

# **3D anisotropic phase shift operators**

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# Introduction

- Through Rayleigh-Sommerfeld modelling and imaging we can:
  - Control up  $\uparrow$  and down  $\downarrow$  and mode conversion:
  - Cause reflection and transmission.
  - Ensure high frequency stability and reduced runtime.
- Extrapolate wavefields with phaseshift operators one planewave at a time.
- Handle heterogeneous media with Gabor windows [Margrave et al., 2002].
- Anisotropy?

## Phaseshift operators in TTI media

- Phaseshift operators are defined relative to the recording surface.
- TTI media are defined relative to the axis of symmetry.
- To reconcile, build operators in the TI coordinate system, map to the data coordinates, then apply.

## Planewave vs. angle, parameters vs. coefficients

- Monochromatic spectrum  $\varphi$  is phaseshifted to a new depth  $\Delta z$  according to

$$\varphi_{\Delta z} = \varphi e^{i\omega q \Delta z}. \quad (1)$$

- In TI media  $q$  is [Daley and Hron, 1977]:

$$q \rightarrow q(p_I, C_{11}, C_{12}, C_{13}, C_{33}, C_{44}). \quad (2)$$

- Alternatively,  $q$  is [Thomsen, 1986]:

$$q \rightarrow q(\theta_I, \alpha_0, \beta_0, \varepsilon, \delta^*, \gamma). \quad (3)$$

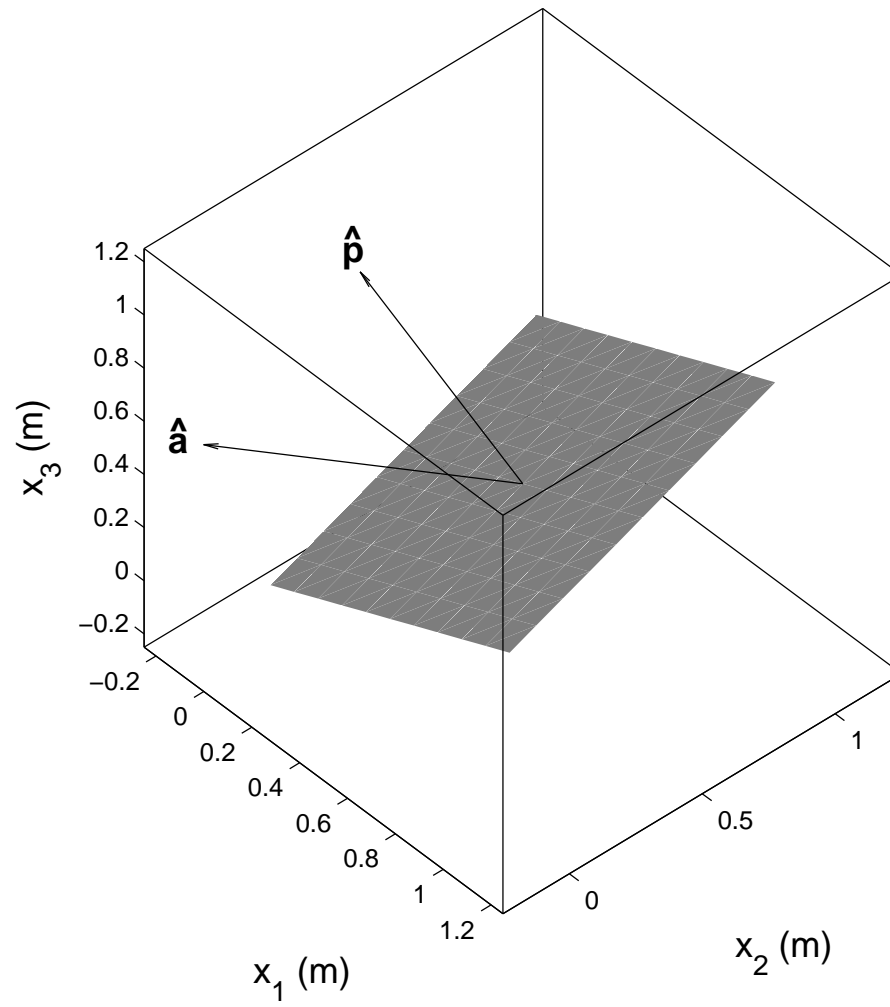
- For RS, we want to use planewaves, but  $\alpha_0, \beta_0, \varepsilon, \delta^*$ , and  $\gamma$  are ubiquitous - reconcile notation [Ferguson and Sen, 2004] so:

$$q \rightarrow q(p_I, \alpha_0, \beta_0, \varepsilon, \delta^*, \gamma). \quad (4)$$

- Matlab's symbolic math toolbox helps.

## What is $p_I$ in a TTI medium?

- For TTI,  $p_I = \frac{\sin \theta_I}{v}$  where  $\theta_I$  is the angle between planewave unit normal  $\hat{\mathbf{p}}$  and normal  $\hat{\mathbf{a}}$  to the symmetry plane.



- Use Snell's Law to eliminate raytracing:

$$p_I = \frac{\sin \theta_I}{v} = |\hat{\mathbf{p}} \times \hat{\mathbf{a}}| |\hat{\mathbf{p}}|. \quad (5)$$

- Recall,  $p_1 = \frac{k_x}{\omega}$ ,  $p_2 = \frac{k_y}{\omega}$ , and  $p_3 = \frac{1}{v} \sqrt{1 - (v p_1)^2 - (v p_2)^2}$ , so given  $\alpha_0$ ,  $\beta_0$ ,  $\varepsilon$ ,  $\delta^*$ , and  $\gamma$ ,  $q_\alpha$ ,  $q_{\beta_{SH}}$  and  $q_{\beta_{SV}}$  are computable for TTI media.
- Note,  $q$  are referenced to  $\hat{\mathbf{a}}$  - map  $q(\hat{\mathbf{a}} \rightarrow \hat{\mathbf{p}})$  to extrapolate  $\varphi(\hat{\mathbf{p}})$  so invoke Snell's Law again so that

$$q_r^2 = q^2 + p_I^2 - p_1^2 - p_2^2, \quad (6)$$

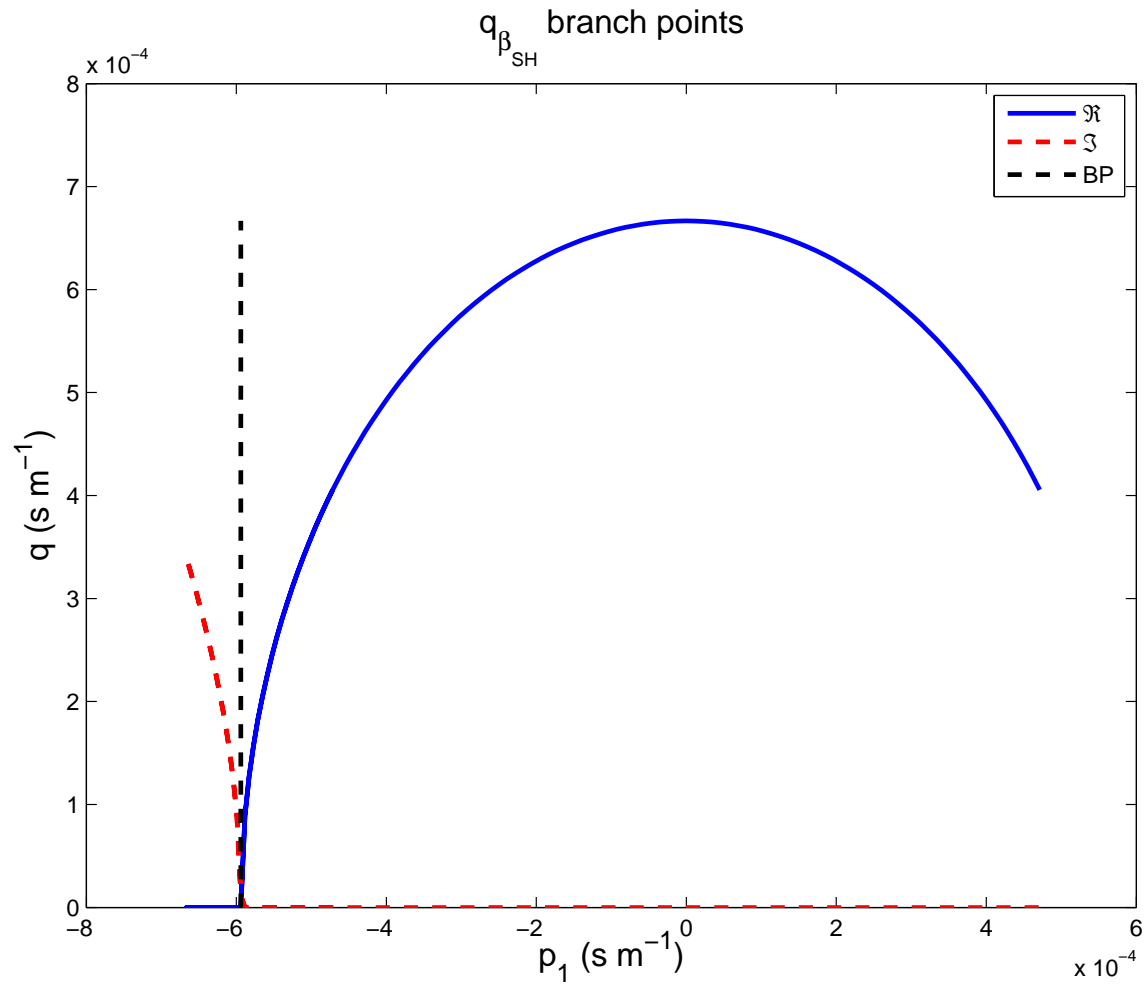
is the general form for the three modes.



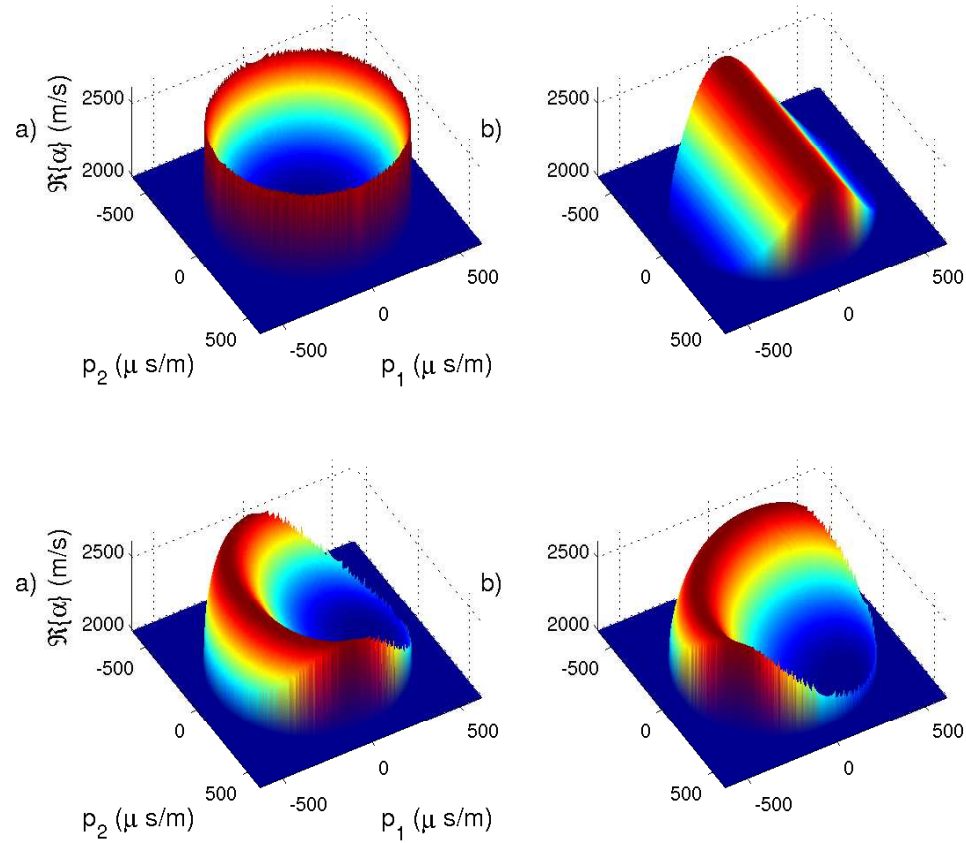
## Branch points

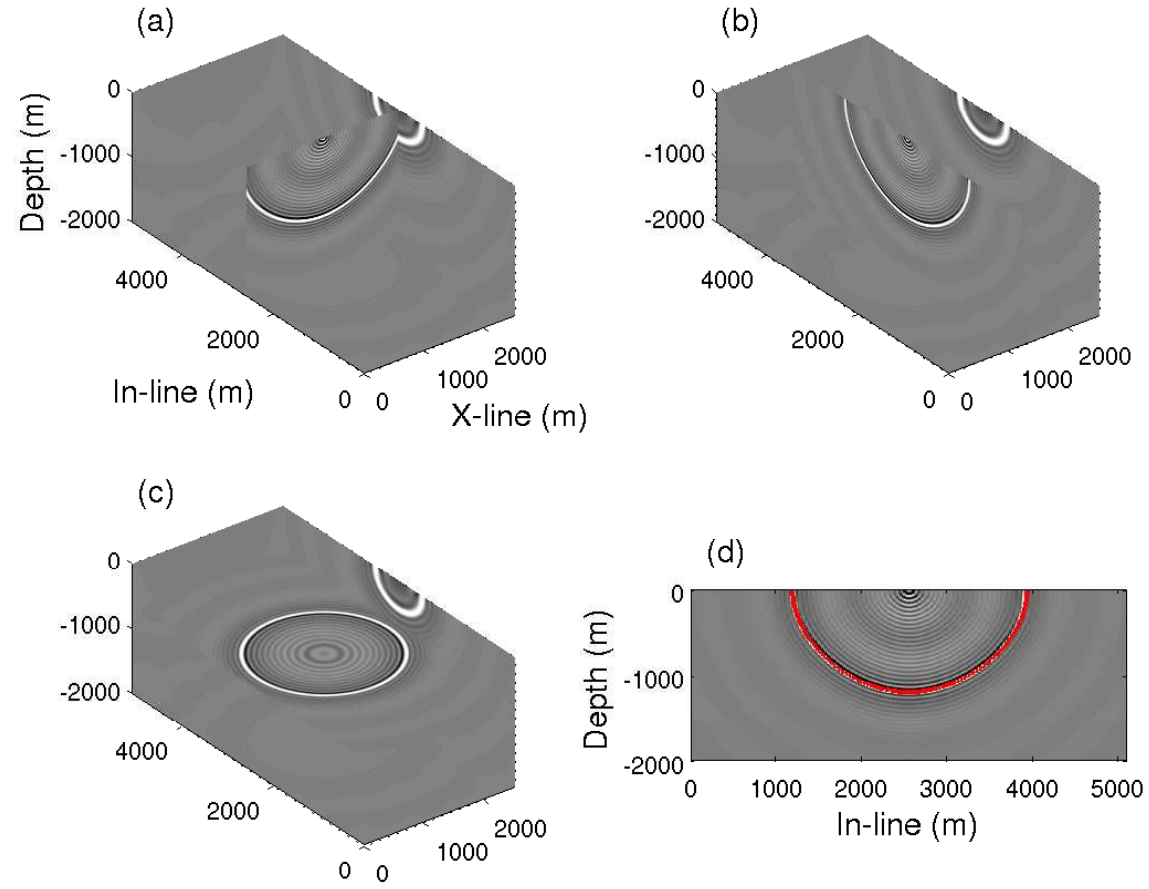
- For VTI and HTI, the evanescent region is defined.
- TTI is more complicated, and branchpoints must be determined, for example, set  $q_{\beta_{SH}} = 0$  to get [Bale, 2006]

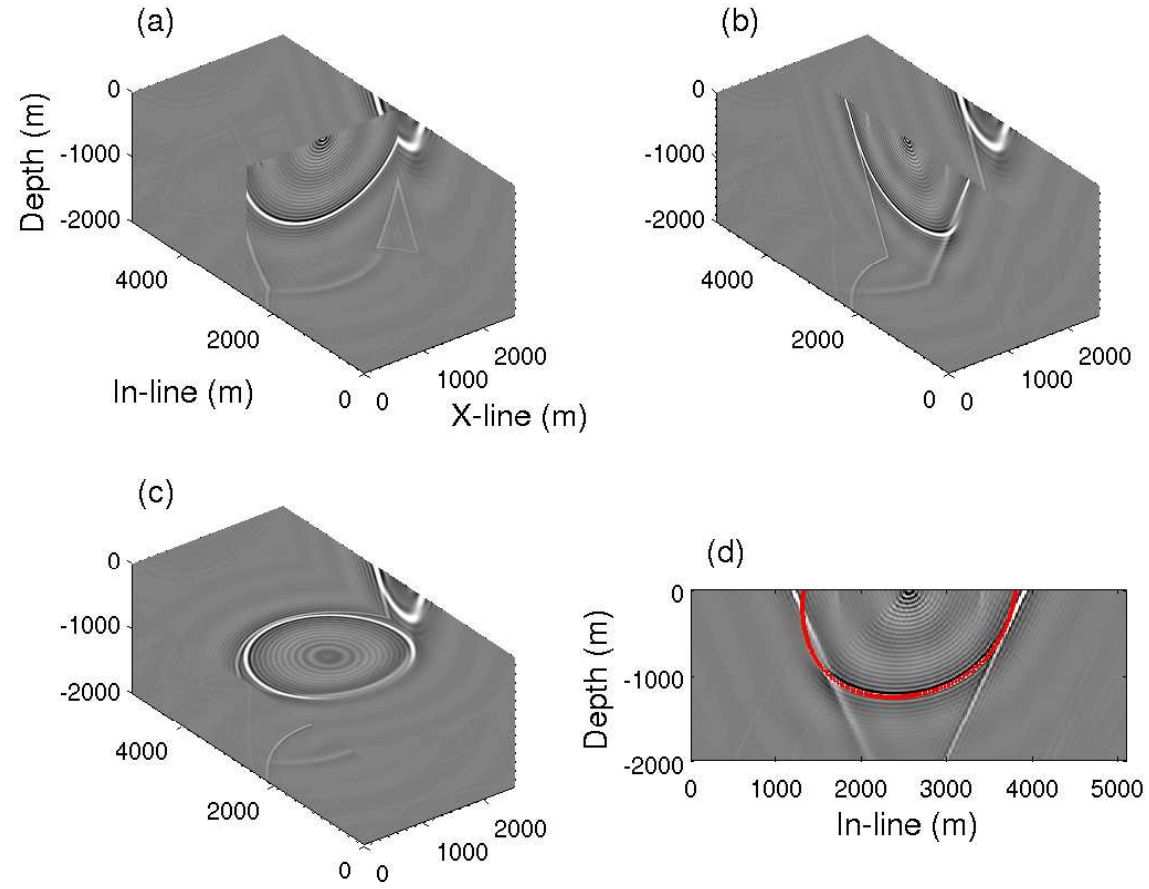
$$1 = \beta_0^2 [p_1^2 + p_2^2] [2\gamma + 1], \quad (7)$$

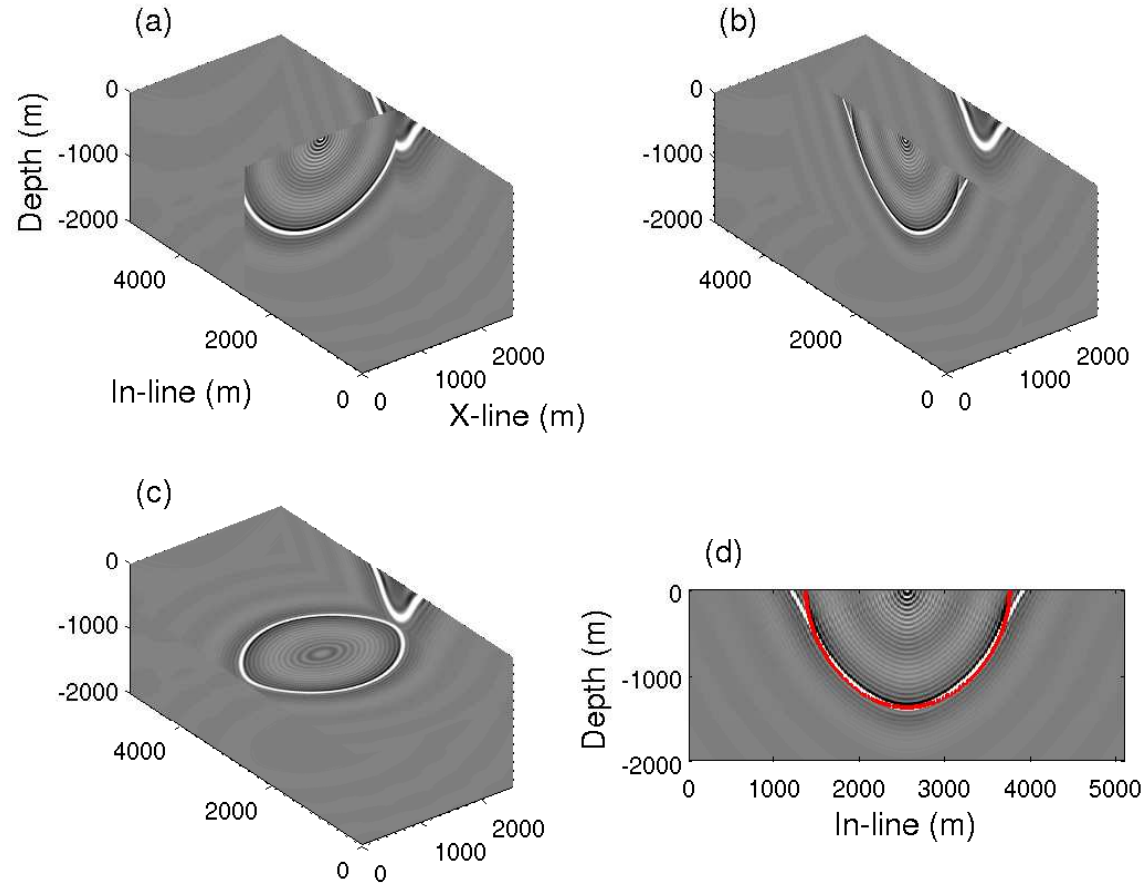


# Example: P-waves

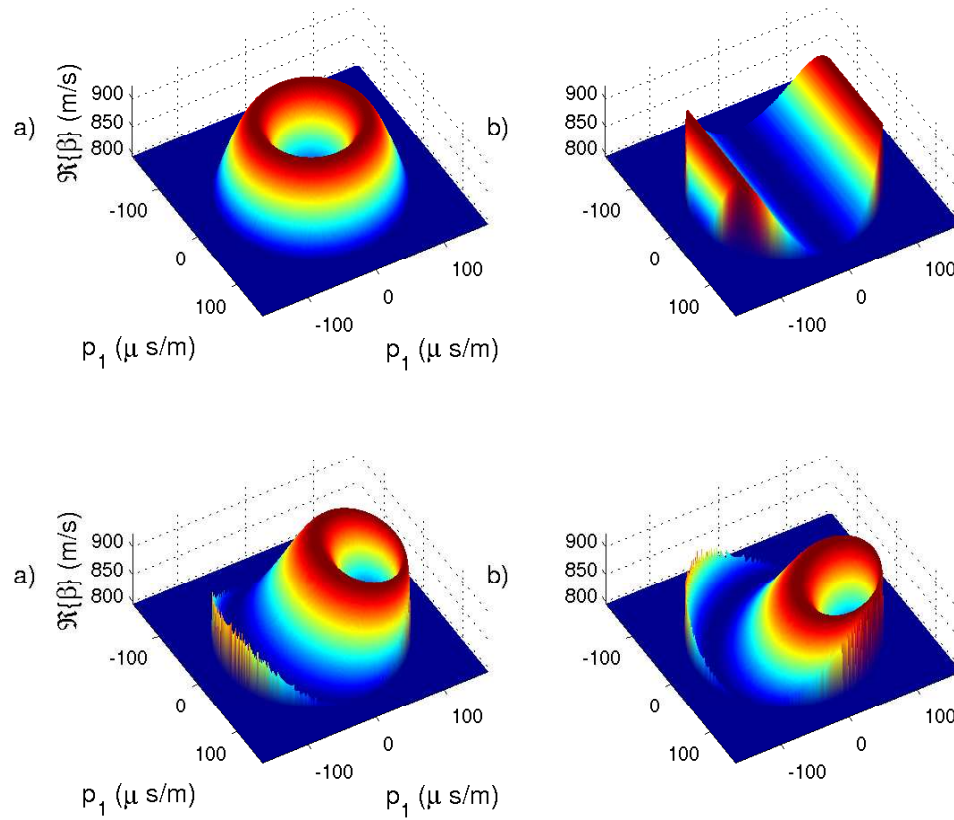


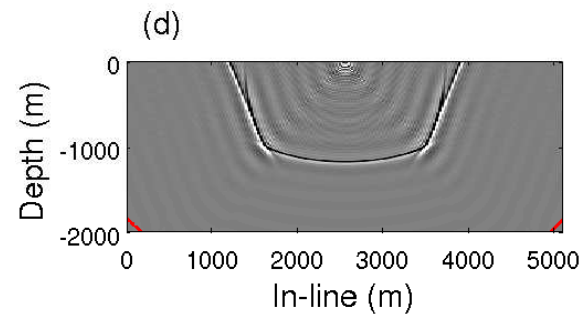
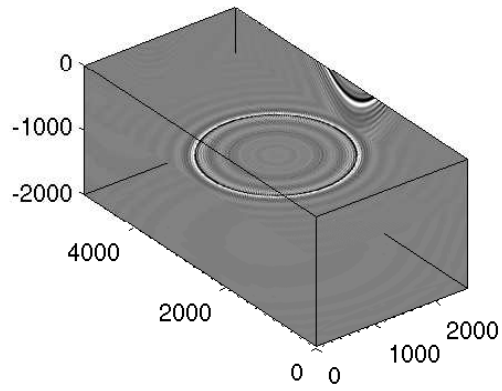
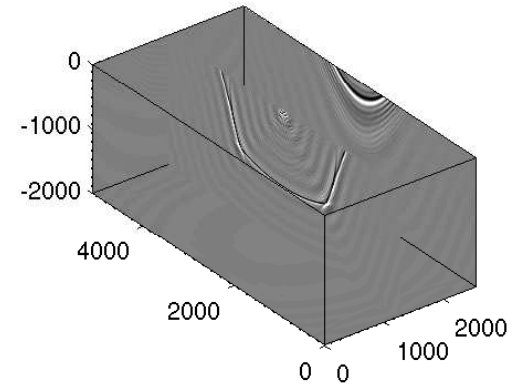
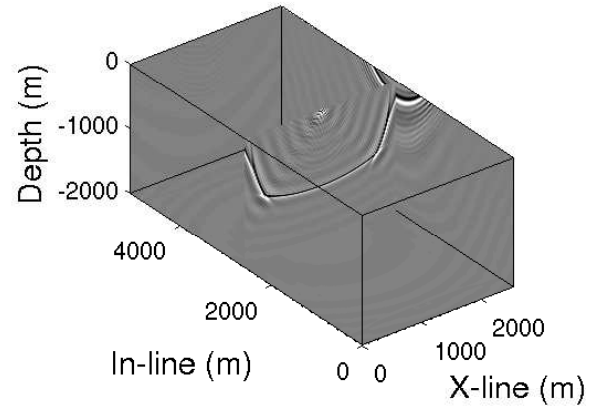




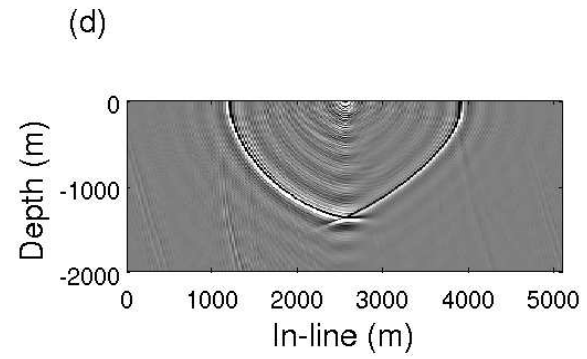
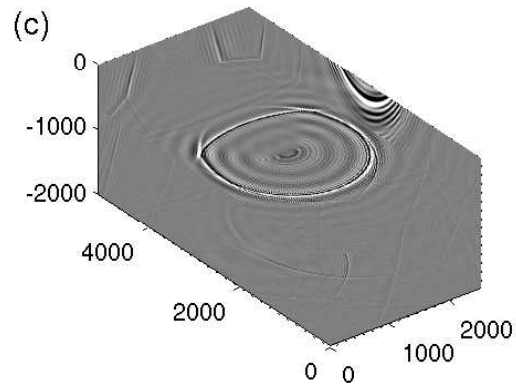
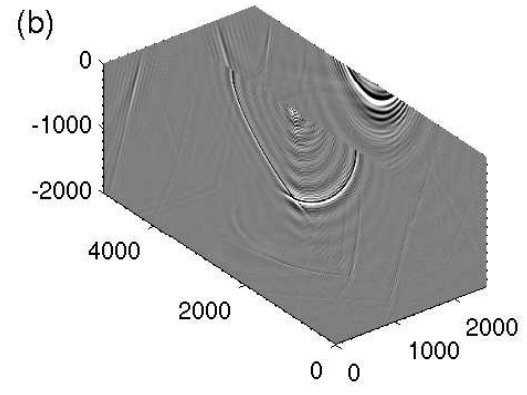
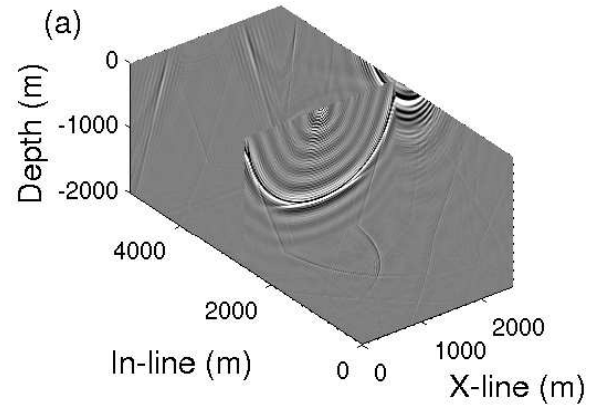


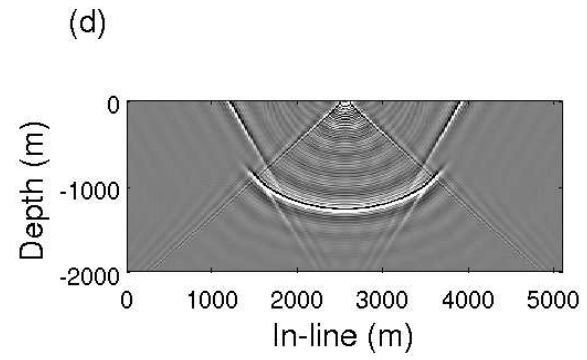
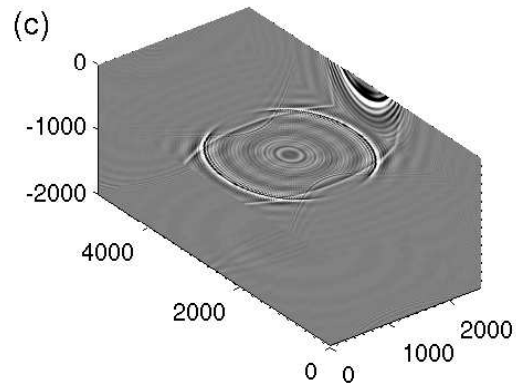
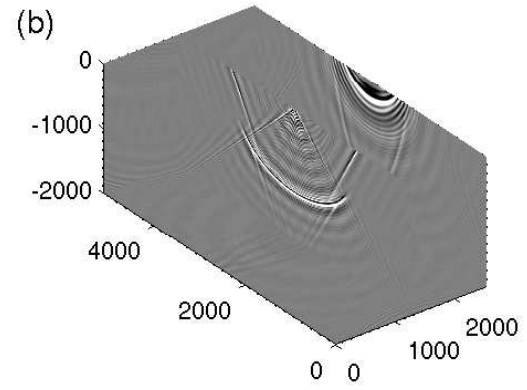
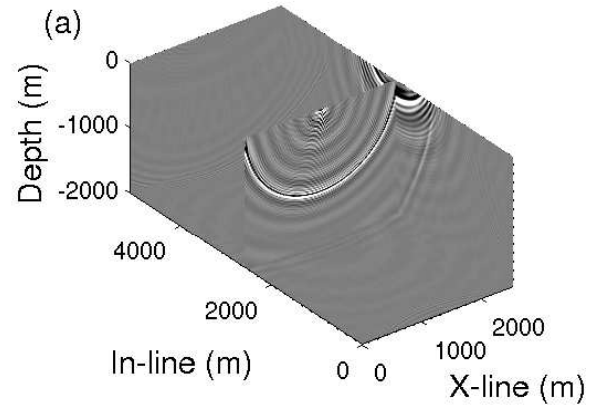
# Example: SV-waves



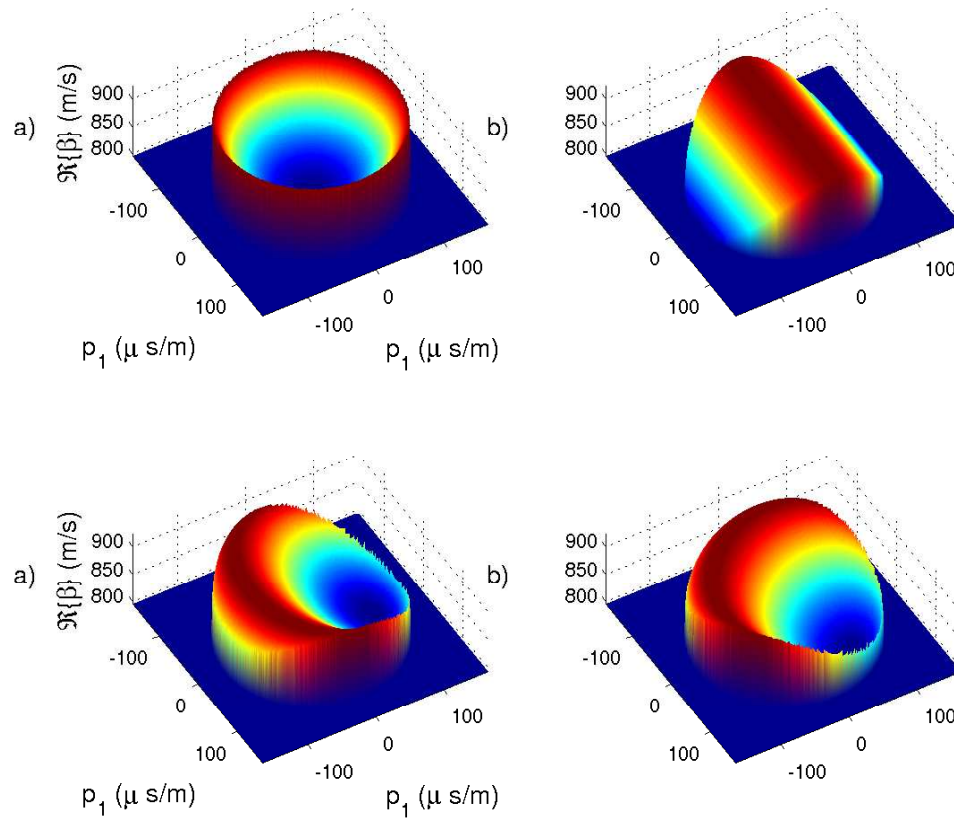








# Example: SH-waves



## Conclusions and future work

- 3D phaseshift operators for homogeneous TTI media:
  - P, SV, and SH.
- Need group direction impulses for SV and SH.
- Compute branch points for P and SV.
- Rather than map  $q_{p_I} \rightarrow q_{\hat{\mathbf{p}}}$  then extrapolate  $\varphi_{\hat{\mathbf{p}}}$ , extrapolate  $\varphi_{p_I}$  then map  $\varphi(p_I) \rightarrow \varphi(\hat{\mathbf{p}})$ .

# Acknowledgements

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## References

- [Bale, 2006] Bale, R. A., 2006, Elastic wave-equation depth migration of seismic data for isotropic and azimuthally anisotropic media: PhD thesis, University of Calgary.
- [Daley and Hron, 1977] Daley, P. F. and F. Hron, 1977, Reflection and transmission coefficients for transversely isotropic media: Bulletin of the Seismological Society of America, **56**, 87–94.
- [Ferguson and Sen, 2004] Ferguson, R. J. and M. K. Sen, 2004, Estimating the elastic parameters of anisotropic media using a joint inversion of p-wave and sv-wave traveltimes: Geophysical Prospecting, **52**, 547–558.
- [Margrave et al., 2002] Margrave, G., M. Lamoureaux, J. Grossman, and V. Iliescu, 2002, Gabor deconvolution of seismic data for source waveform

and Q correction: 72nd Ann. Internat. Mtg, 2190–2193, Soc. of Expl. Geophys.

[Thomsen, 1986] Thomsen, L., 1986, Weak elastic anisotropy: Geophysics, **51**, 1954–1966. Discussion in GEO-53-04-0558-0560 with reply by author.