

Why seismic-to-well ties are difficult

Gary F. Margrave









It's Q's Fault

Gary F. Margrave







Outline

- Well tying: standard techniques
- The easy case that is mostly irrelevant.
- The nonstationary trace model
- Physical effects of attenuation (Q)
- Nonstationary seismogram from well logs
- Failure of stationary deconvolution
- Nonstationary deconvolution: Inverse Q filters and Gabor deconvolution
- Inclusion of noise
- Conclusions

Well Tying













We seek a minimal sequence of adjustments that establish an "acceptable" match.

- depth (m) Wiener match filters are "too much"
 - Adjustments should have a physical motivation
 - Common steps are: timing adjustments, amplitude spectral shaping, phase rotations



Wavelet Estimation: Amplitude spectrum Estimate amplitude spectrum of wavelet by smoothing the spectrum of the seismic data.



Wavelet Estimation: phase spectrum

Assume a simple phase: constant phase, minimum phase, or time-variant constant phase.

Example: Constant phase scan: $err(\theta) = \sum (s_0 - phsrot(s_1, \theta))^2$



The Easy, Nonphysical Case

The easy but nonphysical case

Noise-free convolutional seismogram with a minimum phase wavelet and white reflectivity

$s(t) = w(t) \Box r(t)$



The deconvolved seismogram will tie the reflectivity almost exactly.

The easy but nonphysical case



What's wrong with this? $s(t) = w(t) \Box r(t)$ *A: Almost everything!*

- Real reflectivity is not white
- The wavelet, while possibly minimum phase, evolves continuously
- Convolution cannot be stationary
- Noise is fundamentally important

Nonstationary trace model (CREWES q_tools)

Nonstationary Convolution Model Better trace model

 $s = W_0 W_Q r$



Nonstationary Convolution Model Better trace model



Nonstationary Convolution Model Better trace model



Factoring W₀W_Q



Nonstationary Q matrix

 W_Q

Note: $W_0 W_Q - W_Q W_0 \neq 0$ but is usually very small

Physical Effects of Anelastic Attenuation (CREWES q_tools)

Wavelet evolution

1D attenuating earth



Impulse in

Blob out



Wavelet evolution and drift Frequency dependence of phase velocity leads to "drift"



Wavelet evolution and drift Frequency dependence of phase velocity leads to "drift"



Nonstationary seismogram from well logs

Hussar 12-27 Logs



Synthetic seismograms from Hussar 12-27 Logs, logging depths doubled



Failure of Stationary Deconvolution

Stationary decon of stationary seismogram the unphysical case



Stationary decon of nonstationary seismogram



Nonstationary catastrophe



Understanding the Nonstationary catastrophe



Nonstationary catastrophe Can spectral shaping help?

Impose the amplitude spectrum from the ideal case



Nonstationary catastrophe Can phase rotations help? Find best constant phase rotation in local Gabor windows (drift correction should be done first)



Nonstationary catastrophe Can AGC help?

Global: max corr=0.22737, lag=-0.9 samples Local: max corr=0.29345, lag=-4 samples Early: max corr=0.55126, lag=-0.8 samples Late: max corr=0.17597, lag=-11.2 samples



Stationary deconvolution Nonstationary band-aid



Phase error analysis time-variant constant phase estimates



Inverse Q filter and Gabor deconvolution

Inverse Q filter renders the nonstationary trace stationary



Inverse Q filter then Wiener decon avoiding the nonstationary catastrophe



Gabor decon avoiding the nonstationary catastrophe



Phase rotation analysis Inverse Q->deconw compared to Gabor decon



Inclusion of noise

Noisy seismograms S2N=2 in design window of nonstationary trace Identical noise added to both seismograms



Inverse Q filter on noisy data



Inverse Q filter then Wiener decon on noisy data



Gabor deconvolution on noisy data



Phase rotation analysis Inverse Q->deconw compared to Gabor decon on noisy data



Conclusions

- Anelastic attenuation, which is always present, ensures that the convolutional model is approximate at best.
- Real seismic data does not have a single "wavelet" but rather an evolving wavelet determine by the Q structure.
- The nonstationary convolutional model captures the firstorder effects of the evolving wavelet.
- Applying stationary deconvolution to a nonstationary seismogram is reasonable in the design widow but produces severe distortions elsewhere.
- Standard well-tying procedures of wavelet shaping and phase rotation have limited success outside the design window.
- Nonstationary deconvolution processes are required for better results but have their own problems.
- Noise affects the nonstationary trace severely and ensures that the signal band is time-variant.



Acknowledgements

My thanks to all CREWES sponsors for supporting this research. Thanks also the NSERC, CMC, and the UofC for additional support.







Well versus synthetic reflectivities



Well versus synthetic reflectivities



Wavelet Estimation: Constant phase rotations



 $W_0 W_Q \sim W_Q W_0$ Almost commutative



Stationary decon of stationary seismogram for reference



And the state of the

Stationary decon of stationary seismogram for reference



Stationary decon of stationary seismogram for reference

