

Extending Sommerfeld integral based spherical wavefield computations beyond two layers

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Outline

- ◆ Introduction
- ◆ Sommerfeld Integrals and seismic wave fields
- ◆ Point sources in layered half spaces: Ewing, Jardetzky and Press
- ◆ Interfaces and boundary conditions
- ◆ P-wave multiples in a three-layer system
- ◆ Class 1 AVO-response of a three-layer system
- ◆ Conclusions
- ◆ Acknowledgements

$$u_p(\omega) = i\omega e^{-i\omega t} \int_0^{\infty} \left[\frac{p^2}{\xi} J_1(\omega pr) \sin(\theta) - ip J_0(\omega pr) \cos(\theta) \right] e^{i\omega z \xi} dp$$

Direct-Wave Particle Motion Based on Sommerfeld Integral (Aki and Richards, 1980)

$$\varphi_j = \int_0^{\infty} Q_j' J_0(kr) e^{-\nu_j z} dk + \int_0^{\infty} Q_j'' J_0(kr) e^{\nu_j z} dk$$

$$\psi_j = \int_0^{\infty} S_j' J_0(kr) e^{-\nu_j z} dk + \int_0^{\infty} S_j'' J_0(kr) e^{\nu_j z} dk$$

Potential of Down-Going and Up-Going Waves in
Layer j (Ewing, Jaretzky and Press, 1957)

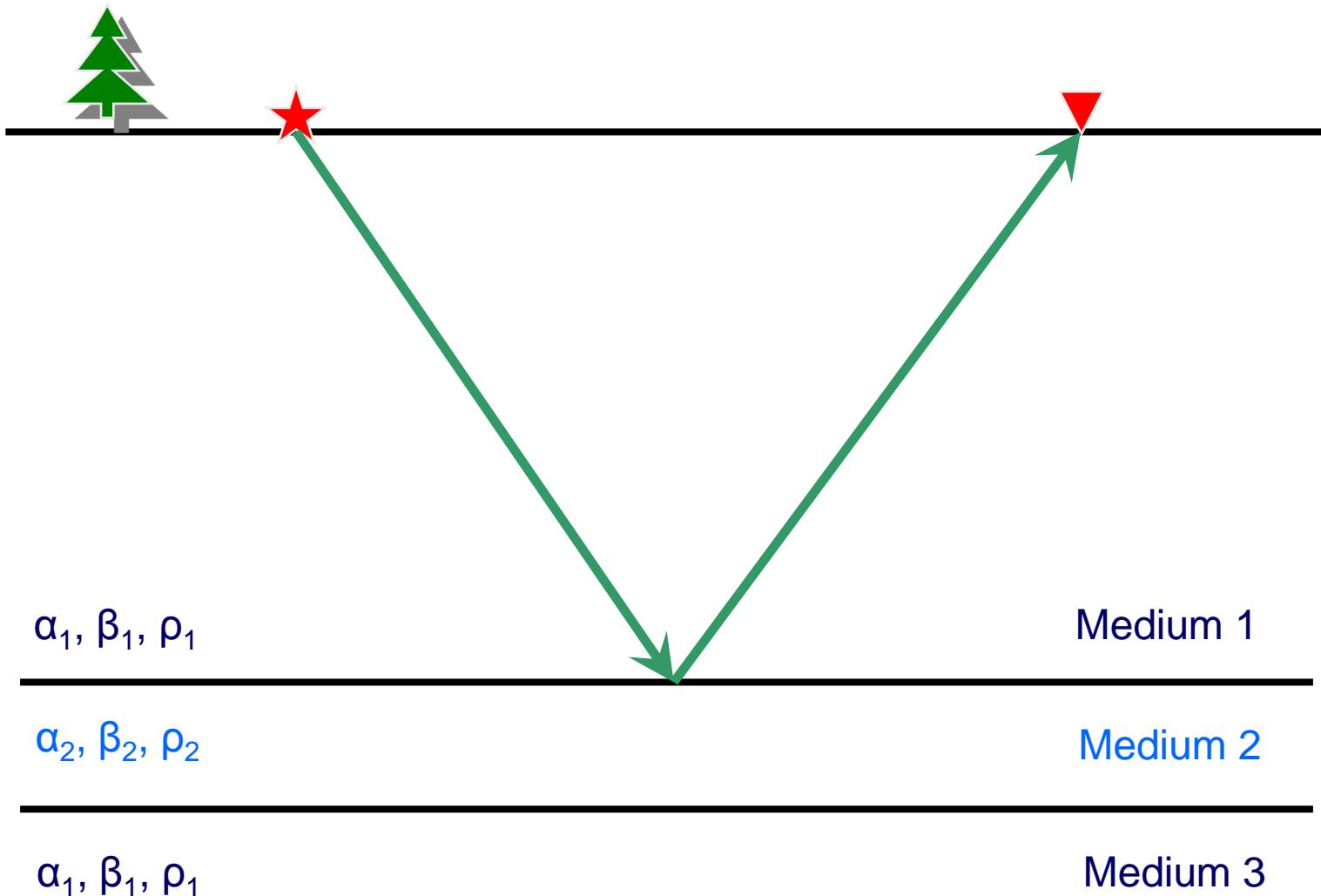
$$A_{R1}\alpha_1 p + B_{R1}\beta_1 \eta_1 - (A_{T1} + A_{R2}e^{-i\omega\xi_2 z_2})\alpha_2 p - (B_{T1} + B_{R2}e^{-i\omega\eta_2 z_2})\beta_2 \eta_2 = -A_{I1}\alpha_1 p$$

$$-A_{R1}\alpha_1 \xi_1 + B_{R1}\beta_1 p - (A_{T1} - A_{R2}e^{-i\omega\xi_2 z_2})\alpha_2 \xi_2 + (B_{T1} - B_{R2}e^{-i\omega\eta_2 z_2})\beta_2 p = -A_{I1}\alpha_1 \xi_1$$

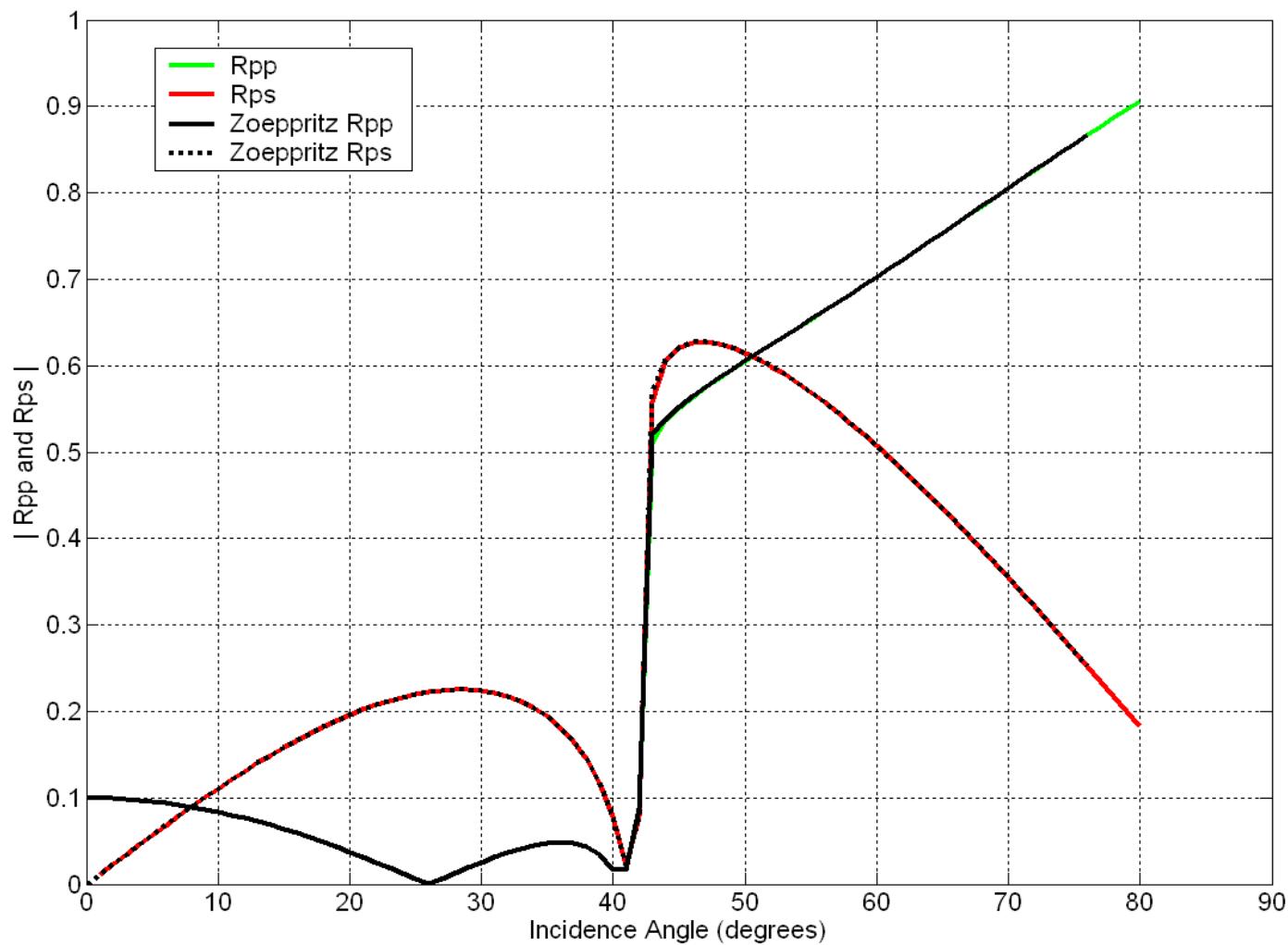
$$\begin{aligned} -2A_{R1}\alpha_1\beta_1^2\rho_1 p\xi_1 - B_{R1}\beta_1\rho_1(1-2\beta_1^2 p^2) - (A_{T1} - A_{R2}e^{-i\omega\xi_2 z_2})2\alpha_2\beta_2^2\rho_2 p\xi_2 \\ -(B_{T1} - B_{R2}e^{-i\omega\eta_2 z_2})\beta_2\rho_2(1-2\beta_2^2 p^2) = -2A_{I1}\alpha_1\beta_1^2\rho_1 p\xi_1 \end{aligned}$$

$$\begin{aligned} A_{R1}\alpha_1\rho_1(1-2\beta_1^2 p^2) - 2B_{R1}\beta_1^3\rho_1 p\eta_1 - (A_{T1} + A_{R2}e^{-i\omega\xi_2 z_2})\alpha_2\rho_2(1-2\beta_2^2 p^2) \\ +(B_{T1} + B_{R2}e^{-i\omega\eta_2 z_2})2\beta_2^3\rho_2 p\eta_2 = -A_{I1}\alpha_1\rho_1(1-2\beta_1^2 p^2) \end{aligned}$$

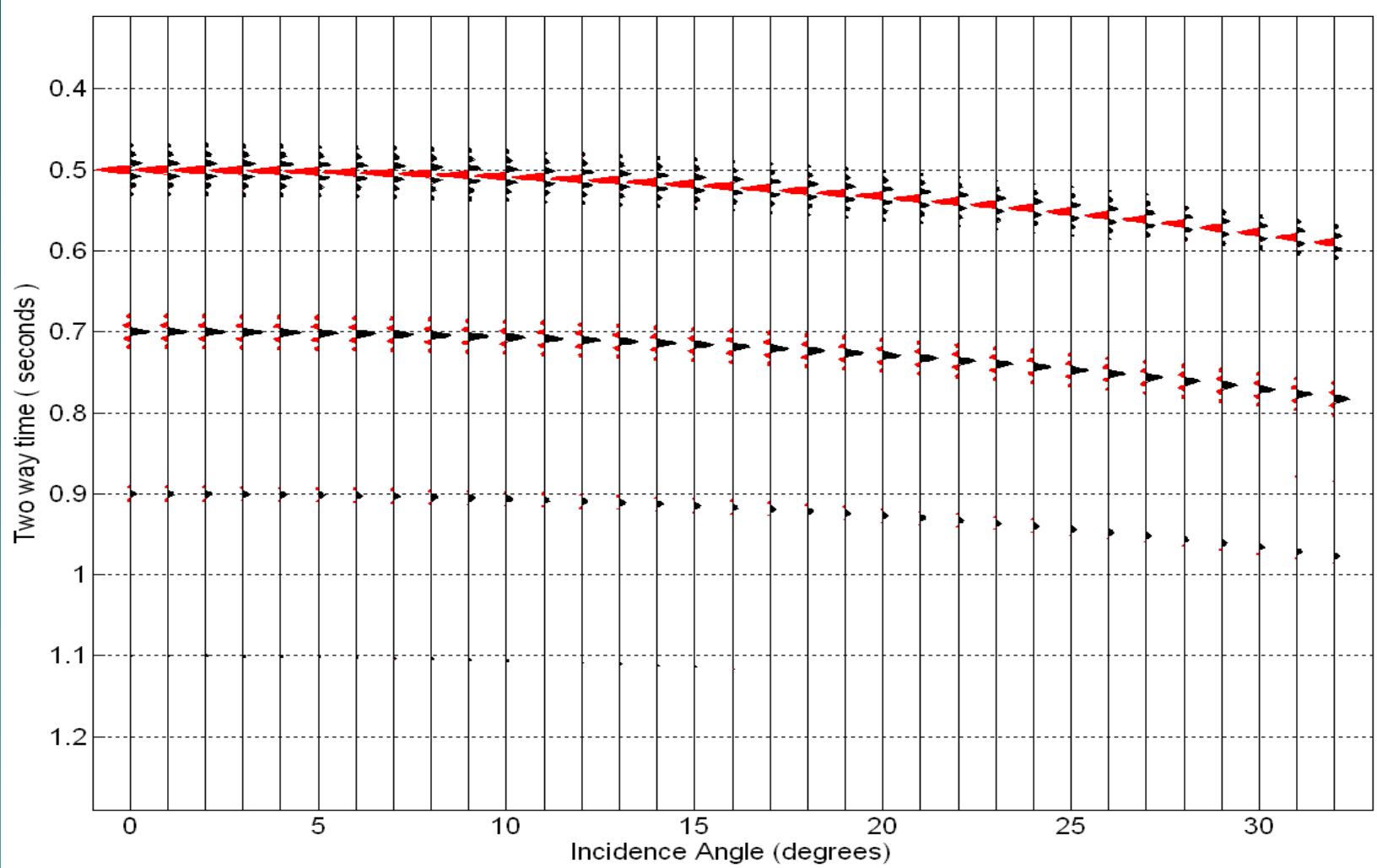
Application of Boundary Conditions at the First Interface ($z_1 = 0$)



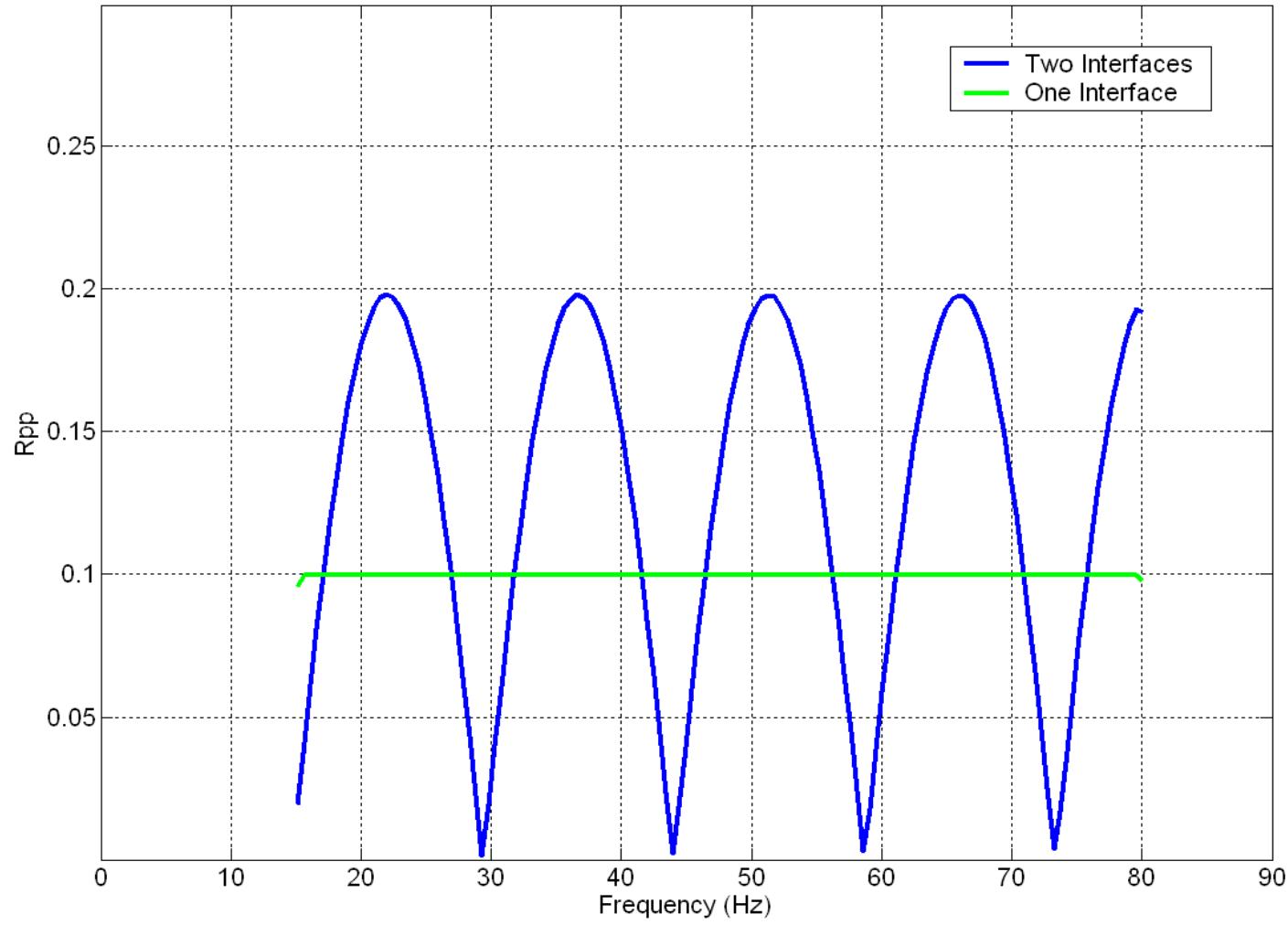
Reservoir Model Geometry



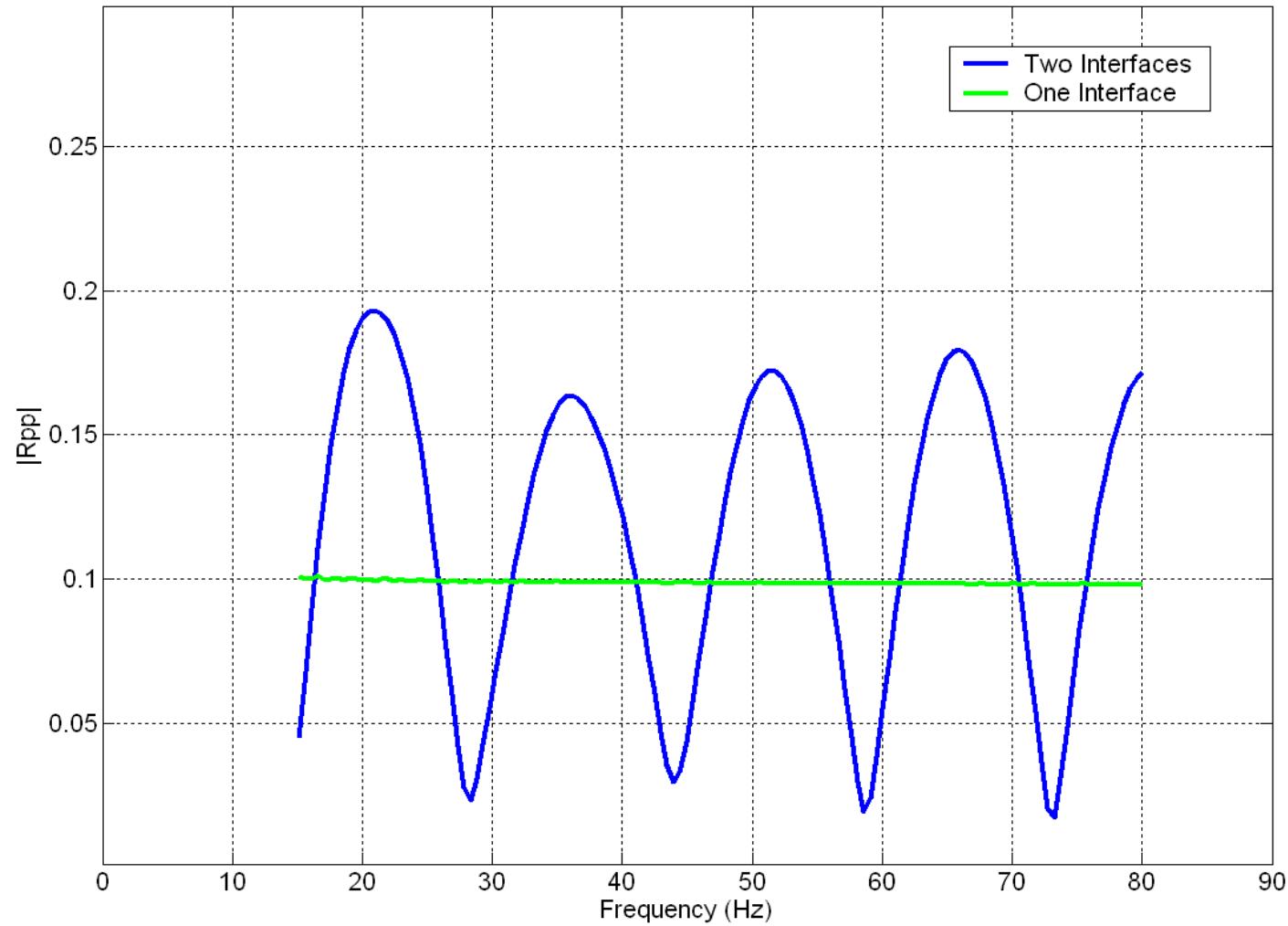
Plane-Wave Class 1 Elastic Reflection Coefficient



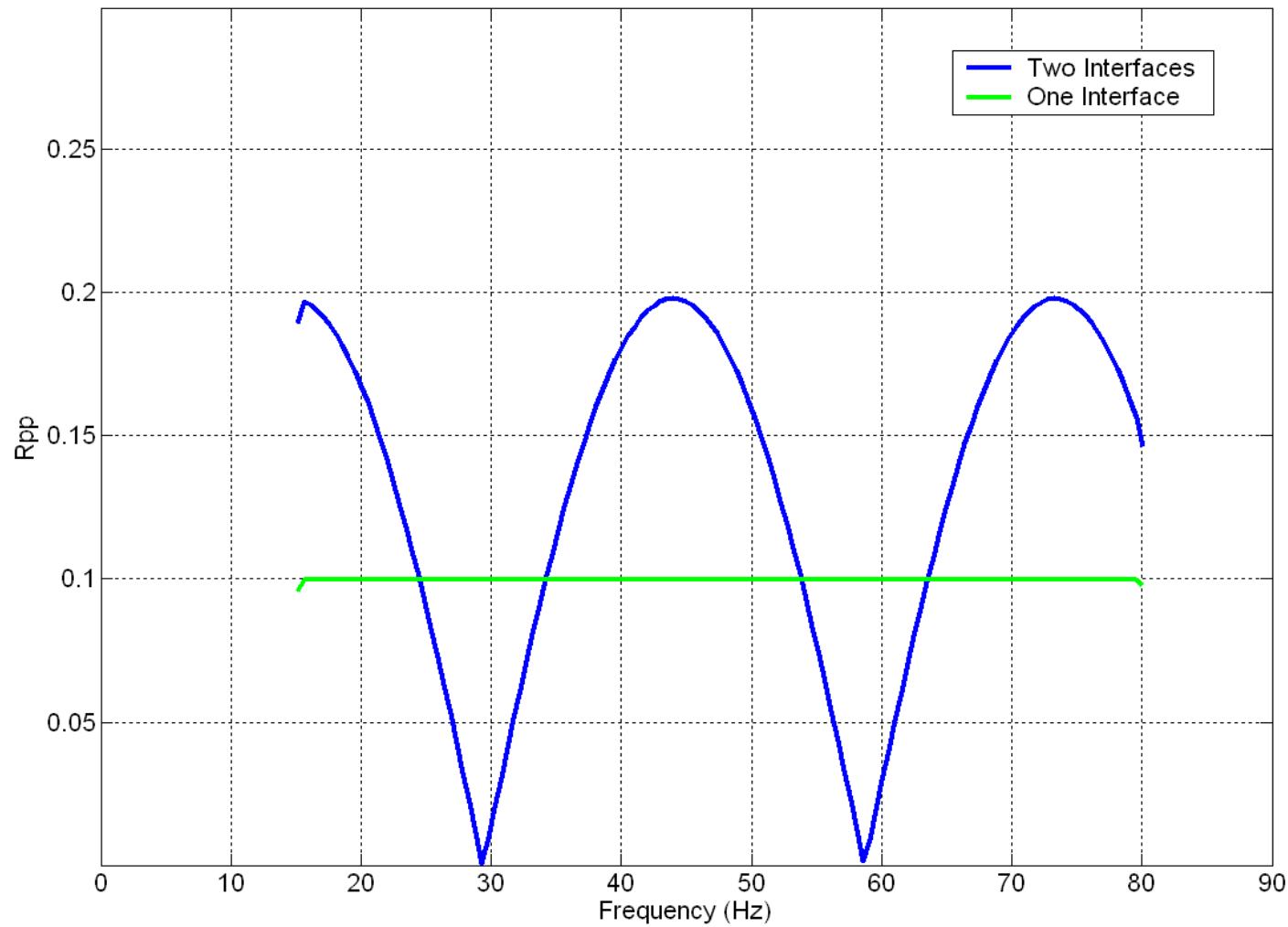
P-Wave Response of a 100 m Reservoir (top at 500 m)



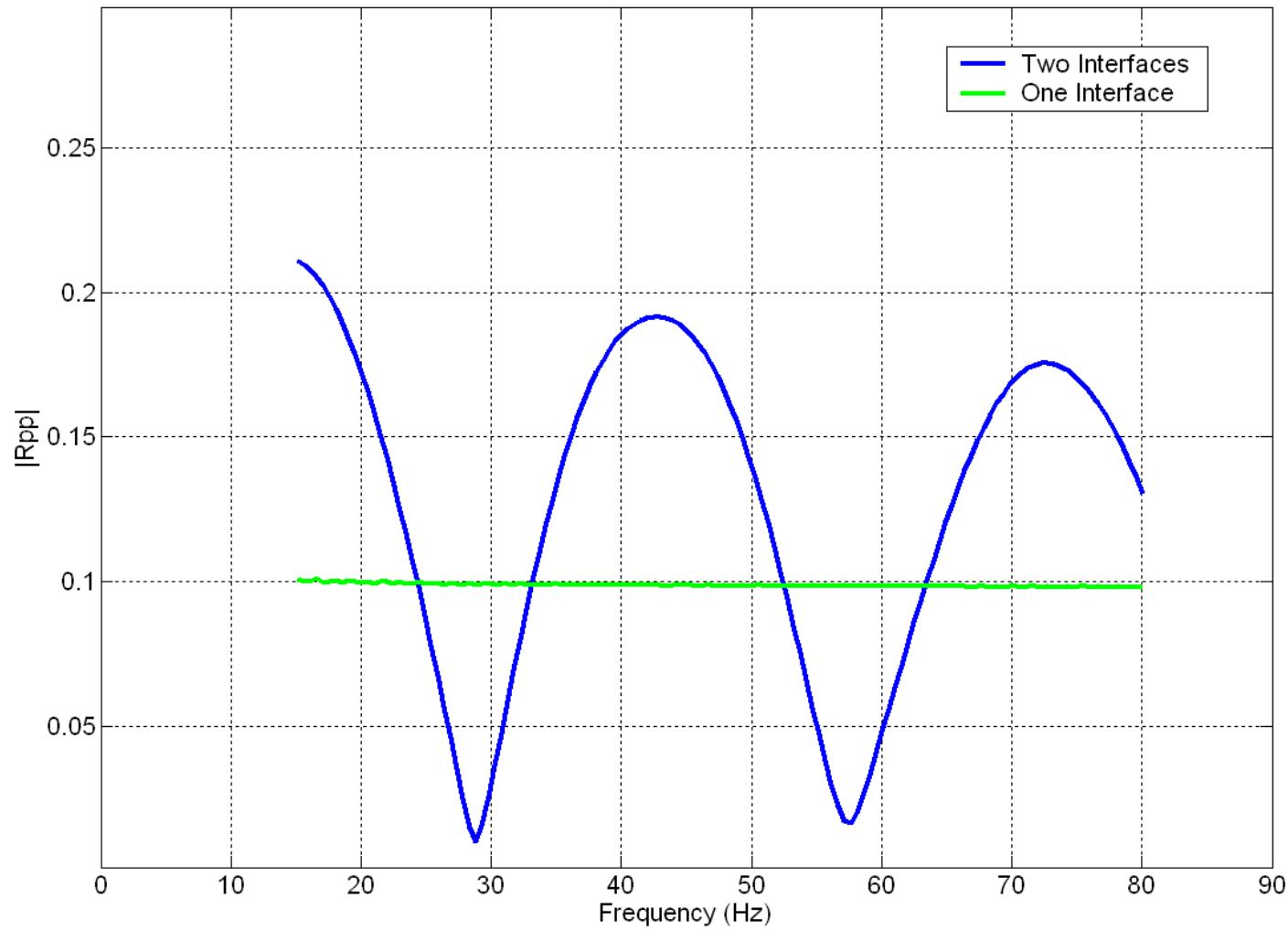
Plane-Wave $R_{pp}(\omega)$ (vertical incidence, 100 m reservoir)



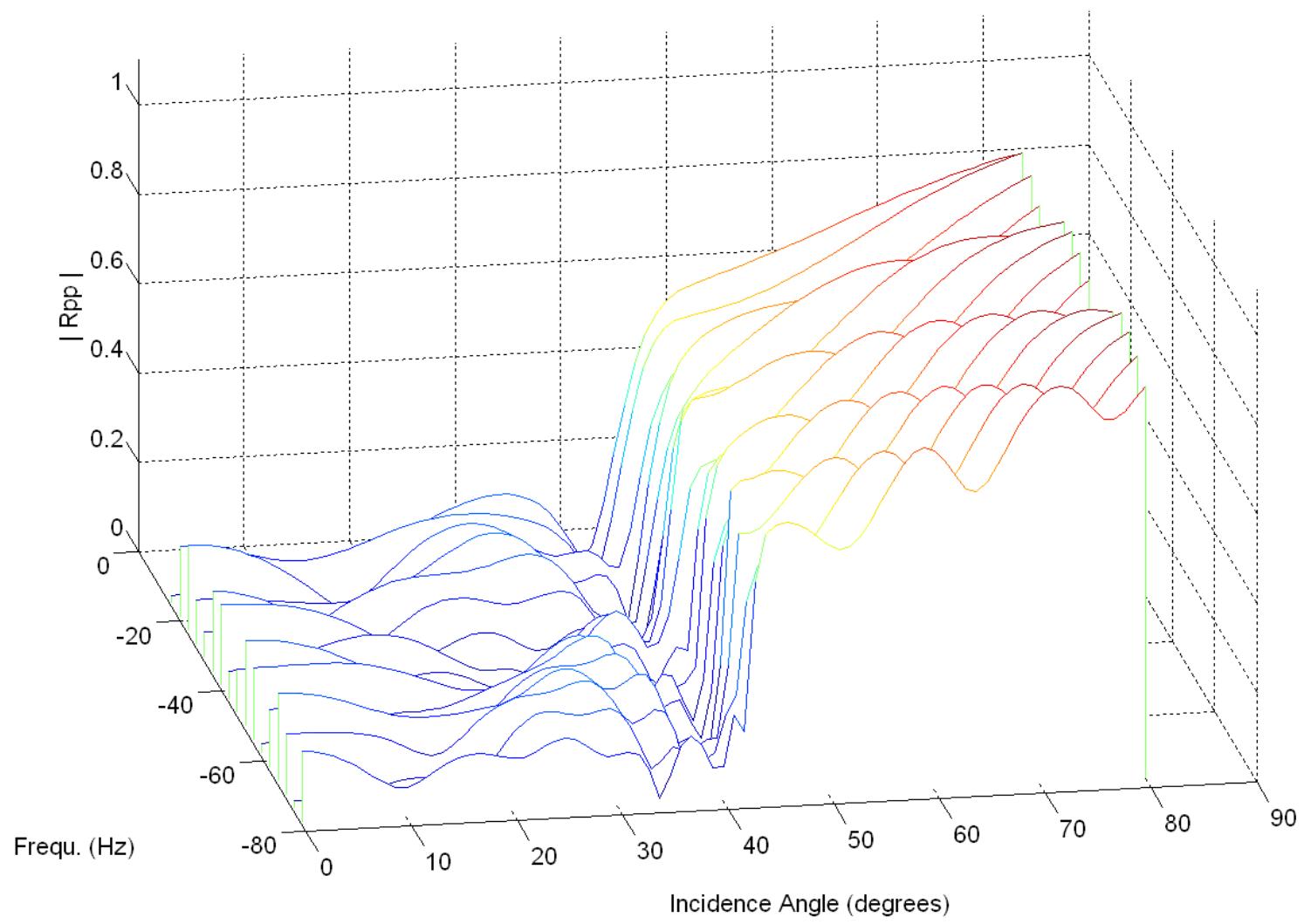
Spher.-Wave $R_{pp}(\omega)$ (vertical incidence, 100 m reservoir)



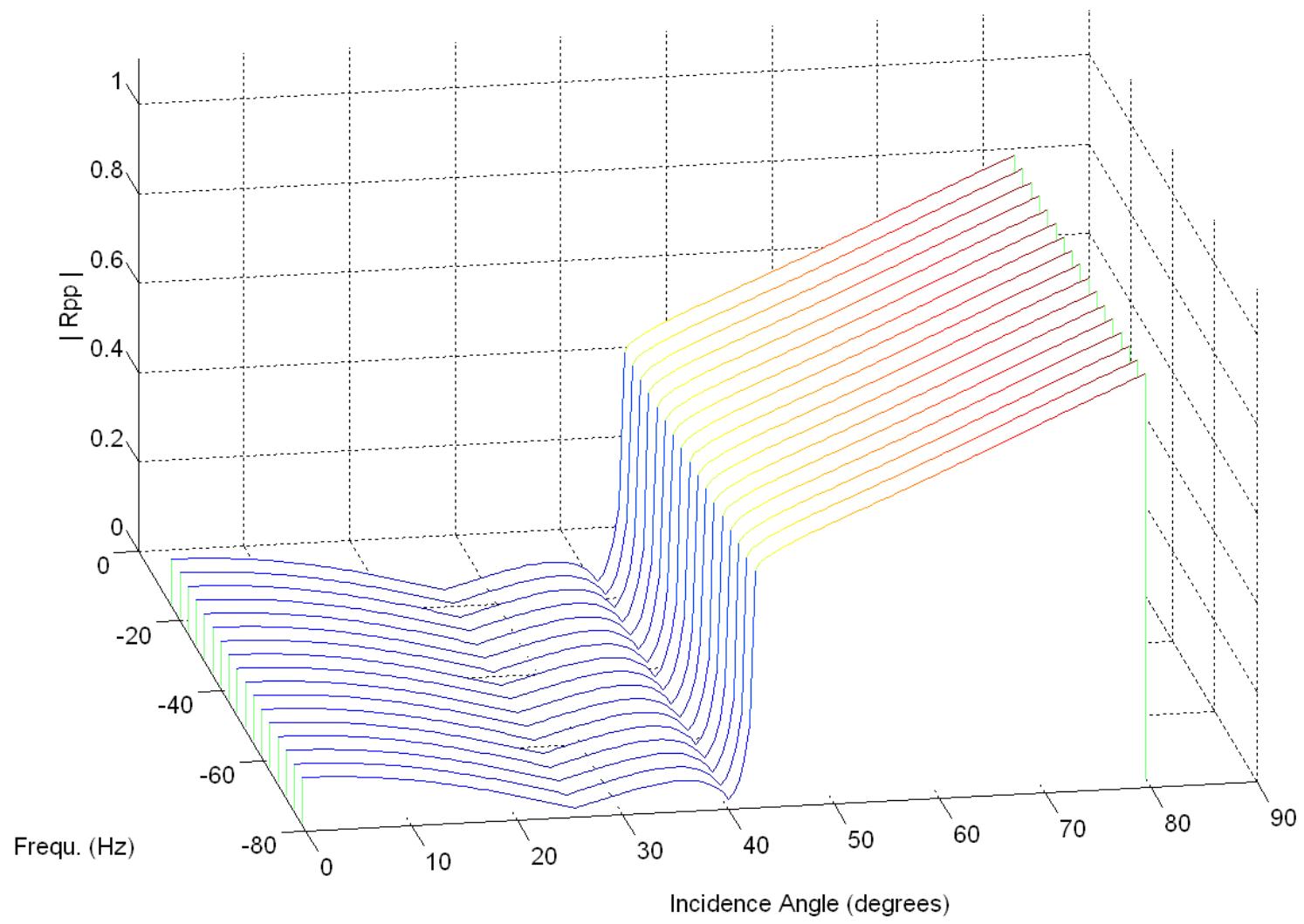
Plane-Wave $R_{pp}(\omega)$ (vertical incidence, 50 m reservoir)



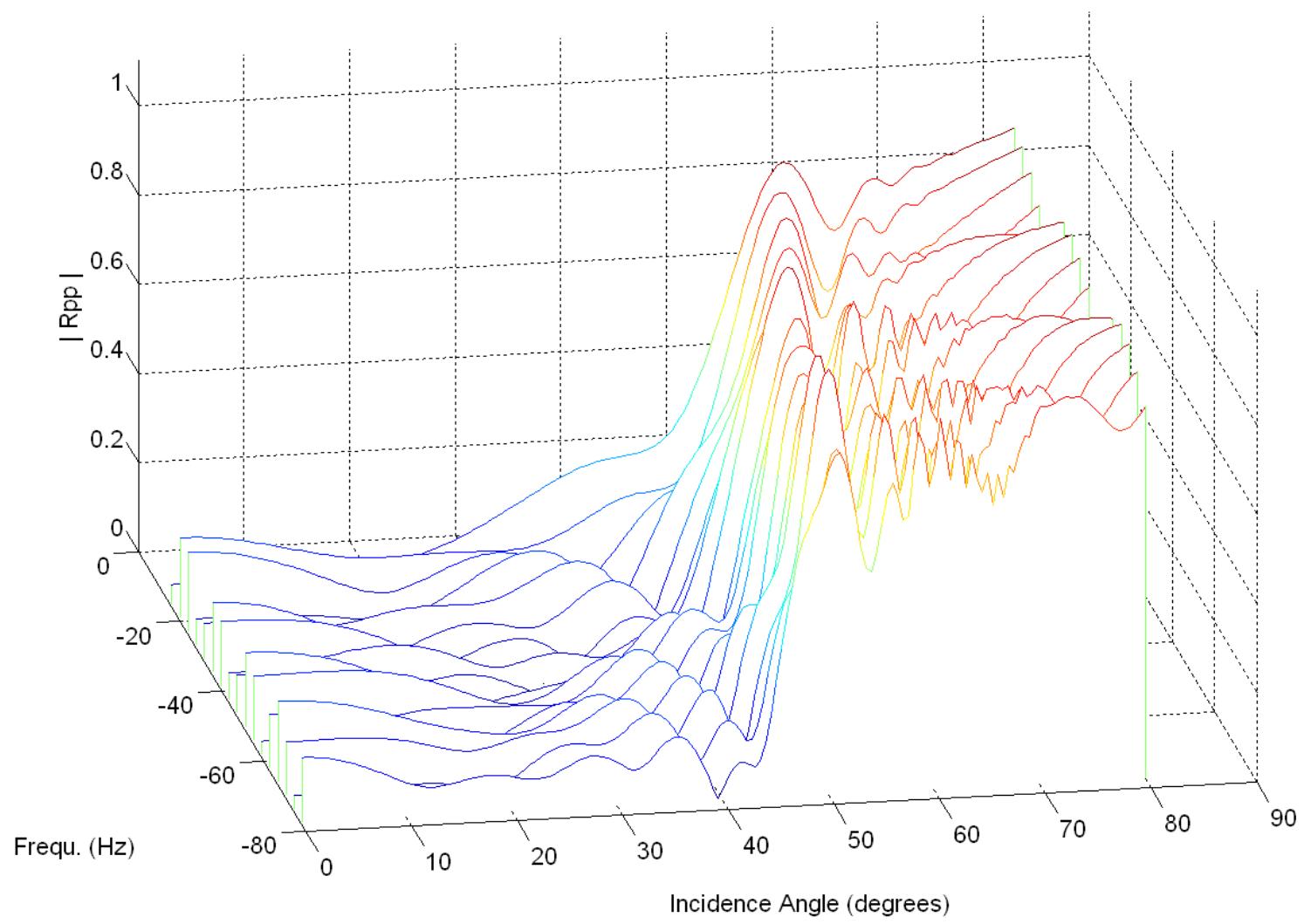
Spher.-Wave $R_{pp}(\omega)$ (vertical incidence, 50 m reservoir)



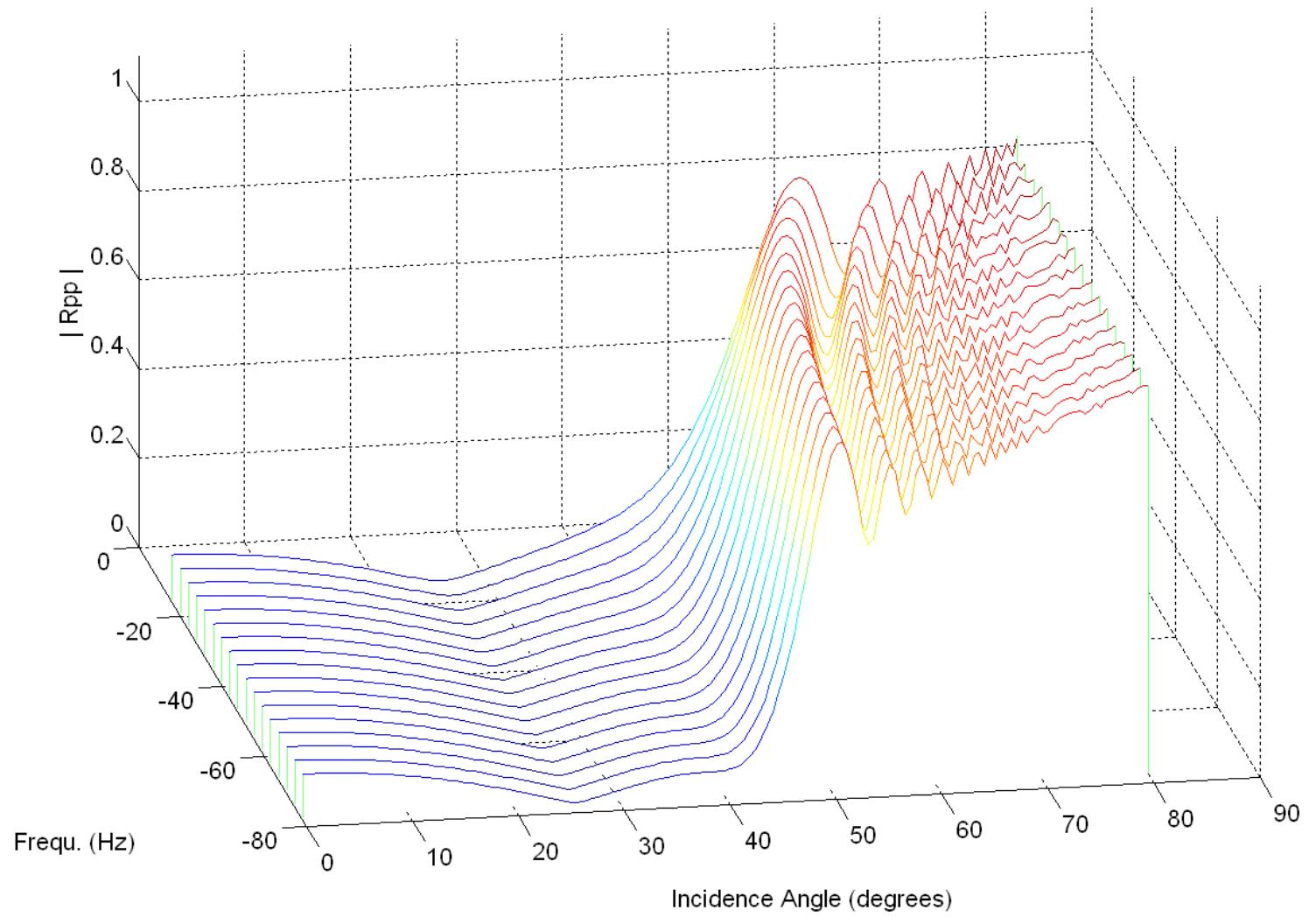
Plane-Wave $R_{pp}(\omega)$ (100 m reservoir)



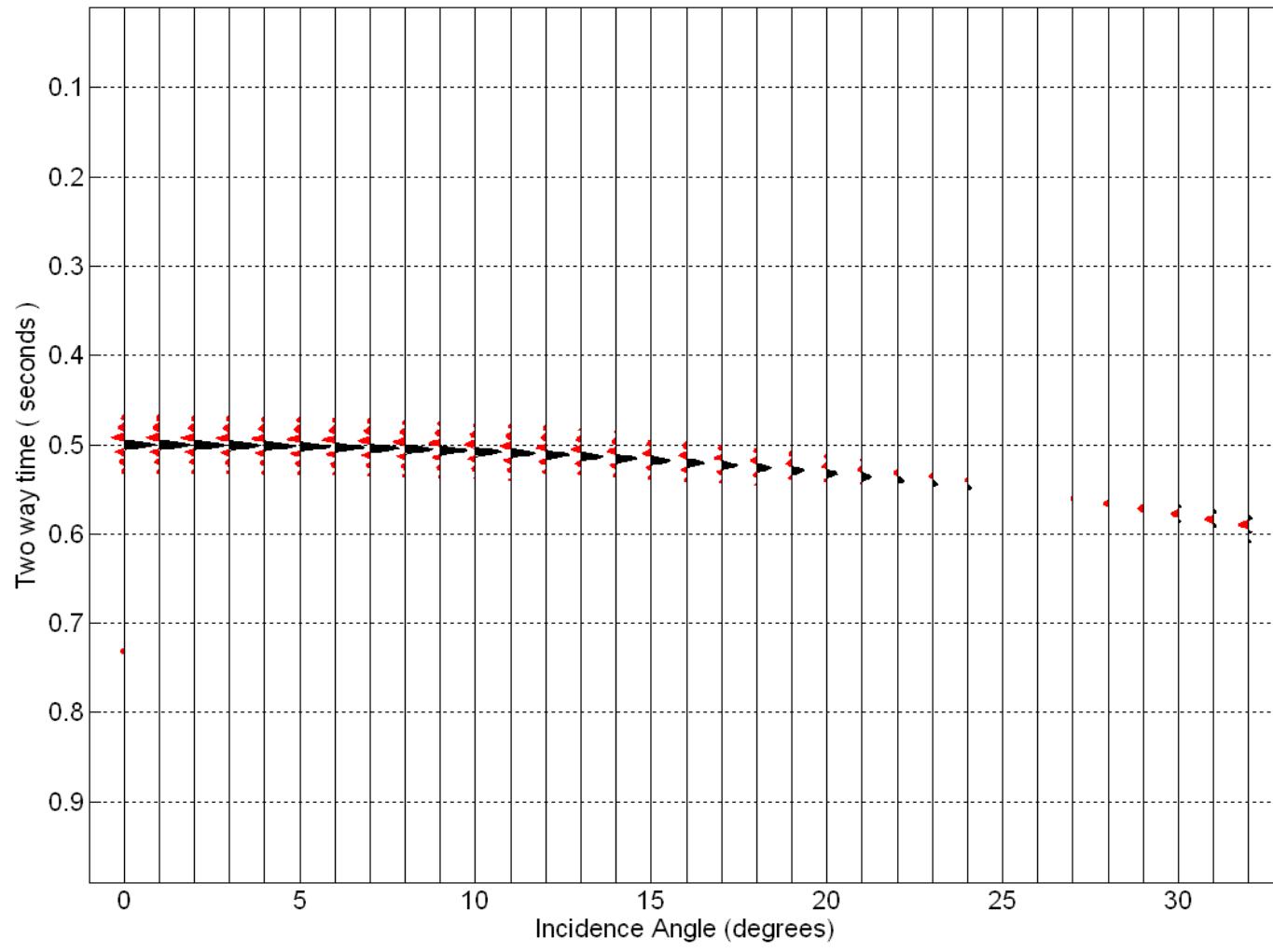
Plane-Wave $R_{pp}(\omega)$ (single interface)



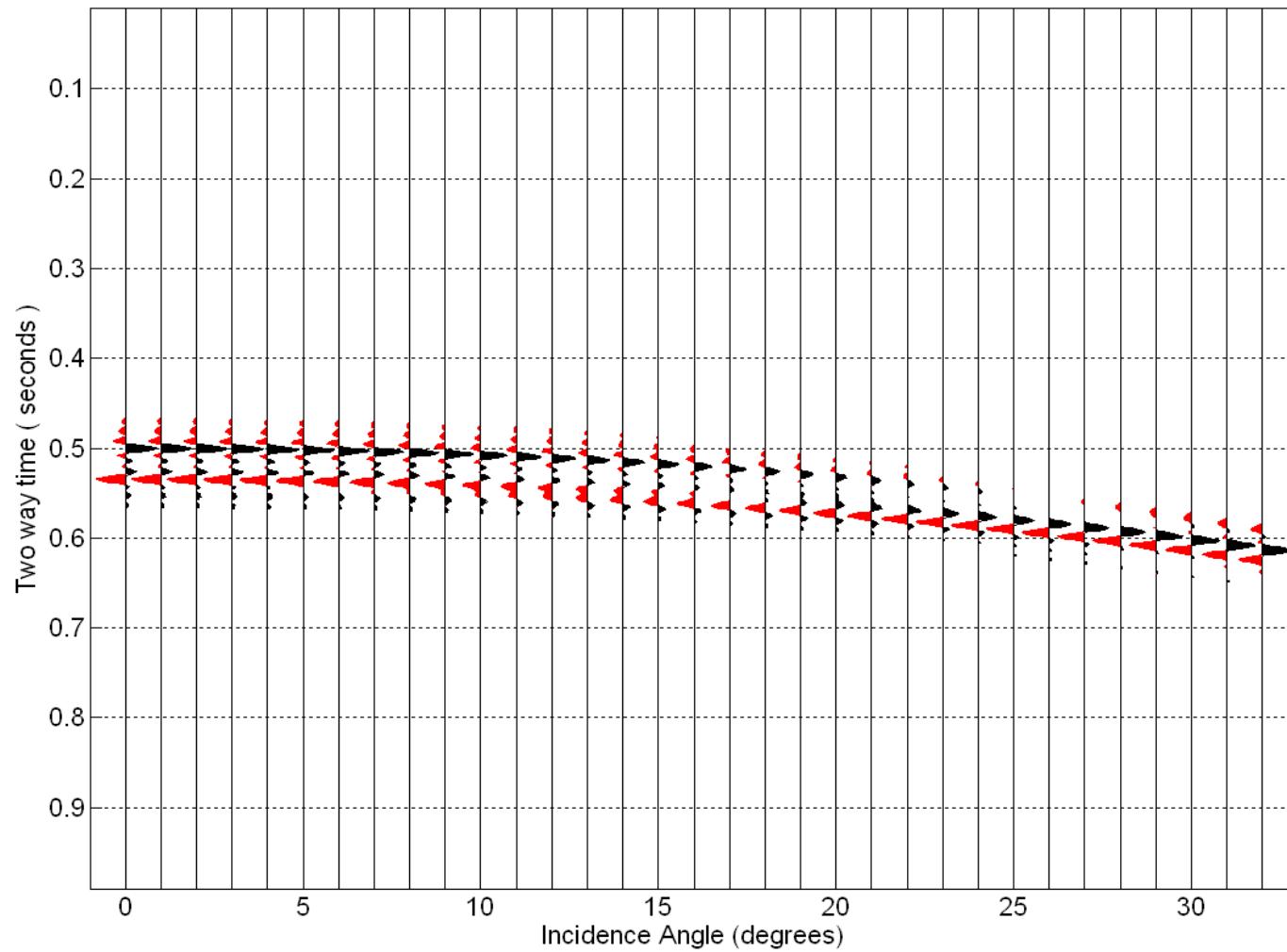
Spherical-Wave $R_{pp}(\omega)$ (100 m reservoir)



Spherical-Wave $R_{pp}(\omega)$ (single interface)



Spherical-Wave PP-Response (single reflector at 500m)



Spherical-Wave PP-Response (50 m reservoir)

Conclusions

- ◆ Reverberations and spherical spreading are modelled by the Ewing-method.
- ◆ Small-offset Class 1 AVO-responses show reservoir layer reverberations in the frequency-domain but at larger offsets the response is obscured.
- ◆ Time-domain responses show an amplitude build-up because of far-offset tuning.

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Thank you for your attention.