median values, phase is unaltered



Edited spectrum is transformed back to seismic trace, sans monochromatic noise



Spectrum of radial trace with ice wave noise



Spectrum of radial trace with ice wave noise after spectral clipping



R-T fan transform of raw shot gather after spectral clipping—monochromatic noise greatly attenuated









Gabor deconvolution applied in X-T domain—R-T domain Gabor deconvolution



2.0

Gabor deconvolution applied in X-T domain—R-T domain spectral clipping



Brute stack of Hansen Harbour line with no ice wave filtering



Brute stack of Hansen Harbour line after ice wave filtering

Conclusions

- Stacking alone is not sufficient to attenuate ice wave noise
- Ice wave should be properly spatially sampled during acquisition
- R-T domain more effective than X-T domain for ice wave attenuation
- Spectral clipping marginally more effective than Gabor decon for strong ice wave

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