

Time-lapse attenuation variations during CO₂ injection using DAS VSP data from the CaMI.FRS

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- DAS VSP survey and data processing
- Time-lapse attenuation measurements
- Accuracy and resolution of estimated attenuation
- Discussion
- Conclusions





Time shifts along the first reflection horizon below CO₂ injection zone

- Injection zone at the FRS is thin (~7 m)
- Injected CO₂ may cause a decrease in P-wave velocity of only a few percent
- Such a decrease is near the seismic noise level for detection
- S/N of VSP data obtained through DAS fiber cable is generally lower than that from geophone
- Compared with velocity or density variations, attenuation is more sensitive to changes in saturation, permeability, and pore fluid



Location map of CaMI FRS in southern Alberta, Canada



- Monitor survey acquired in March 2021
- ~34 tonnes of CO₂ injected prior to this date





- Continuous fiber loop (5km) and raw DAS shot gather
- Tag numbers indicate the corresponding data from different portions of fiber loop
- DAS data from the straight fiber in OBS2
- Identical data processing for different vintages of DAS data



- Construct attenuation-attribute images from reflection signal for monitoring injected CO₂
- Separate downgoing and upgoing wavefields through a median filter followed by a f-k filter
- Upgoing P-wave reflections are clearly isolated and evident over the entire depth coverage

1) Non-stationary convolutional model for reflection signal:

A(t,f) = S(f)R(t,f)F(t,f)G(t)

- *A*: time-frequency variant amplitude of reflection signal
- *S*: spectrum of seismic source
- R: time-frequency response of reflectivity
- F: seismic attenuation $F(t, f) = \exp\left[-\pi f \int_0^t d\tau / Q(\tau)\right]$ (anelastic)
- G: focusing or defocusing $G(t) = \exp\left[-\int_0^t \gamma(\tau) d\tau\right]$ (elastic)

2) Take logarithm and time derivative:

$$\gamma + \pi f Q^{-1} = \frac{\partial}{\partial t} \Big[\ln R(t, f) - \ln A(t, f) \Big]$$

- γ is a "geometric" counterpart for Q^{-1}
- 3) Under the assumption R(t, f) = const:

$$\gamma_A + \pi f Q_A^{-1} = -\frac{\partial \ln A(t, f)}{\partial t}$$

- γ_A and Q_A are the "apparent" attenuation
- Require smooth estimates of *A*(*t*, *f*)



$$s(t_i, \tau) = c_0 \sum_{j \in \text{segment } i} h\left(\frac{t_j - t_i}{T_s}\right) w_j(\tau)$$

- c_0 scales peak amplitude equal to $s(t_i, \tau) = 1$
- Hann window h of the same length T_s with segment i
 - Apply weights to $w_j(\tau)$ located within different parts centered on t_j of each peak
 - Located at the middle of segment has the maximum weight and at the edge has a zero weight
- Temporal resolution of the resulting spectra can be estimated as the half segment length $T_s/2$

Attenuation measurements from monitor data



- Large γ_A and Q_A^{-1} around OBS2 at the BBRS injection zone
- Increased CO₂ saturation and/or pore pressure
- Consistent γ_A and Q_A^{-1} around OBS2 from two different lines

Time-lapse attenuation variations



- Measurements from baseline data show there is no anomaly prior to CO₂ injection
- Time-lapse responses show a clear increase to the SW of OBS2
- Attenuation increase not around the injection well indicates the migration of injected CO₂
- From liquid phase around the injection well to gas phase at the leading edge of the plume



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$$a_k = \gamma + \pi \left(f_k - \overline{f} \right) Q^{-1}$$

Empirical error

Errors for regression coefficients (Q^{-1} and γ) are

$$s_{\hat{\gamma}} = \frac{\sigma_{\hat{a}}}{\sqrt{n}}$$
 (intercept

$$s_{\hat{\mathcal{Q}}^{-1}} = \frac{\sigma_{\hat{a}}}{\pi \sqrt{\sum \left(f_k - \overline{f}\right)^2}} = \frac{s_{\hat{\gamma}}}{\pi \left\langle f_k - \overline{f} \right\rangle_{RMS}}$$
(slope)

• $\langle ... \rangle_{\rm \tiny RMS}$: RMS average

• $\sigma_{\hat{a}}^2$: variance of the estimated spectral quantity

Theoretical error

$$s_{\hat{\gamma}} = \frac{1}{\sqrt{2\Delta t}} \frac{1}{\sqrt{BT}}$$
$$s_{\hat{Q}^{-1}} = \frac{s_{\hat{\gamma}}}{\pi cB} = \frac{1}{\pi\sqrt{2}c} \frac{1}{\sqrt{BT}} \frac{1}{B\Delta t}$$

- *B*: signal frequency bandwidth
- *T*: waveform length in spectral measurement
- Δt : temporal resolution
- Independent of γ or Q^{-1} value
- Key parameter is the desired temporal resolution Δt



- Resolution of $\Delta t = 200 \text{ ms}$
- *Q*⁻¹ is below theoretical error (green line) for most depths above and below the BBRS
- Limited accuracy of *Q*⁻¹ for areas outside of injection zone
- Around the BBRS that is with high attenuation, Q^{-1} is much larger than error
- Accuracy of γ is higher with the theoretical error smaller than γ for all the depths

Relative errors



- Black dashed line: resolution of $\Delta t = 200 \text{ ms}$
- Black solid line: a) estimated Q⁻¹ of 0.08; b) estimated γ of 10 s⁻¹
- Significant trade-off between accuracy of estimated Q^{-1} or γ and temporal resolution
- $Q^{-1} > 0.09$ and $|\gamma| > 6.5 \text{s}^{-1}$ for achieving an error lower than 30%
- Around the BBRS injection zone, relative errors are modest $e_{\hat{Q}^{-1}} \approx 34\%$ and $e_{\hat{\gamma}} \approx 20\%$



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Comparison with time shifts





Time shifts along the first reflection horizon below injection zone

Comparison with impedance variations





- Both Q⁻¹ and its geometric counterpart γ are measured from the DAS VSP data for monitoring the CO₂ injection.
- The variations of Q^{-1} or γ are correlated with injection zone and interpretated as being caused by an increase in CO₂ saturation.
- Evaluations of measurement accuracy show the measured Q^{-1} or γ of large values around the injection zone are reliable.
- Attenuation is a sensitive property that can complement velocity change for monitoring injected CO₂ during geologic CO₂ storage.

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