

# Mathematical details of time-lapse AVO framework

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## Introduction

The study described here focuses on applying the perturbation theory in time-lapse amplitude variation with offset (Time-lapse AVO) method to model a framework to describe the difference data from a baseline survey to monitor survey in a reservoir. Reflection coefficients are derived for the baseline and monitor survey using Zoeppritz equations to calculate the reflection coefficient for difference data.

## Procedure for deriving A-R from Zoeppritz equations

Consider an incident P wave striking on the boundary of a planar interface between two elastic media with rock properties  $V_{P0}$ ,  $V_{S0}$ ,  $\rho_0$  and  $V_{P1}$ ,  $V_{S1}$ ,  $\rho_1$ .

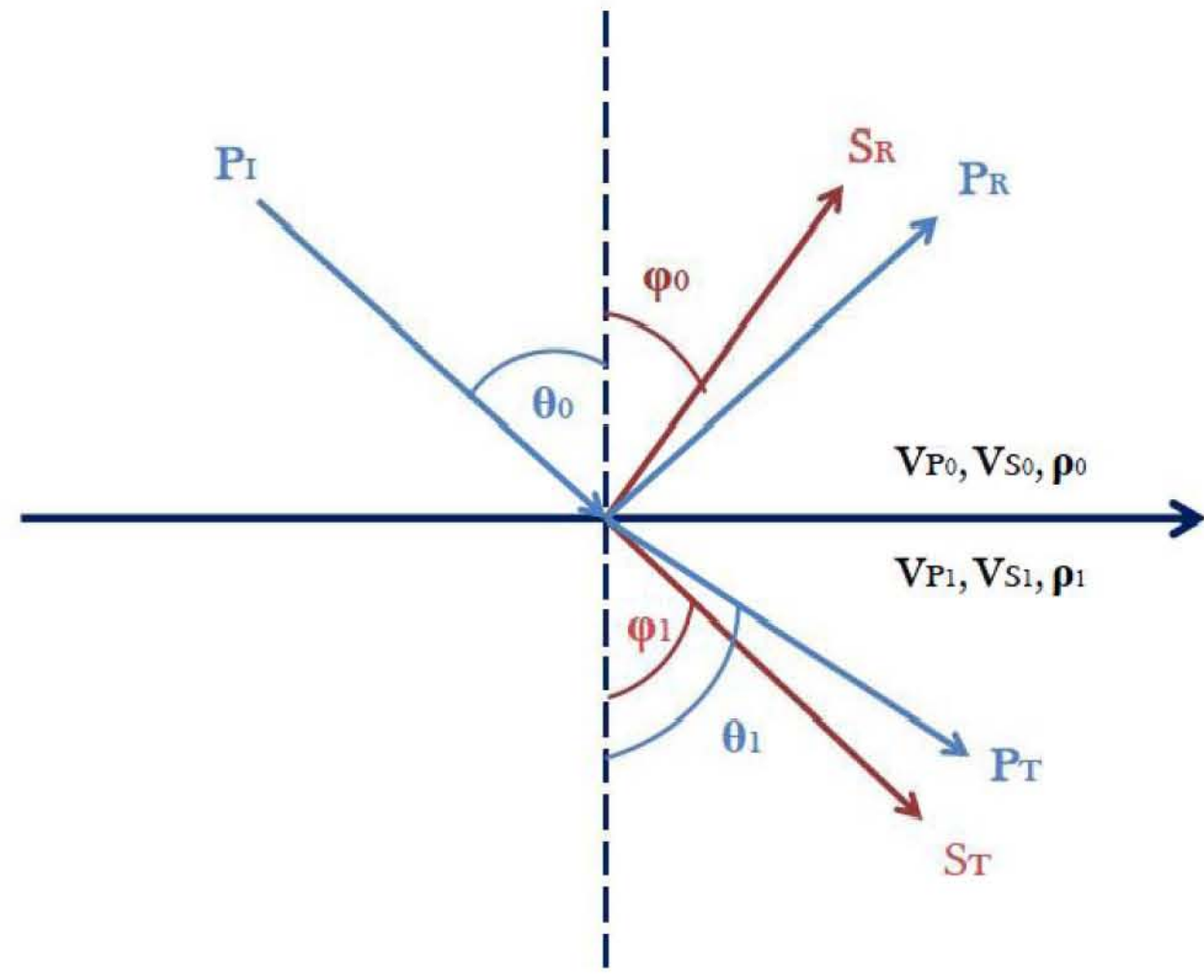


Figure 1: Displacement amplitude of an incident P-wave with related reflected and transmitted P and S waves.

Setting boundary conditions in the problem leads to Zoeppritz equations:

$$\begin{bmatrix} R_{PP} \\ R_{PS} \\ T_{PP} \\ T_{PS} \end{bmatrix} = b_P \quad R_{PP}(\theta) = \frac{\det(P_P)}{\det(P)} \quad b_P \equiv \begin{bmatrix} X \\ \sqrt{1-X^2} \\ 2B^2X\sqrt{1-X^2} \\ 1-2(BX)^2 \end{bmatrix}$$

where

$$X = \sin(\theta_0), \quad A \equiv \frac{\rho_1}{\rho_0}, \quad B \equiv \frac{V_{S0}}{V_{P0}}, \quad C \equiv \frac{V_{P1}}{V_{P0}}, \quad D \equiv \frac{V_{S1}}{V_{P0}}, \quad E \equiv \frac{V_{P1}}{V_{S0}}, \quad F \equiv \frac{V_{S1}}{V_{S0}}.$$

Then

$$P \equiv \begin{bmatrix} -X & -\sqrt{1-(BX)^2} & CX & \sqrt{1-(DX)^2} \\ \sqrt{1-X^2} & -BX & \sqrt{1-(CX)^2} & -DX \\ 2B^2X\sqrt{1-X^2} & B(1-2(BX)^2) & 2AD^2X\sqrt{1-(CX)^2} & AD(1-2(DX)^2) \\ -(1-2(BX)^2) & 2B^2X\sqrt{1-(BX)^2} & AC(1-2(DX)^2) & -2AD^2X\sqrt{1-(DX)^2} \end{bmatrix}$$

We next seek a way to expand these solutions about the contrasts across the interface. Now let's introduce perturbation parameters the wave experience traveling from medium one to two (Figure 1):

$$a_{VP} = 1 - \frac{V_{P0}^2}{V_{P1}^2}, \quad a_{VS} = 1 - \frac{V_{S0}^2}{V_{S1}^2}, \quad a_\rho = 1 - \frac{\rho_0}{\rho_1},$$

The elastic parameters are expressed in terms of perturbations parameters:

$$A = (1 - a_\rho)^{-1}, \quad C = (1 - a_{VP})^{-\frac{1}{2}}, \quad D = B \times (1 - a_{VS})^{-\frac{1}{2}}$$

$$(1 - a_\rho)^{-1} = 1 + a_\rho + a_\rho^2 + \dots \quad (1 - a_{VP})^{-\frac{1}{2}} = 1 + \frac{1}{2}a_{VP} + \frac{3}{8}a_{VP}^2 + \dots \quad (1 - a_{VS})^{-\frac{1}{2}} = 1 + \frac{1}{2}a_{VS} + \frac{3}{8}a_{VS}^2 + \dots$$

Substituting these parameters in Zoeppritz equations and expanding them about X and perturbation parameters, reflection coefficient are calculated as:

$$R_{PP}(\theta_0) = R^{(1)}_{PP}(\theta_0) + R^{(2)}_{PP}(\theta_0) + \dots$$

$$R^{(1)}_{PP}(\theta_0) = \frac{1}{4}(1 + X^2)a_{VP} - 2(BX)^2a_{VS} + \left(\frac{1}{2} - 2(BX)^2\right)a_\rho$$

$$R^{(2)}_{PP}(\theta_0) = \frac{1}{4}\left(\frac{1}{2} + X^2\right)a_{VP}^2 + (B^3X^2 - 2(BX)^2)a_{VS}^2$$

$$+ \left(\frac{1}{4} - \frac{1}{4}BX^2 - (BX)^2 + B^3X^2\right)a_\rho^2 + (2B^3X^2 - (BX)^2)a_\rho a_{VS}$$

First order is equivalent to the Aki-Richards approximation.

## A framework for time-lapse AVO

We will consider two seismic experiments involved in a time-lapse survey, the baseline survey, followed by a monitoring survey.

In a time-lapse study, there are two groups of perturbation parameters; from the first medium to the second medium in the baseline survey and from baseline to monitor survey:

$$a_{VP} = 1 - \frac{V_{P0}^2}{V_{PBL}^2}, \quad a_{VS} = 1 - \frac{V_{S0}^2}{V_{SBL}^2}, \quad a_\rho = 1 - \frac{\rho_0}{\rho_{BL}},$$

$$b_{VP} = 1 - \frac{V_{PBL}^2}{V_{PM}^2}, \quad b_{VS} = 1 - \frac{V_{SBL}^2}{V_{SM}^2}, \quad b_\rho = 1 - \frac{\rho_{BL}}{\rho_M},$$

The elastic parameters are:

$$A \equiv (1 - b_\rho)^{-1} \times (1 - a_\rho)^{-1}, \quad C \equiv (1 - a_{VP})^{-\frac{1}{2}} \times (1 - b_{VP})^{-\frac{1}{2}}, \quad D \equiv B \times (1 - a_{VS})^{-\frac{1}{2}} \times (1 - b_{VS})^{-\frac{1}{2}}$$

The same AVO method is used to calculate the reflection coefficient for the baseline and monitor survey. Difference data is then calculated as:

$$\Delta R_{PP}(\theta) = R_{PP}^M(\theta) - R_{PP}^{BL}(\theta)$$

$$\Delta R_{PP}^{(1)}(\theta_0) = \frac{1}{4}(1 + X^2)b_{VP} - 2(BX)^2b_{VS} + \left(\frac{1}{2} - 2(BX)^2\right)b_\rho$$

$$\Delta R_{PP}^{(2)}(\theta_0) = \frac{1}{4}\left(\frac{1}{2} + X^2\right)b_{VP}^2 + (B^3X^2 - 2(BX)^2)b_{VS}^2 + \left(\frac{1}{4} - \frac{1}{4}BX^2 - (BX)^2 + B^3X^2\right)b_\rho^2$$

$$+ (2B^3X^2 - (BX)^2)b_\rho b_{VS} + (2B^3X^2 - 2(BX)^2)a_{VS}b_{VS} + \left(\frac{1}{4}X^2\right)a_{VP}b_{VP}$$

$$+ (2B^3X^2 - (BX)^2)a_\rho b_{VS} + (2B^3X^2 - (BX)^2)b_\rho a_{VS} + (2B^3X^2 - \frac{1}{2}BX^2)a_\rho b_\rho$$

## Time-lapse AVO in terms of relative parameters

More recognizable and possibly numerically more accurate way is computing time-lapse difference data in terms of relative changes:

$$\frac{\Delta V_P}{V_P} = 2 \frac{V_{Pb} - V_{P0}}{V_{Pb} + V_{P0}}, \quad \frac{\Delta V_S}{V_S} = 2 \frac{V_{Sb} - V_{S0}}{V_{Sb} + V_{S0}}, \quad \frac{\Delta \rho}{\rho} = 2 \frac{\rho_b - \rho_0}{\rho_b + \rho_0}$$

$$\frac{\delta V_P}{V_P} = 2 \frac{V_{Pm} - V_{Pb}}{V_{Pm} + V_{Pb}}, \quad \frac{\delta V_S}{V_S} = 2 \frac{V_{Sm} - V_{Sb}}{V_{Sm} + V_{Sb}}, \quad \frac{\delta \rho}{\rho} = 2 \frac{\rho_m - \rho_b}{\rho_m + \rho_b}$$

Linear and second order difference data can be recalculated in terms of relative changes using the relations below:

$$a_{VP} = 2 \left( \frac{\Delta V_P}{V_P} \right) - 2 \left( \frac{\Delta V_P}{V_P} \right)^2 + \frac{3}{2} \left( \frac{\Delta V_P}{V_P} \right)^3 - \dots \quad b_{VP} = 2 \left( \frac{\delta V_P}{V_P} \right) - 2 \left( \frac{\delta V_P}{V_P} \right)^2 + \frac{3}{2} \left( \frac{\delta V_P}{V_P} \right)^3 - \dots$$

$$a_{VS} = 2 \left( \frac{\Delta V_S}{V_S} \right) - 2 \left( \frac{\Delta V_S}{V_S} \right)^2 + \frac{3}{2} \left( \frac{\Delta V_S}{V_S} \right)^3 - \dots \quad b_{VS} = 2 \left( \frac{\delta V_S}{V_S} \right) - 2 \left( \frac{\delta V_S}{V_S} \right)^2 + \frac{3}{2} \left( \frac{\delta V_S}{V_S} \right)^3 - \dots$$

$$a_\rho = \left( \frac{\Delta \rho}{\rho} \right) - \frac{1}{2} \left( \frac{\Delta \rho}{\rho} \right)^2 + \frac{1}{4} \left( \frac{\Delta \rho}{\rho} \right)^3 + \dots \quad b_\rho = \left( \frac{\delta \rho}{\rho} \right) - \frac{1}{2} \left( \frac{\delta \rho}{\rho} \right)^2 + \frac{1}{4} \left( \frac{\delta \rho}{\rho} \right)^3 + \dots$$

The linear term is in agreement with Landrø's work (Landrø et al., 2001):

$$\Delta R_{PP}^{(1)}(\theta) = \frac{1}{2} \left( \frac{\delta \rho}{\rho} + \frac{\delta V_P}{V_P} \right) - 2 \frac{V_S^2}{V_P^2} \left( \frac{\delta \rho}{\rho} + 2 \frac{\delta V_S}{V_S} \right) \sin^2 \theta + \frac{\delta V_P}{2V_P} \sin^2 \theta$$

Also second, and third order terms are calculated and can be found in the reference (Jabbari and Innanen, 2012).

## Conclusions

Time-lapse AVO measurements provide a tool to monitor the dynamic changes in subsurface properties during the time of the exploitation of a reservoir. Results showed that in plausible large contract time-lapse scenarios higher order term corrections is necessary.

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## References

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