



UNIVERSITY OF  
CALGARY



# FLUID SUBSTITUTION AND SEISMIC MODELLING IN A SANDSTONE AQUIFER

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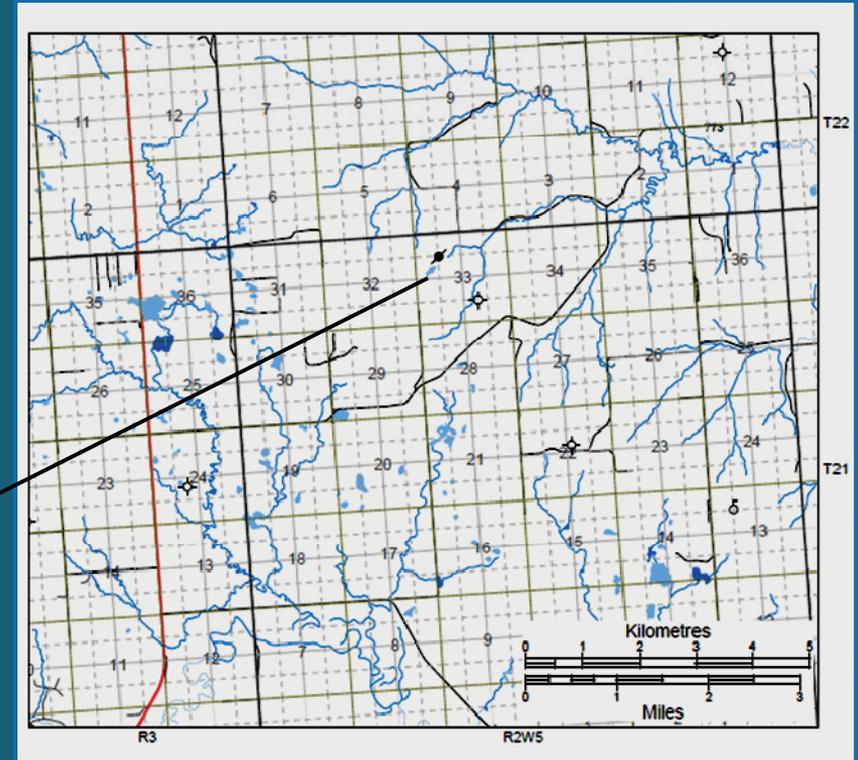
# Outline

- **Objectives**
- **Area of Study and Stratigraphy**
- **Fluid Substitution**
- **Data and Methodology**
- **Results and Conclusions**

# Objectives

- To evaluate the Paskapoo Formation as a potential CO<sub>2</sub> geological storage site.
- To apply fluid substitution and seismic modelling in order to identify and analyze the effects of CO<sub>2</sub> on rock properties and seismic patterns.

# Area of Study



Wellbore Data: MILLAR 12-33-21-2W5

Rocky Mountains  
Foothills, 20 kilometres  
Southwest of Calgary

(Bachu et al., 2000)

# Stratigraphy

ERA	PERIOD	FORMATION OR GROUP	LITHO.	AVERAGE SEISMIC VELOCITIES (m/s)	
MESOZOIC	Tertiary (TERT)	Paskapoo		3200-3700	
	Cretaceous	Upper Brazeau	Edmonton (EDMN)		
			Bearpaw Shale		
		Lower Brazeau	Belly River (BLRV)		3800-4100
			Wapiabi		
		Alberta Group	Cardium (CRDM)		3900-4150
			Blackstone		
	Lower	Blairmore (BMGP)		4200-4400	
		Kootenay			
	PALEOZOIC (PAL)	Jurassic	Fernie		4100-4300
Mississippian (MSSP)			Mount Head		
			Turner Valley		
		Shunda			
		Pekisko			
Devonian (DEVN)		Banff			
		Exshaw			
Palliser					
Cambrian	Fairholme		5500-6000		
				Detachment Horizon ←	

## Paskapoo Fm:

- Composed of mudstone, siltstone and sandstone, with subordinate limestone and coal.

## Foreland Deposits

- Important ground water reservoir target
- High porosity coarse-grained sandstone channels

# Fluid Substitution

## Gassmann equation

$K$  = Bulk Modulus

$$K = \frac{\delta \epsilon}{\delta \sigma}$$

Stress increment

Volume strain

$K_{sat}$  = Saturated Rock Bulk Modulus

$$K_{sat} = K * \frac{(1 - \frac{K*}{K_o})^2}{\frac{\varphi}{K_{fl}} + \frac{1 - \varphi}{K_o} - \frac{K*}{K_o^2}}$$

Matrix Properties



$K_0$  = Matrix Bulk Modulus

$\rho_0$  = Matrix Density

Fluid Properties



$K_{fl}$  = Fluid Bulk Modulus

$\rho_{fl}$  = Fluid Density

Rock Properties



$K_{100sat}$  = Rock Bulk Modulus 100% water sat

$K^*$  = Dry Rock Bulk Modulus

$G$  = Shear Modulus

$\rho_b$  = Bulk Density

$\phi$  = Porosity

$$\phi = \frac{\rho_0 - \rho_b}{\rho_0 - \rho_{fl}}$$

# Fluid Substitution

## Gassmann equation

Constantes during substitution:  
 $\varphi, G, \rho_0, K_0, K^*$

Variables during substitution:  
 $K_{fl}, \rho_{fl}, \rho_b$

$$K_{sat} = K^* \frac{\left(1 - \frac{K^*}{K_0}\right)^2}{\frac{\varphi}{K_{fl}} + \frac{1 - \varphi}{K_0} - \frac{K^*}{K_0^2}}$$

$$v_p = \sqrt{\frac{K_{sat} + \frac{4G}{3}}{\rho_b}}$$

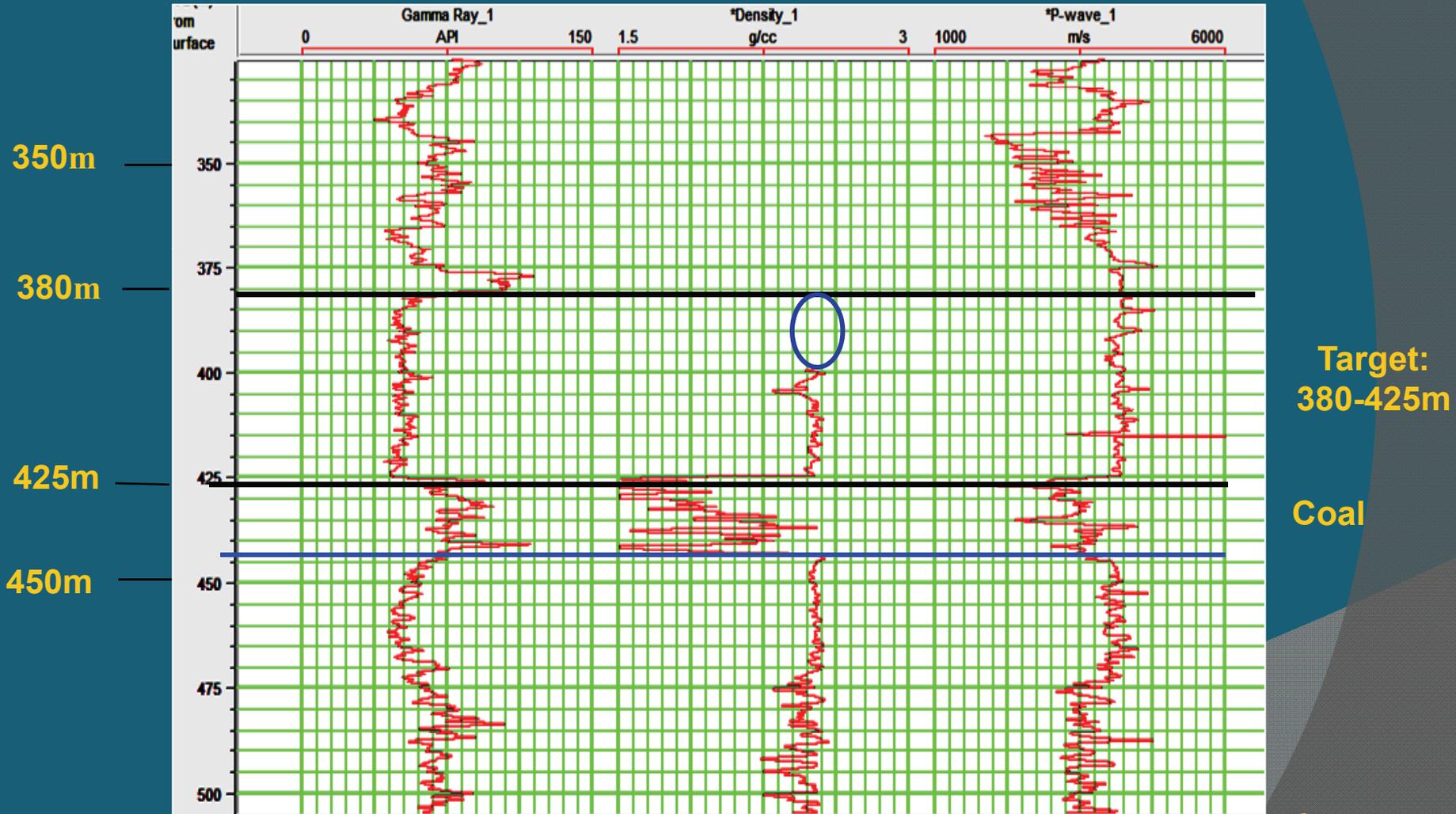
$$v_s = \sqrt{\frac{G}{\rho_b}}$$

## Wellbore data: MILLAR 12-33-21-2W5

GR

Density

P-wave



Target:  
380-425m

Coal

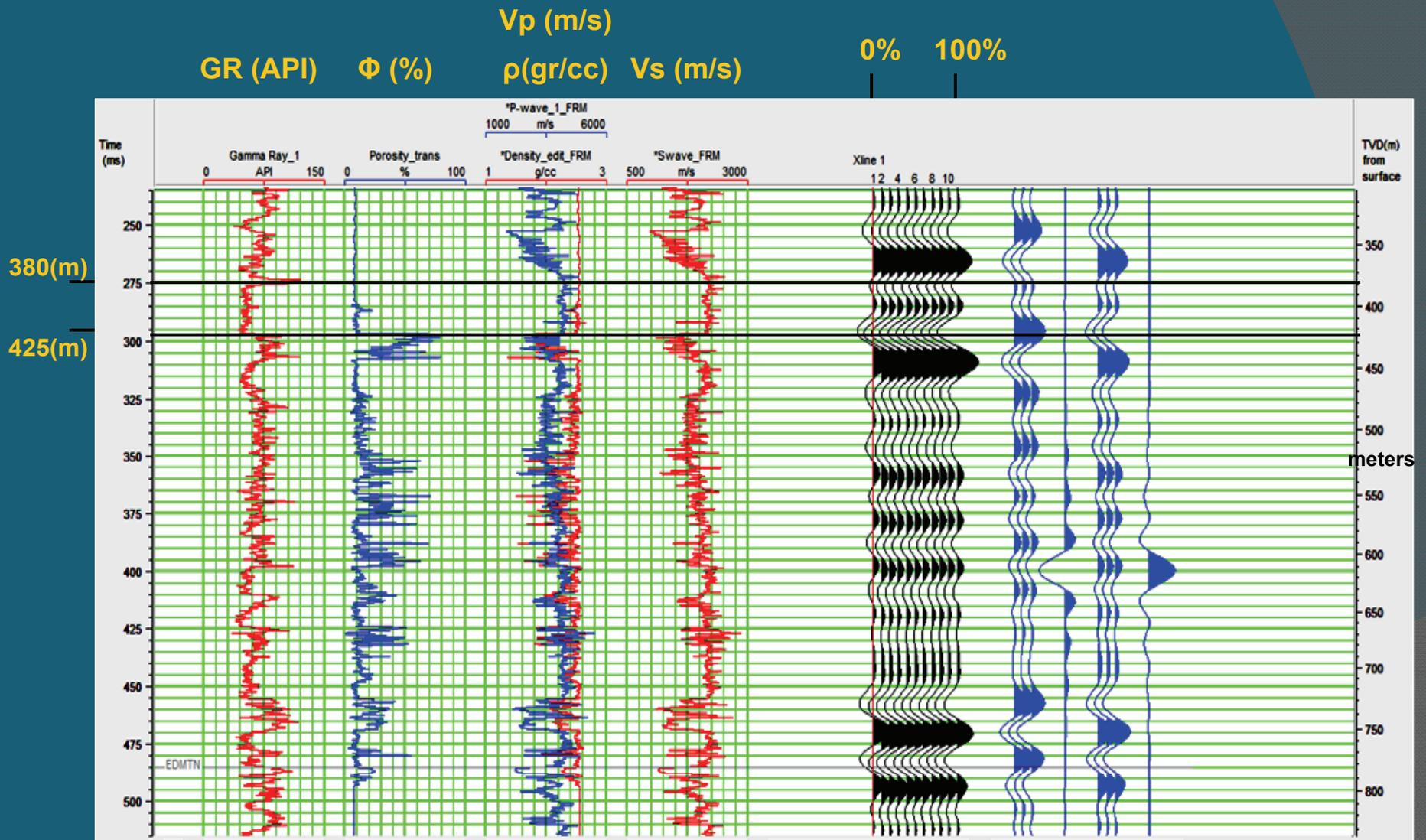
# Methodology

**Fluid substitution using Gassmann's equation and synthetic seismogram generation.**

## **Seismic Modelling :**

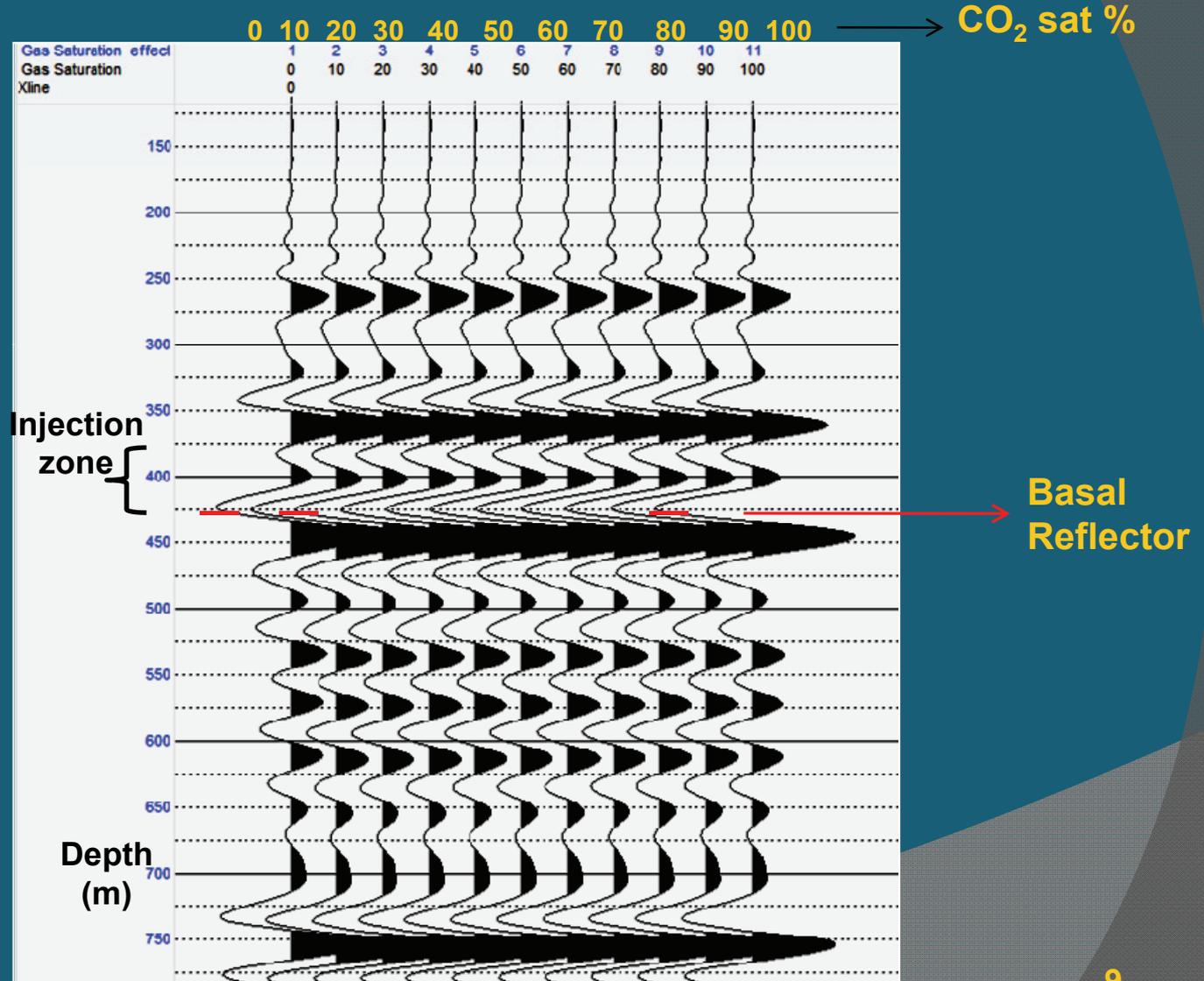
- 2D geological model
- Ray-tracing and synthetic seismogram generation
- Processing of the seismic sections
- Comparison between different seismic sections
- Recognition of differences in seismic pattern and AVO analysis

## Logs and synthetic traces

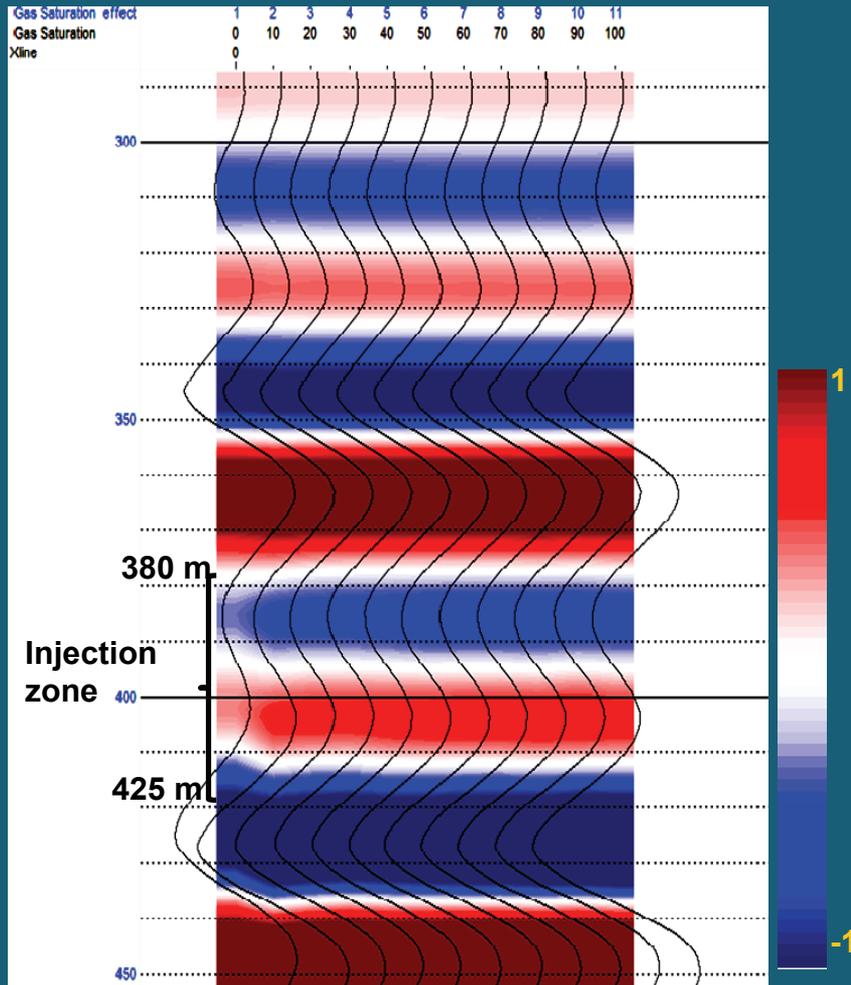


# Results

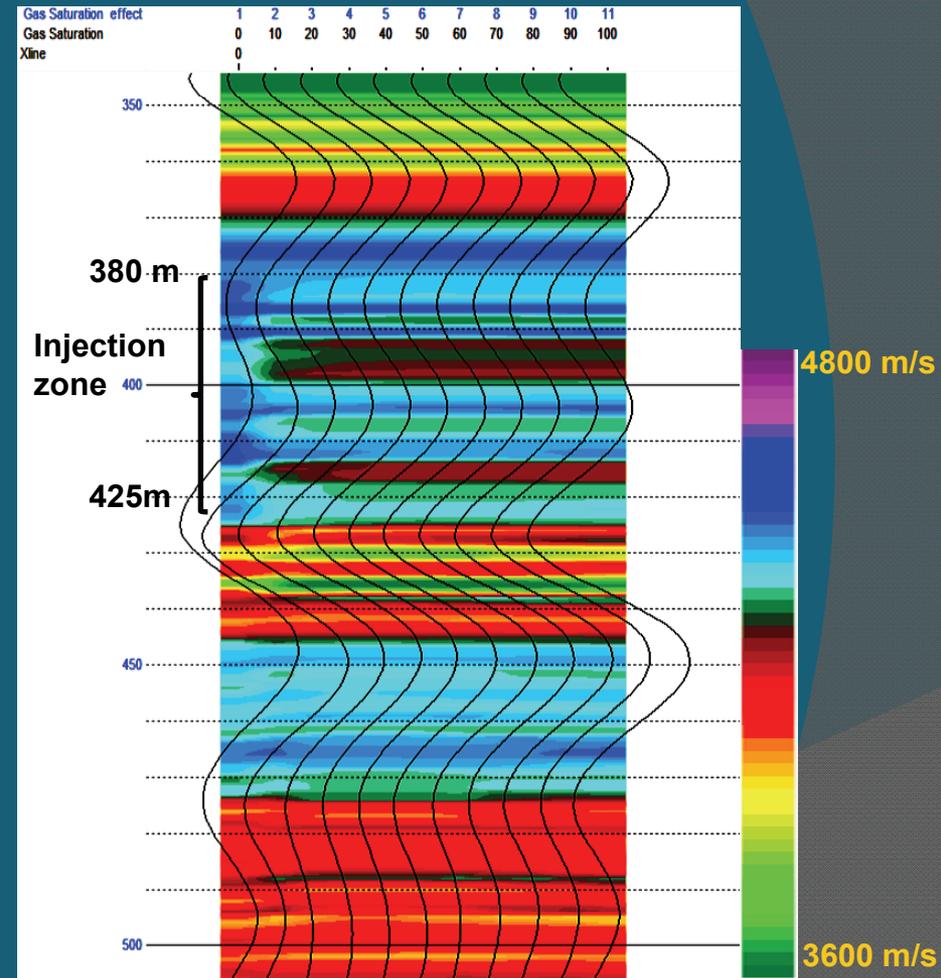
## Synthetic Traces. CO<sub>2</sub> Saturation



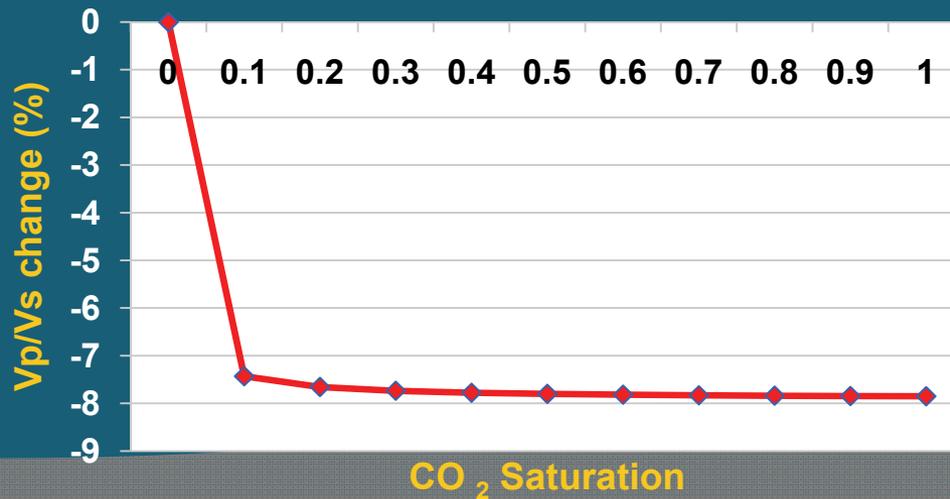
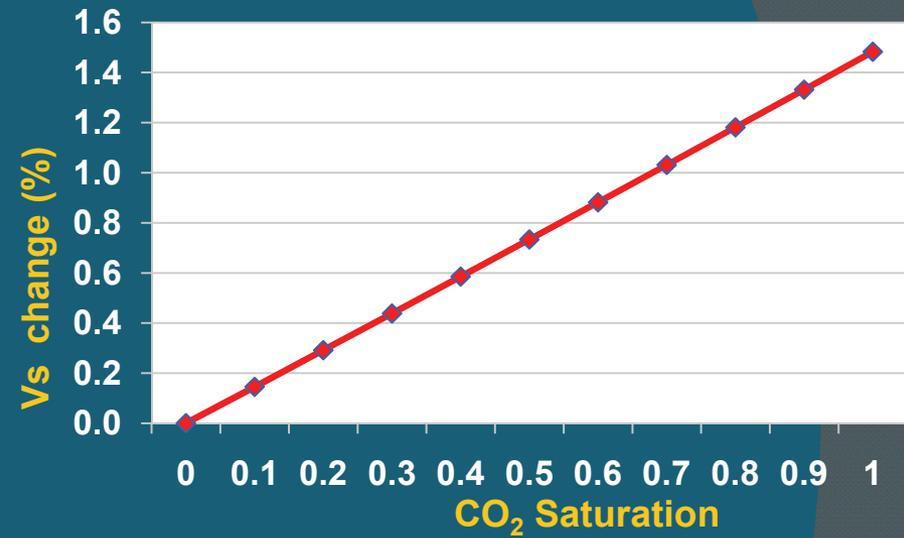
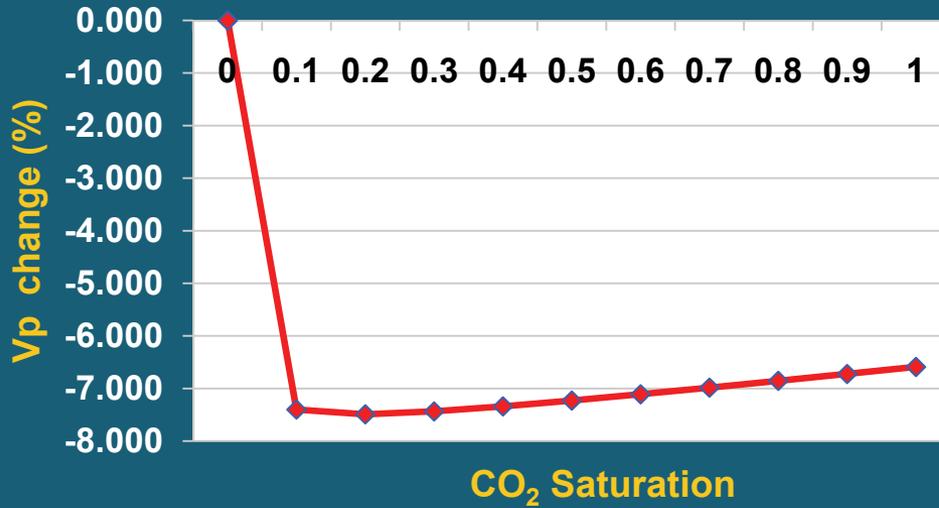
## Amplitude variations



## Velocity variations

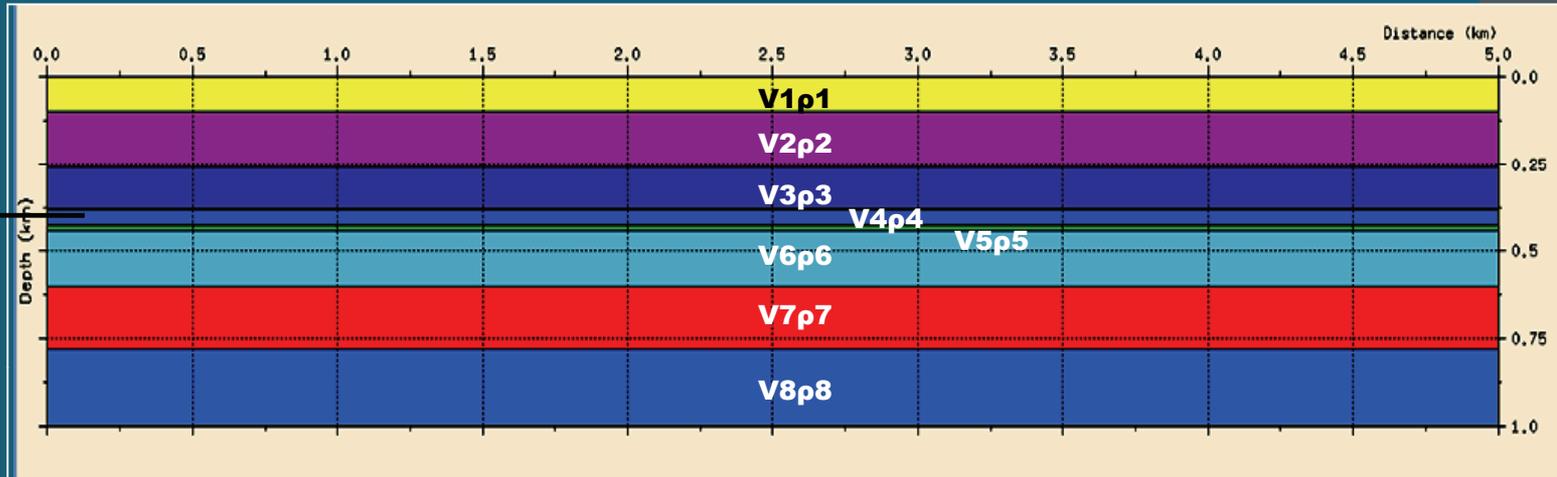


## $V_p$ and $V_s$ changes

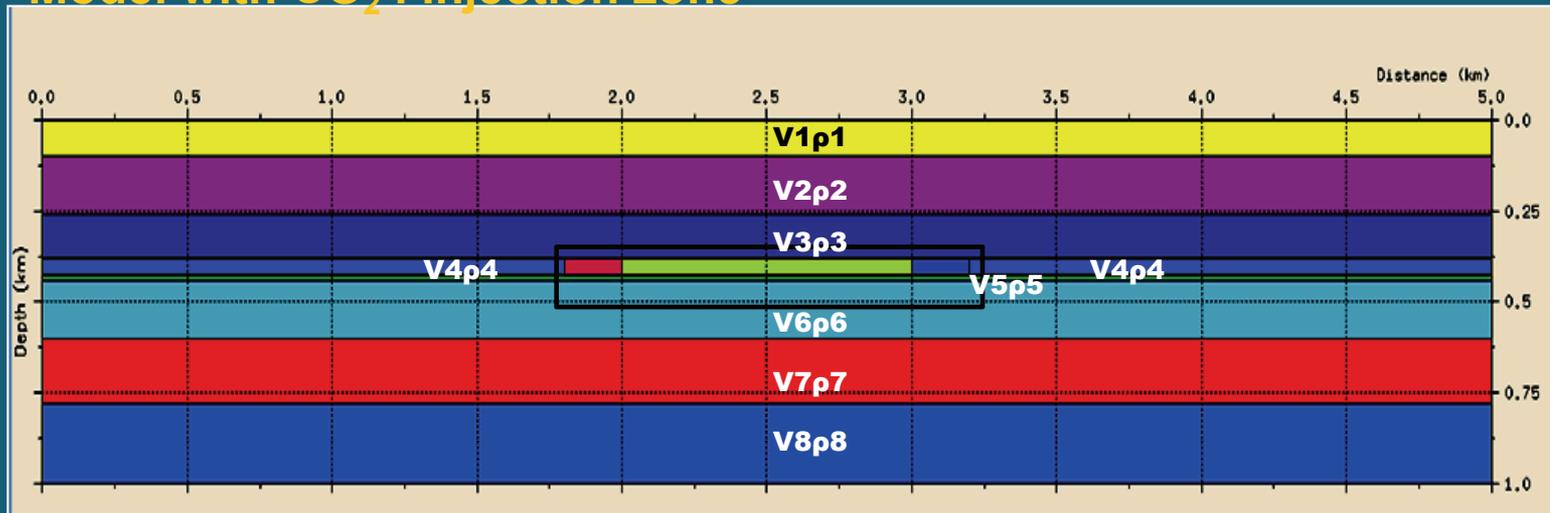


## Geological model

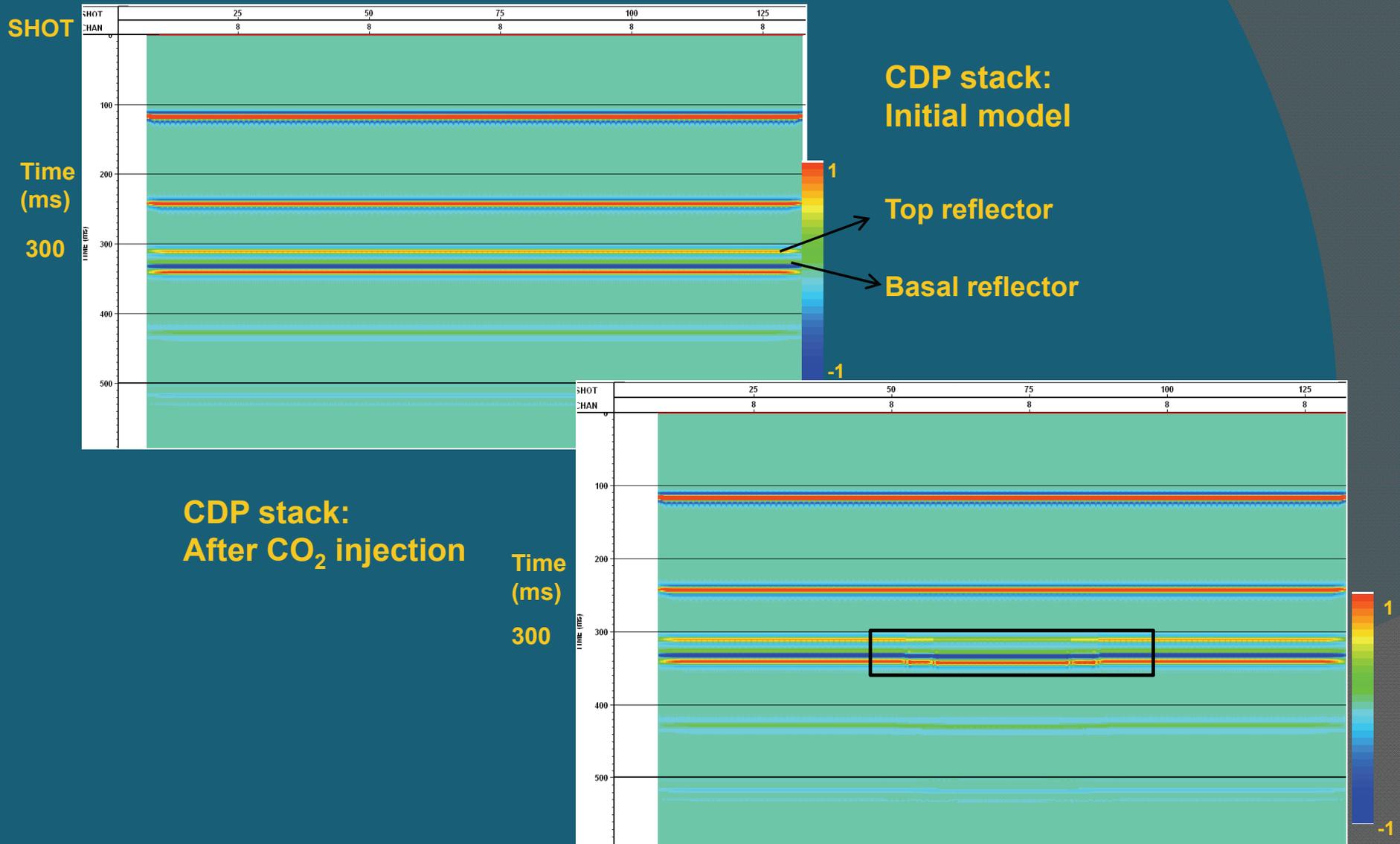
### Initial model



### Model with CO<sub>2</sub> . Injection zone

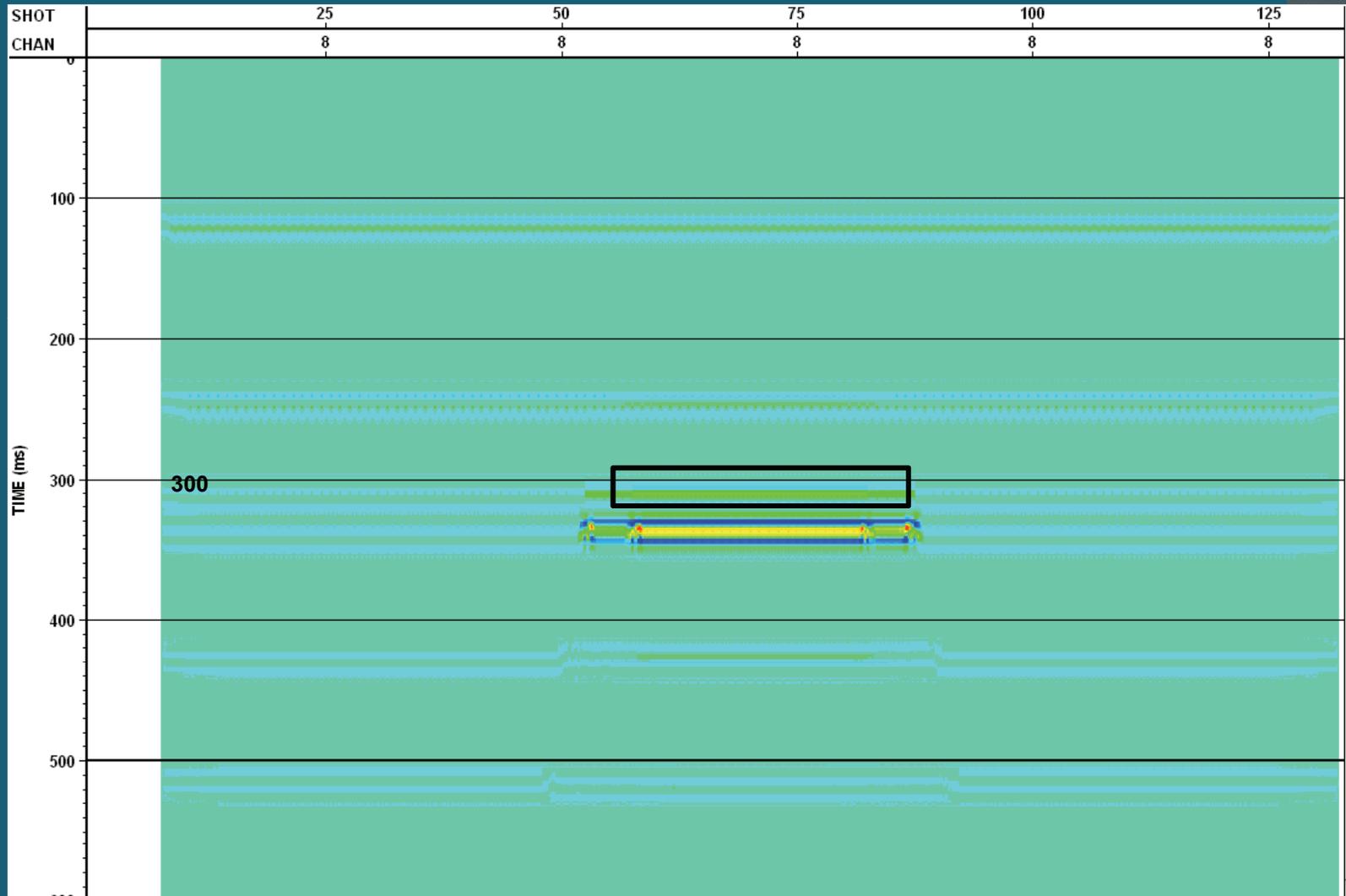


## Seismic modelling



## Difference between sections

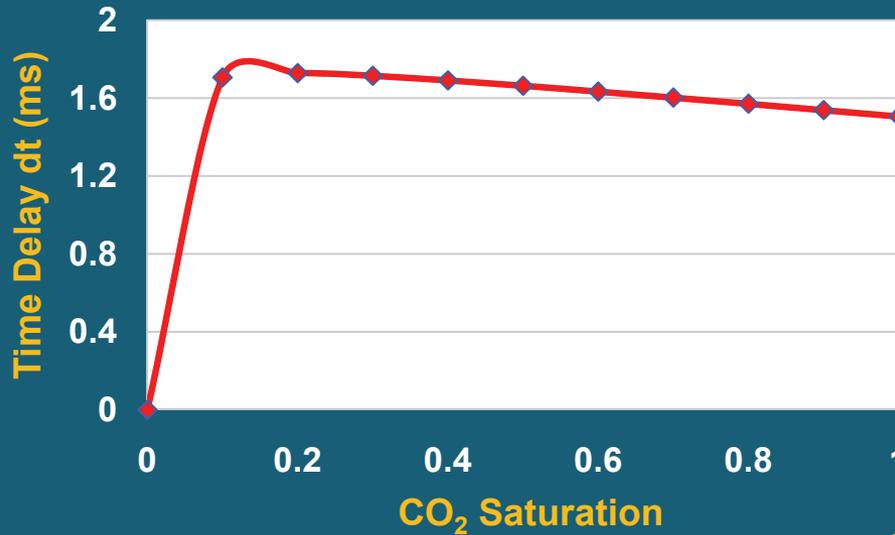
SHOT



Time (ms)

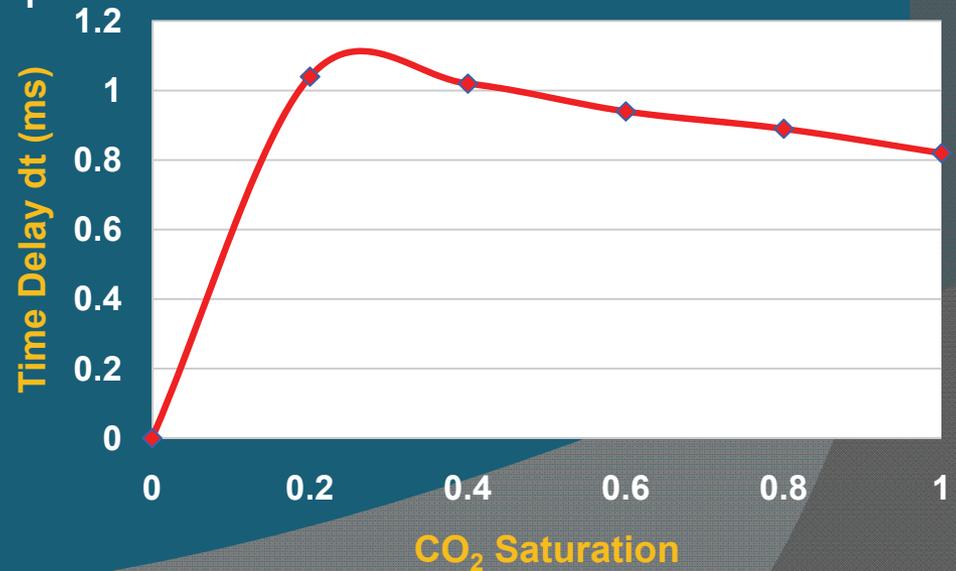
# Results

## Time delay (basal reflector) vs. CO<sub>2</sub> saturation



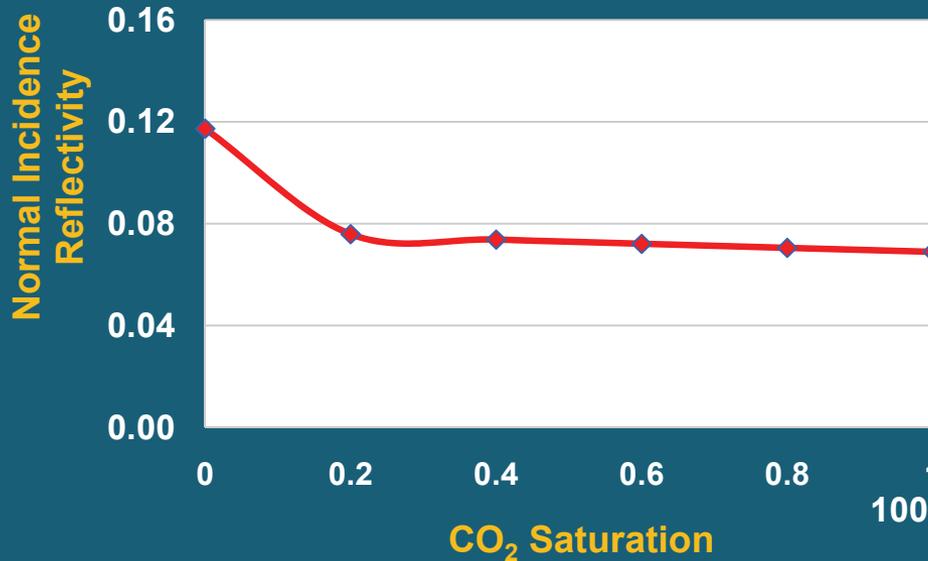
Measured from the section

Theoretical calculation:  
$$\Delta T = T_2 - T_1 = 2H (1/V_2 - 1/V_1)$$



# Results

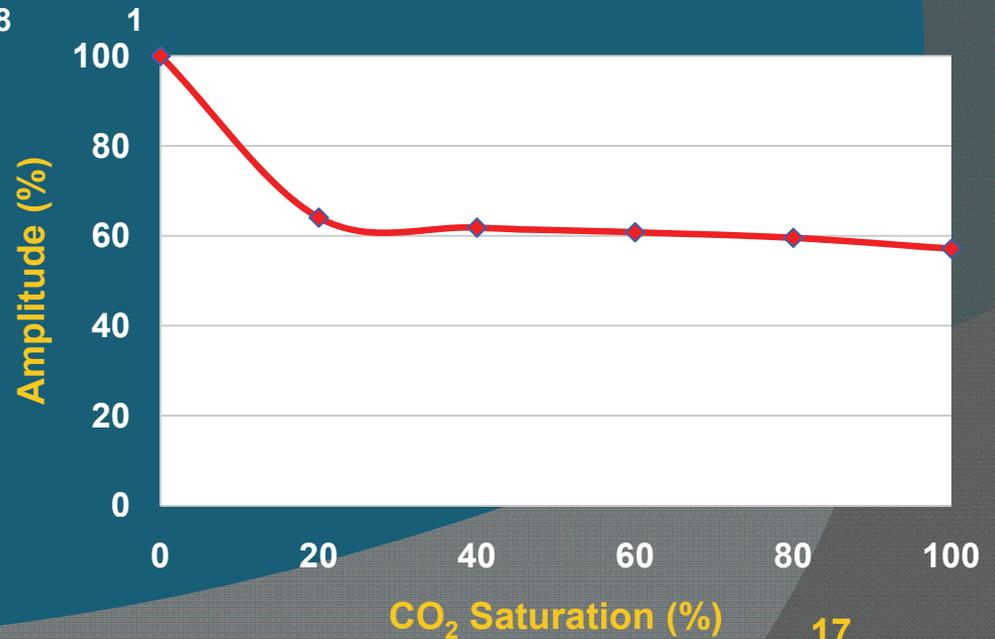
## Amplitude (top reflector) vs. CO<sub>2</sub> saturation



Measured from  
the section

Theoretical calculation:

$$R_{ij} = (\rho_j V_j - \rho_i V_i) / (\rho_j V_j + \rho_i V_i)$$



## AVO: Shuey's approximation

$$R_{pp} = A + B \sin^2 \theta_i + C(\tan^2 \theta_i - \sin^2 \theta_i)$$

$R_{pp}$  = P-wave Reflectivity

A = First Coefficient, normal Incidence Reflectivity

B = Second Coefficient, small angle (<30 degrees)

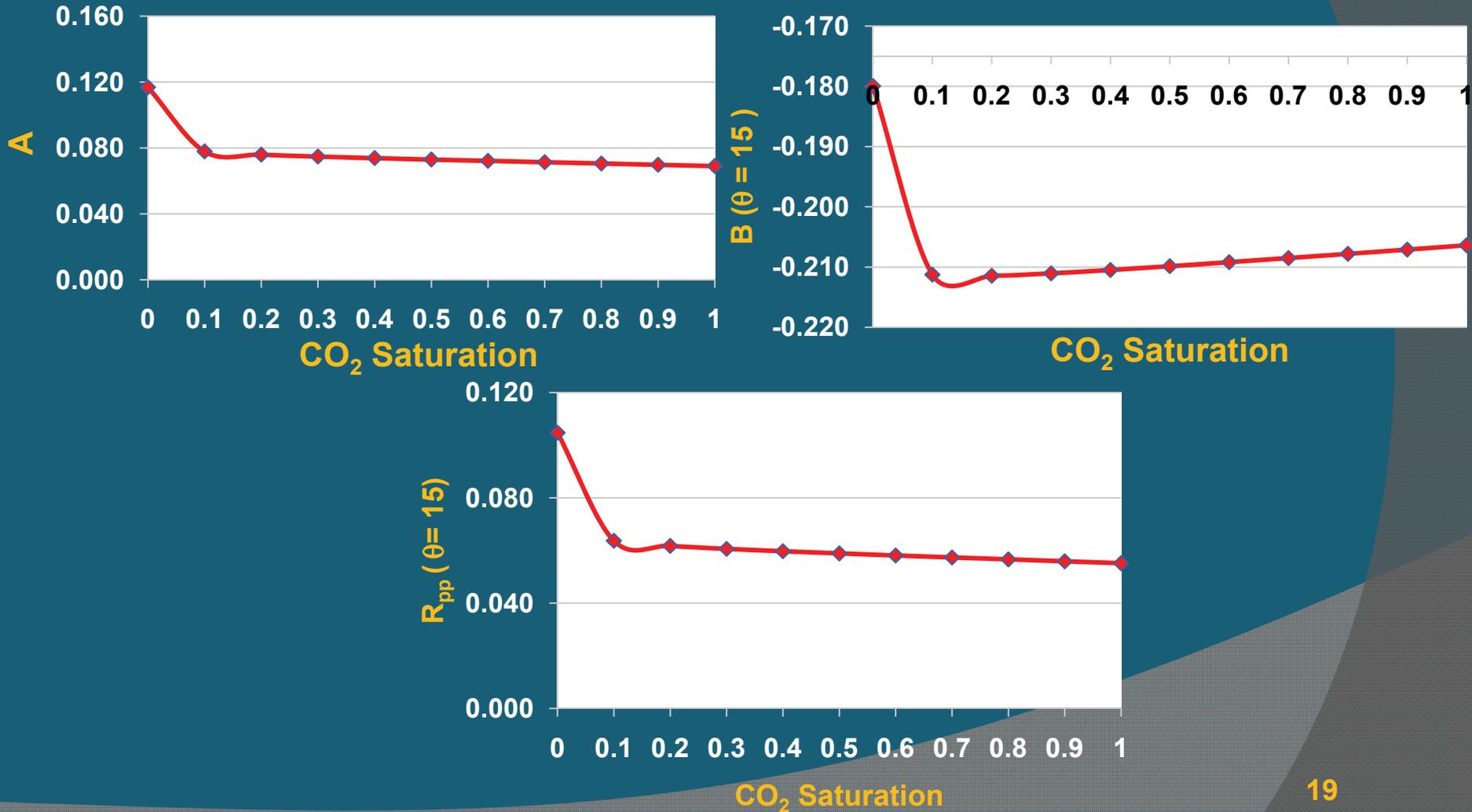
C = Third Coefficient, larger angle (>30 degrees)

$\theta_i$  = Angle of Incidence

$$R_{pp} = A + B \sin^2 \theta_i$$

## AVO: Shuey's approximation

Having an angle of Incidence,  $\theta = 15$  degrees

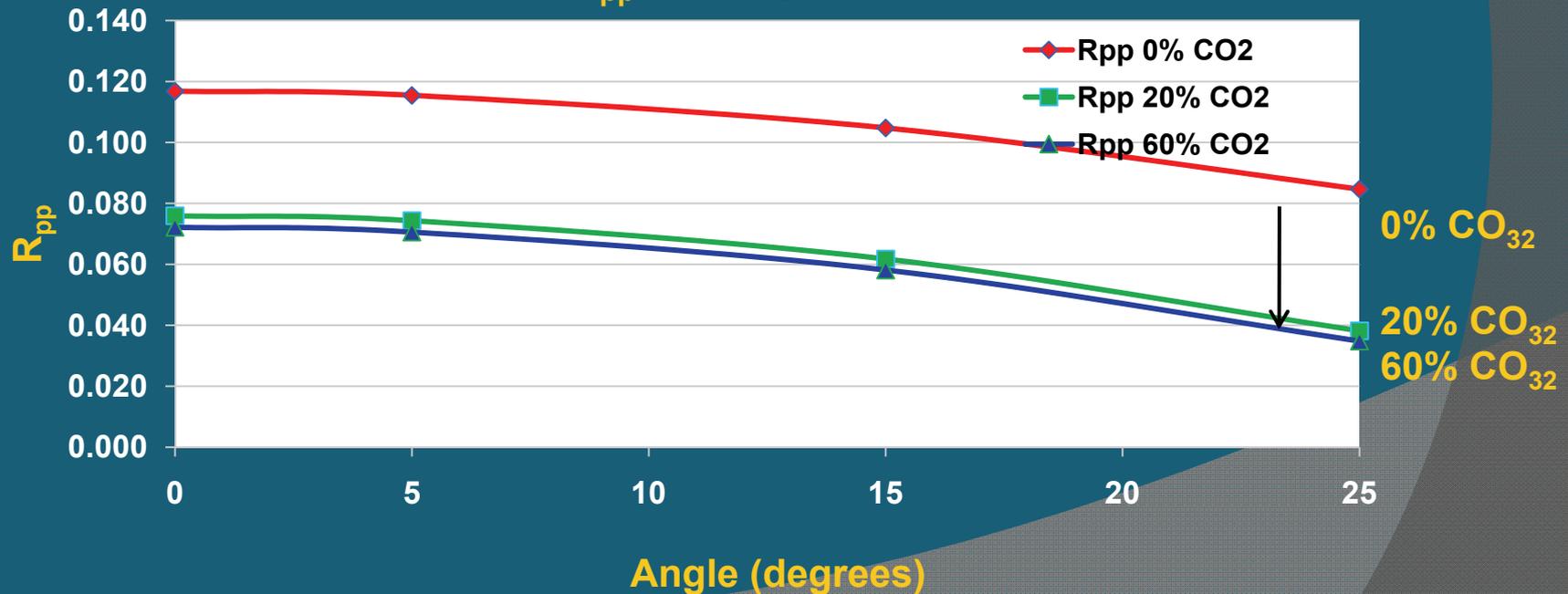


## AVO: Shuey's approximation

### AVA (Amplitude Vs. Angle)

Angle (Degrees)	Rpp(0% CO <sub>2</sub> )	Rpp(20% CO <sub>2</sub> )	Rpp(60% CO <sub>2</sub> )
0	0.117	0.0759	0.0721
5	0.115	0.0743	0.0706
15	0.105	0.0618	0.0581
25	0.085	0.0382	0.0348

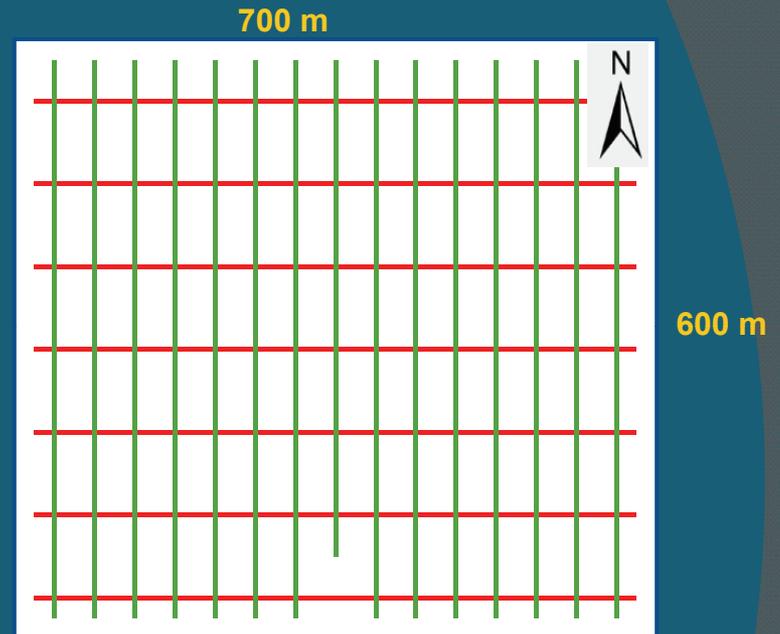
R<sub>pp</sub> Vs Angle



# Recent Progress

## Field acquisition

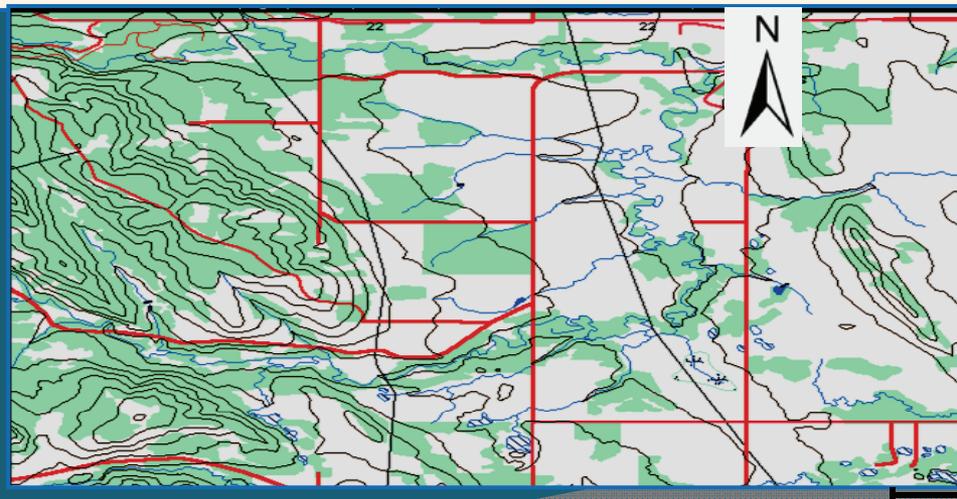
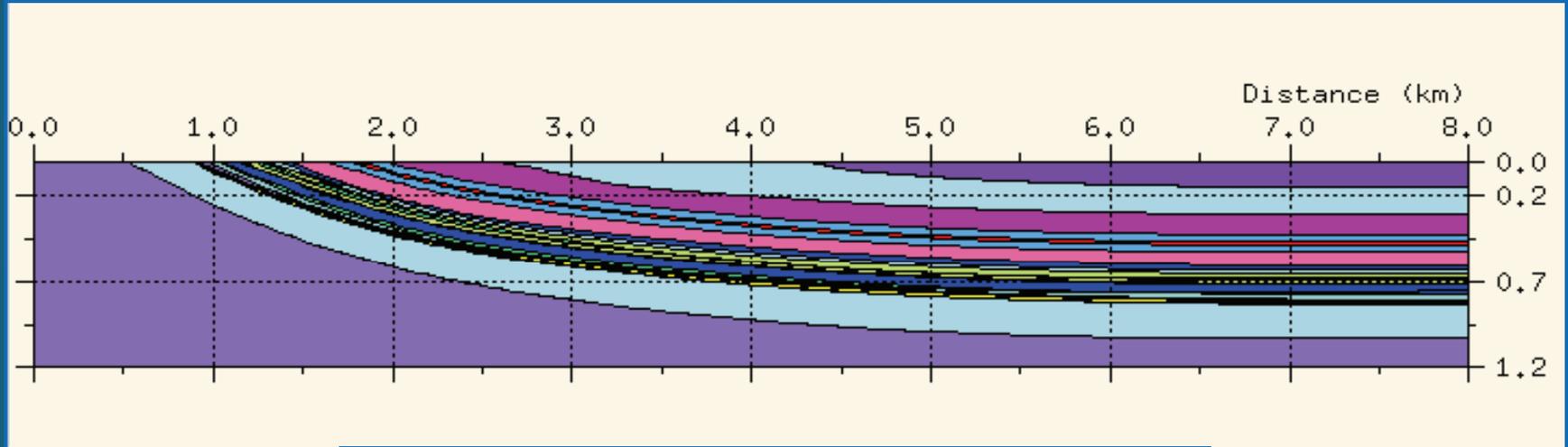
**May, 2010: 3D 3C Seismic Survey**  
(15 receivers lines and 7 source lines)



**August, 2010: 2D Seismic Survey**  
(One 2.5 km Line)

# Recent Progress

## Extended geological model



1 km

# Conclusions

- Paskapoo Formation has suitable qualities for a test CO<sub>2</sub> geological storage site.
- Gassmann theory is a practical and useful tool in fluid substitution models.
- There is a recognizable difference in rock properties and seismic response due to CO<sub>2</sub>:
  - 1) P-wave velocity drops ~7% from 0% to 20% CO<sub>2</sub> saturation

# Conclusions

- 2) S-wave velocity increase is directly proportional to CO<sub>2</sub> saturation (average  $V_s$  increase 0.8 %).
- 3)  $V_p / V_s$  shows a decrease of 8%
- 4) Time delay of the bottom reflector is ~ 1.6 ms
- 5) AVO gradient will decrease with CO<sub>2</sub> saturation, particularly between 0% and 20 % CO<sub>2</sub> saturation.

# Acknowledgements

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# Thanks!

