

Direct measurement of frequency-dependent phase velocities from Snowflake data

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CREWES sponsor meeting



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Department of Earth, Energy, and Environment



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- Motivation
- Background
- Method
- Discussion
 - Semi-automatic picking procedure
 - Automatic picking procedure
- Summary



- In a dispersive medium, seismic waves of varying frequencies travel with different velocities; leading to frequency-dependent attenuation.
- Understanding the dispersive properties in the near-surface by estimating the frequency-dependent phase velocity helps in determining the quality factor(Q).



CMC Newell County Facility

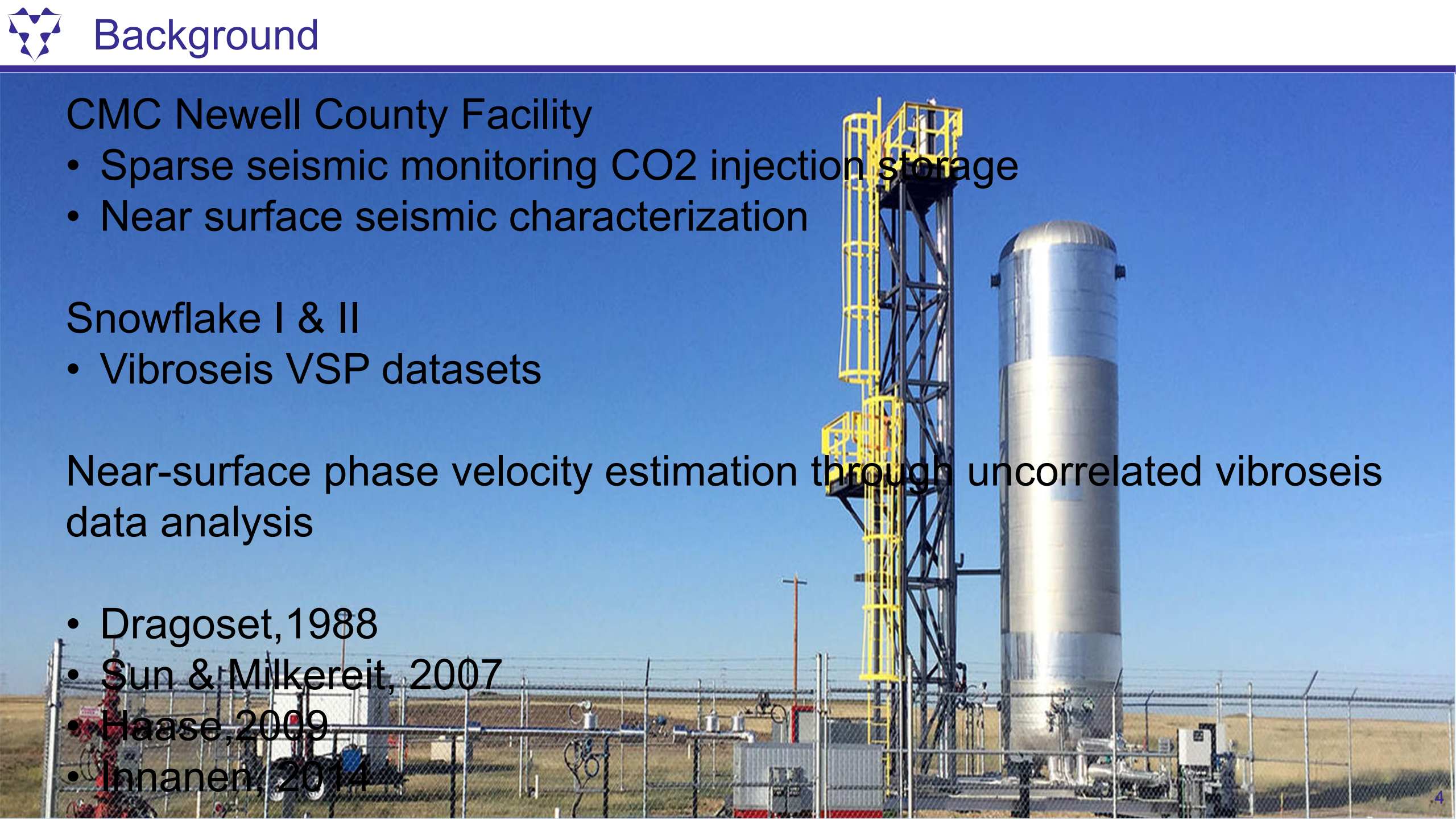
- Sparse seismic monitoring CO₂ injection storage
- Near surface seismic characterization

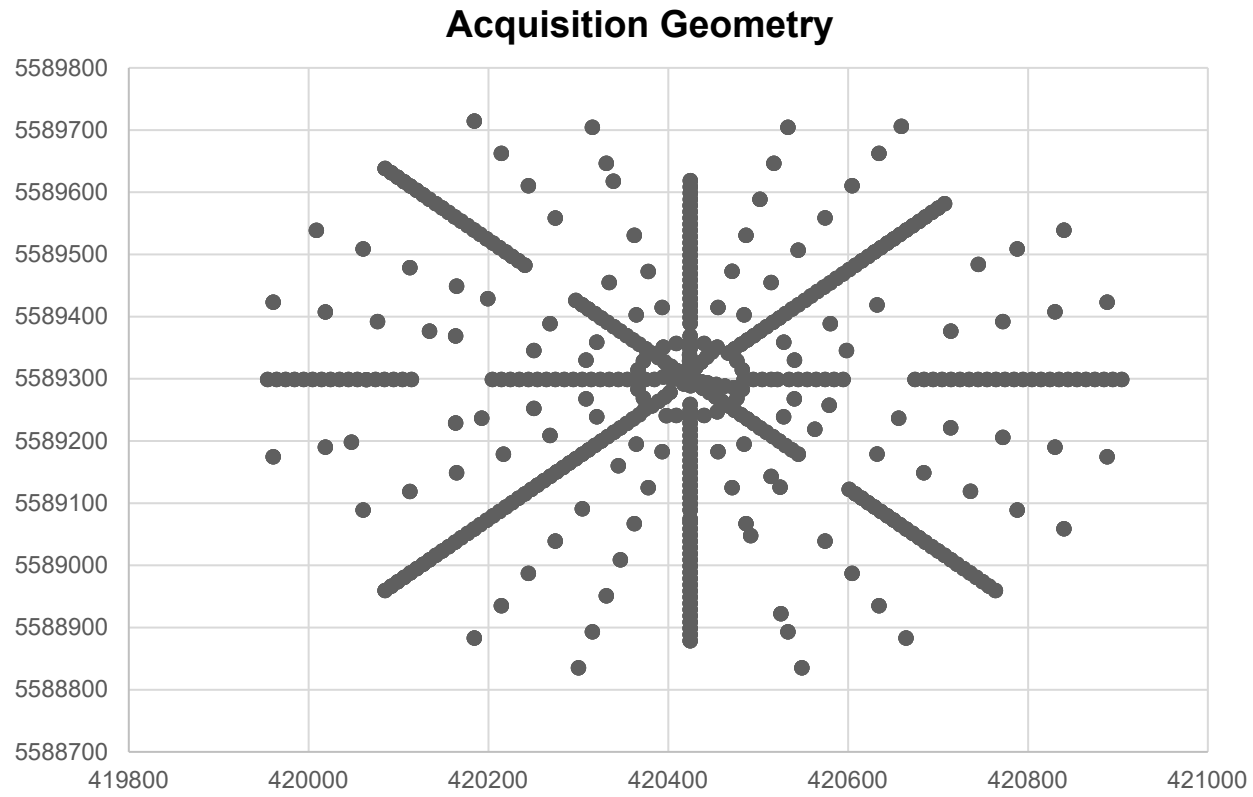
Snowflake I & II

- Vibroseis VSP datasets

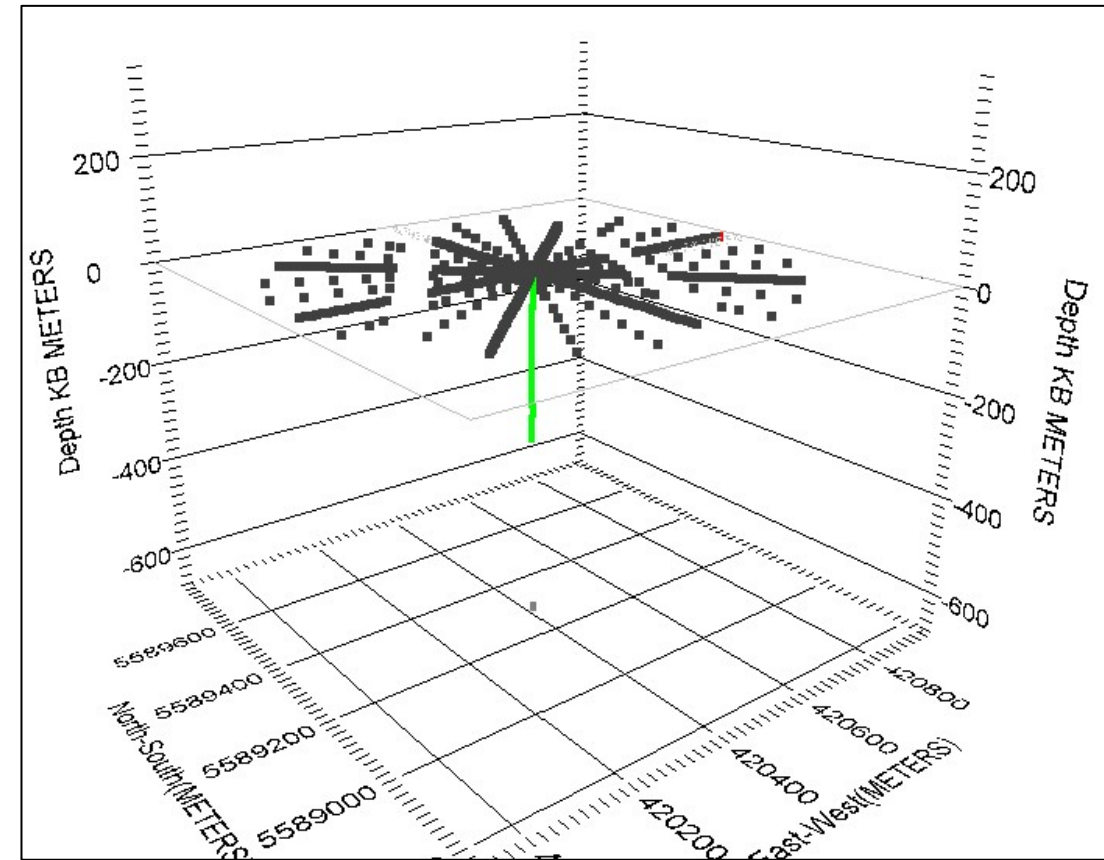
Near-surface phase velocity estimation through uncorrelated vibroseis data analysis

- Dragoset, 1988
- Sun & Milkereit, 2007
- Haase, 2009
- Innanen, 2014





Planar View



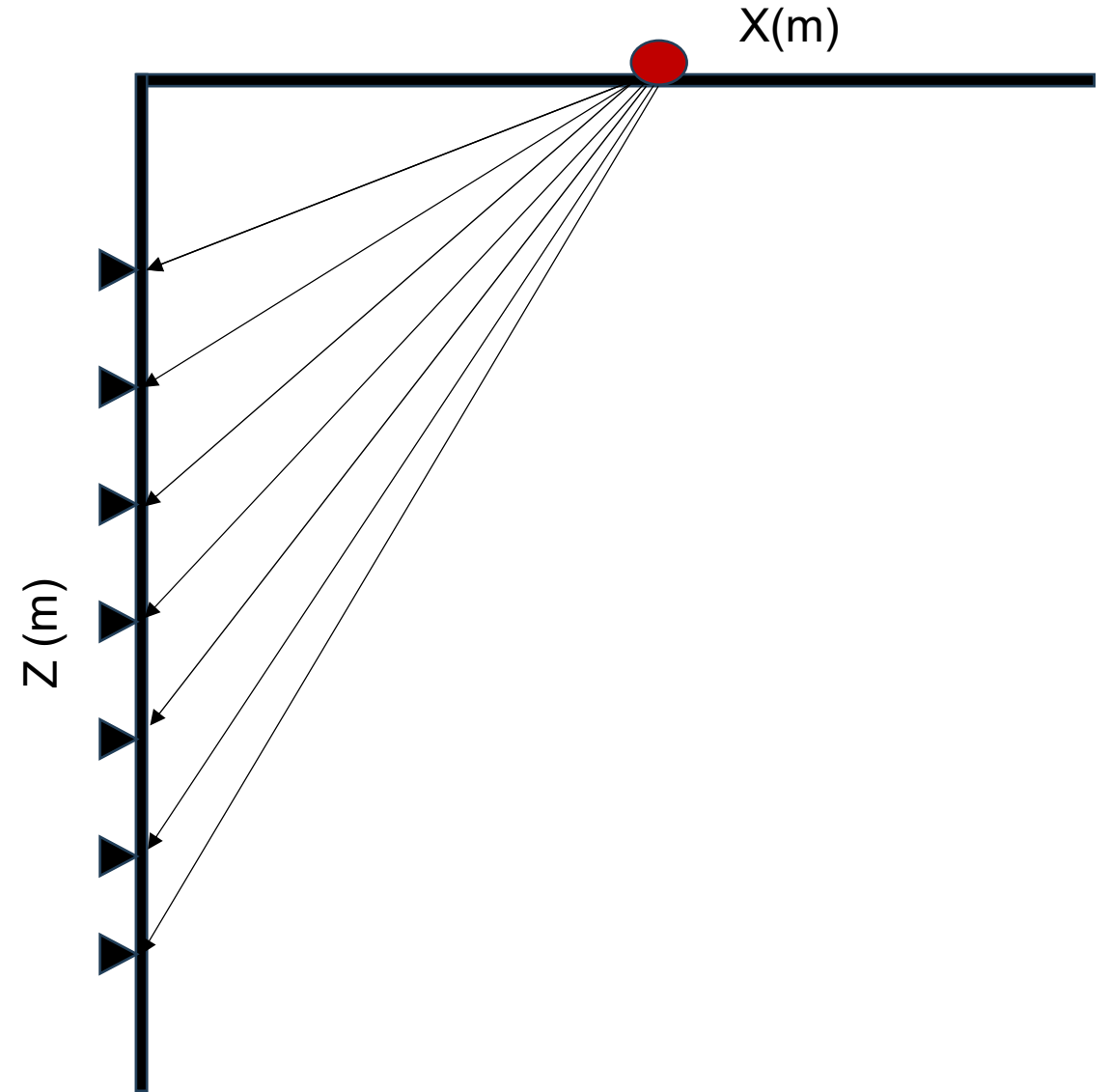
Side view



$$V_p(f) = \frac{L}{\Delta T(f)}$$

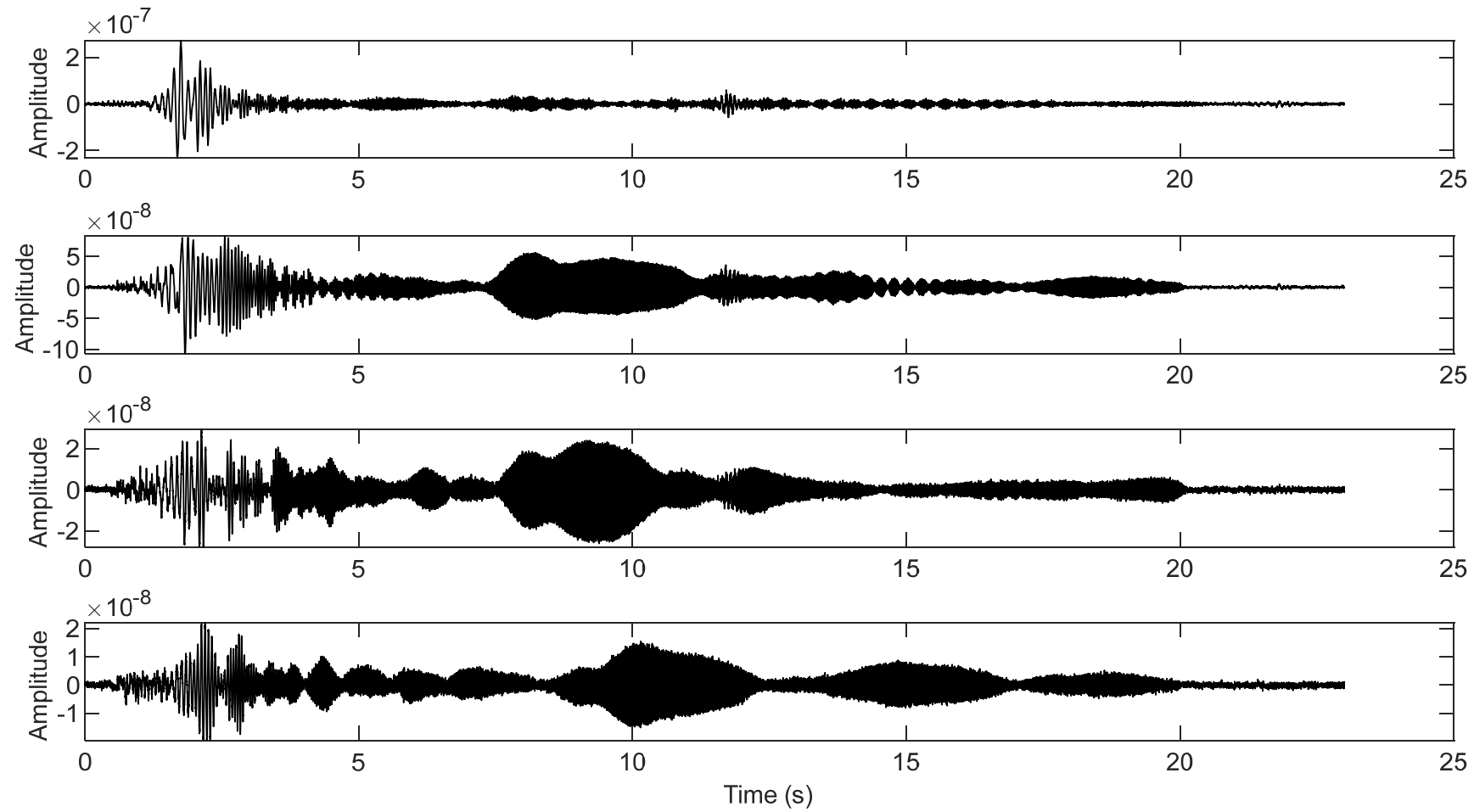
$$L = \sqrt{x^2 + z^2}$$

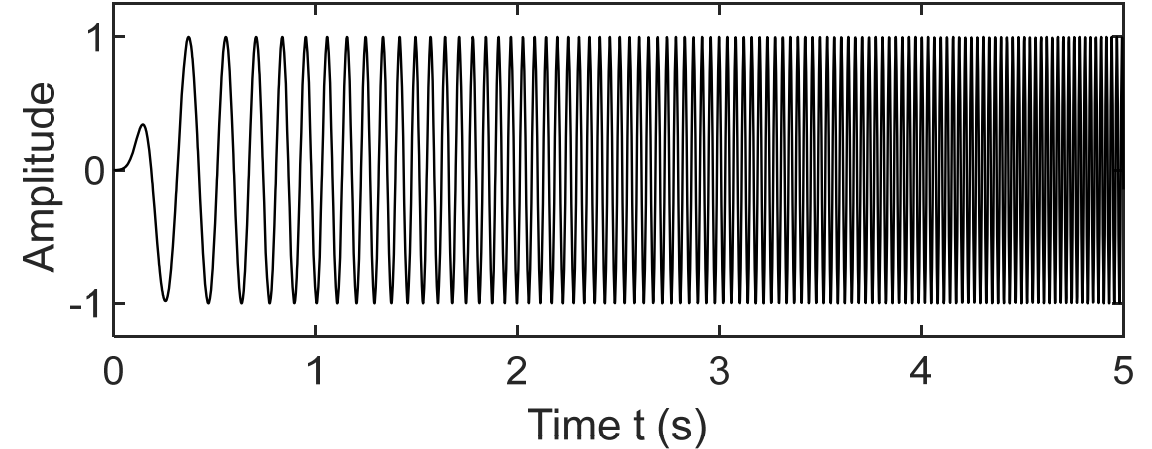
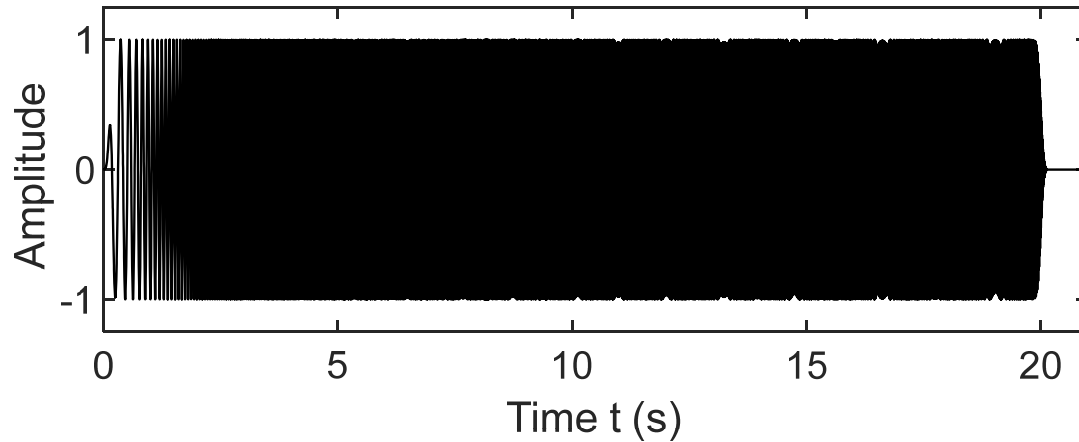
$$\Delta T(f) = T_M(f) - T_s(f)$$





Extracting uncorrelated seismic traces from a shot record



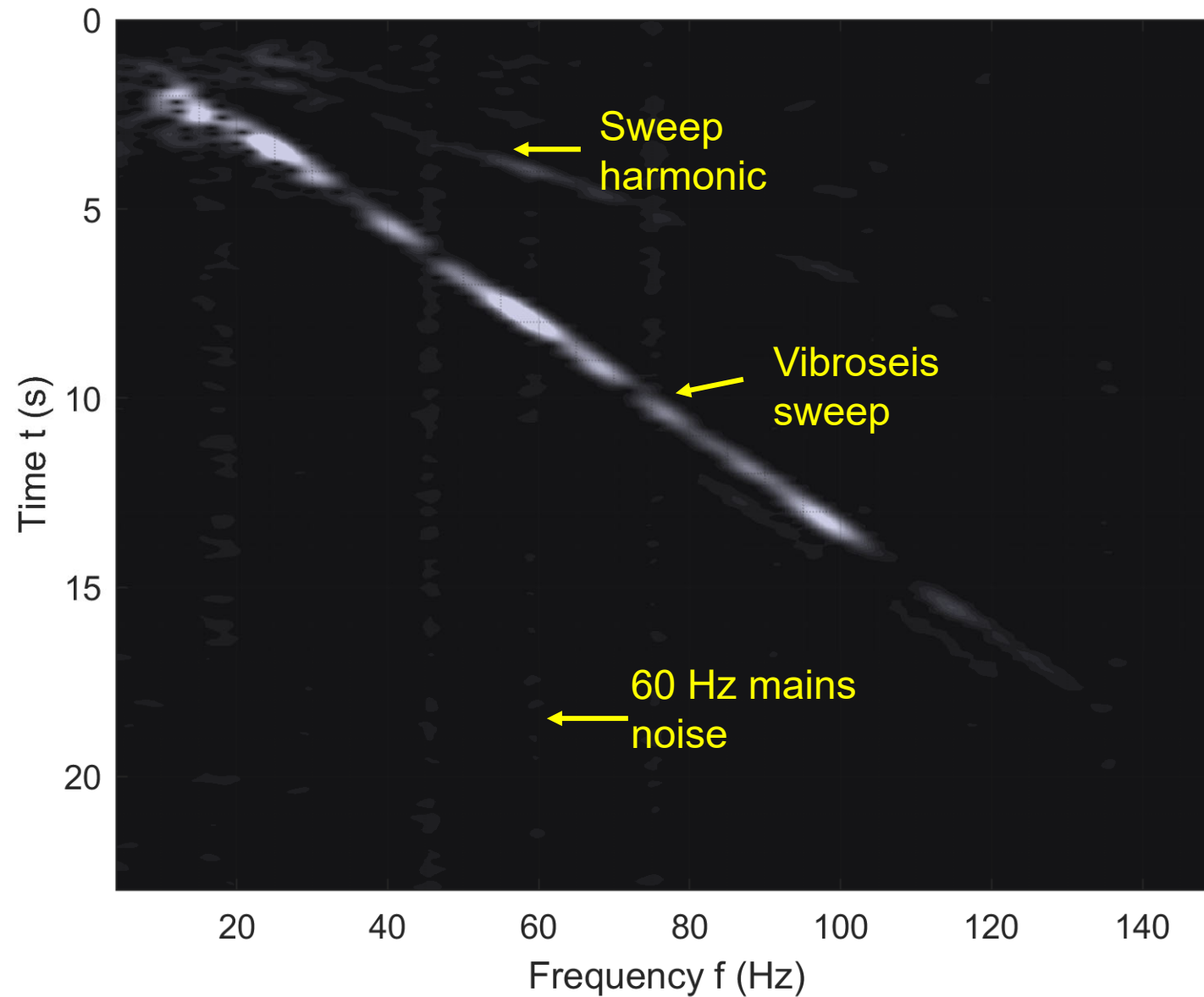


$$f_{min} = 2\text{Hz}$$
$$f_{max} = 150\text{Hz}$$
$$T = 20\text{s}$$

$$S(t) = \text{Im}[a(t)e^{i\phi(t)}] \quad ; \quad \phi(t) = 2\pi f(t)t \quad ; \quad f(t) = f_{min} + \frac{(f_{max} - f_{min})t}{T}$$

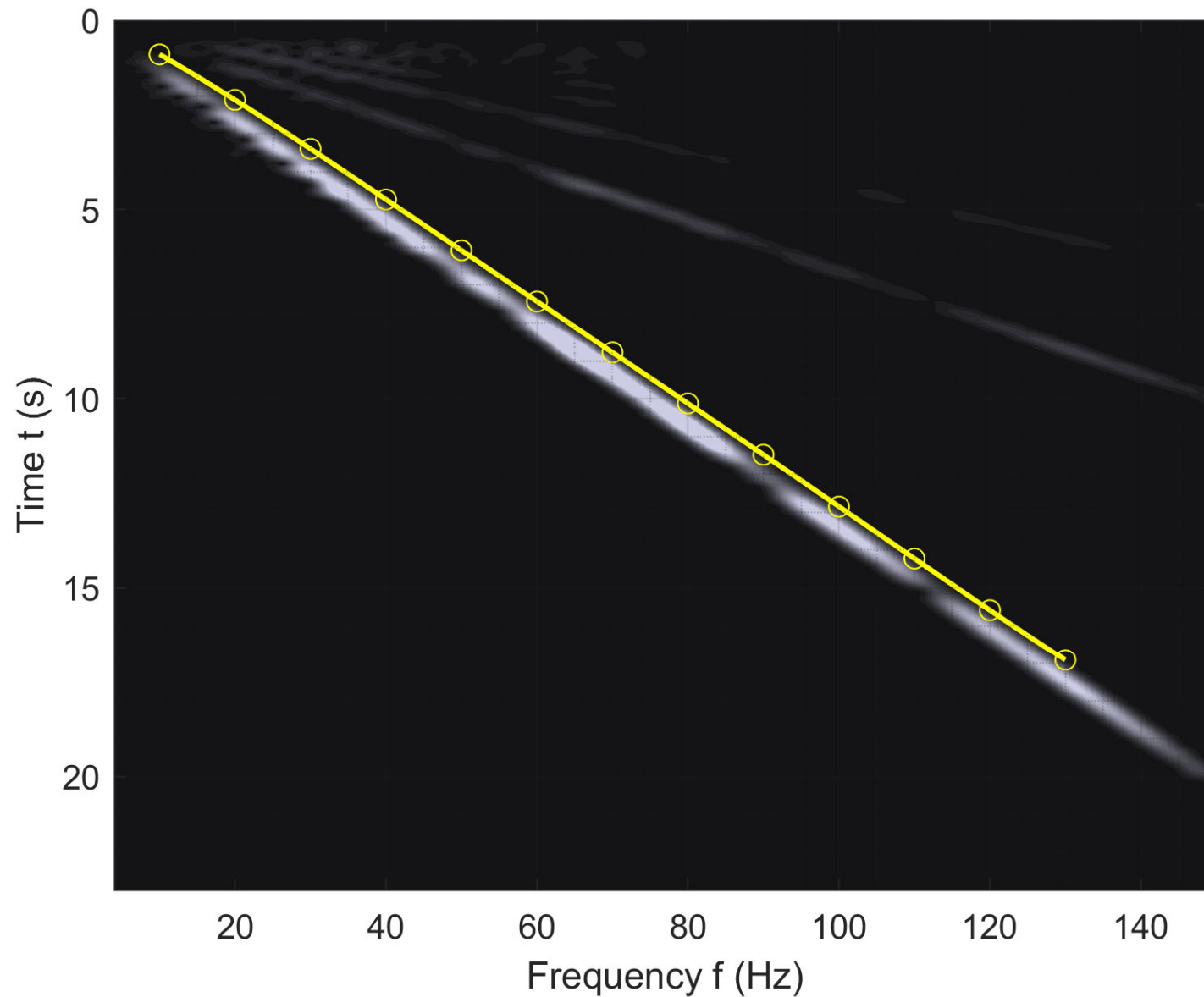


Gabor spectrum of a seismic trace





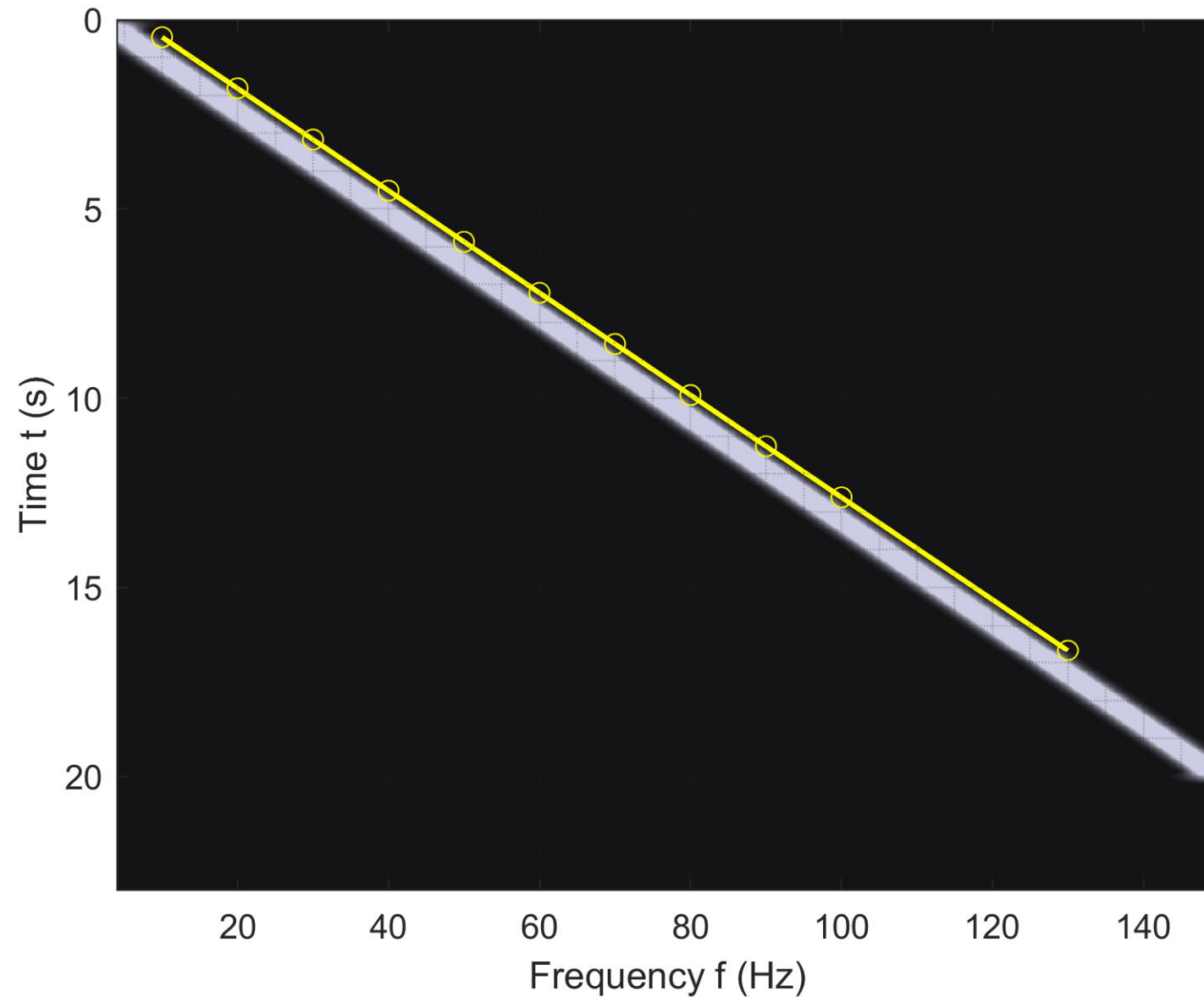
Arrival time picks on the seismic trace Gabor spectrum



Depth = 140m

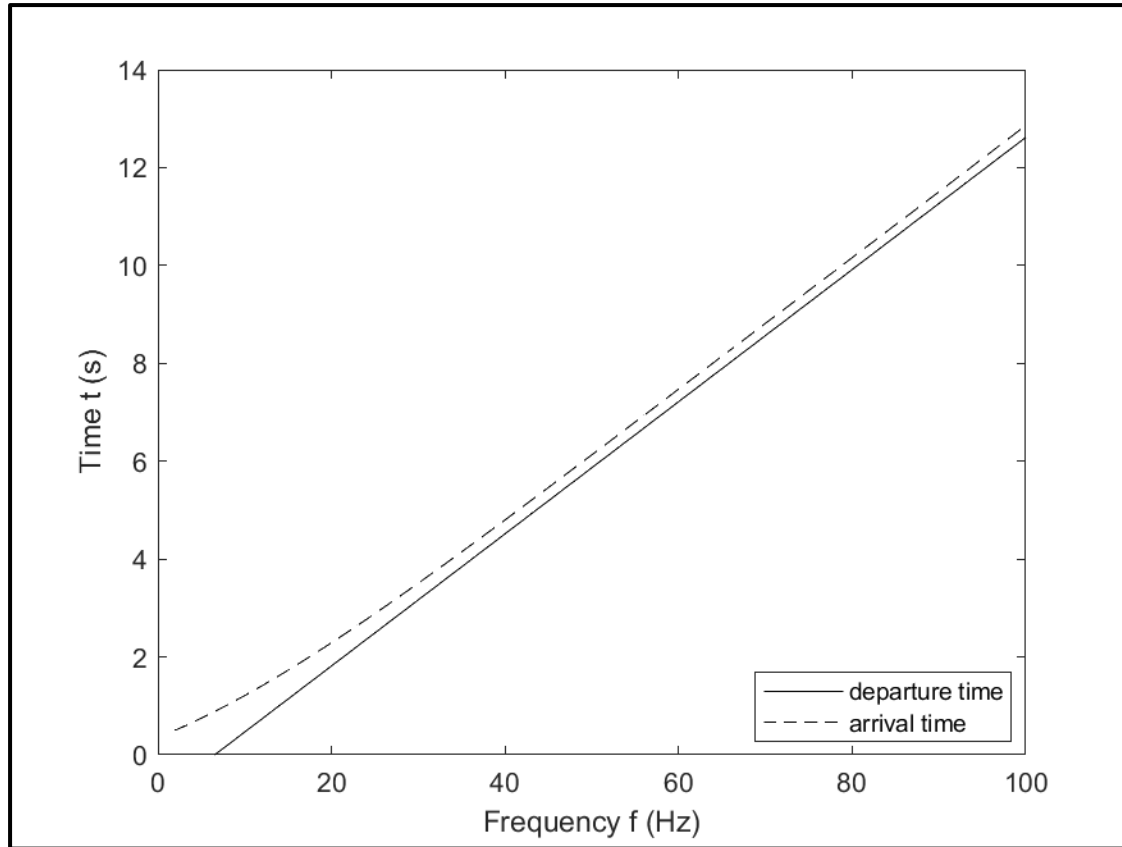


Departure time picks on the linear sweep Gabor spectrum

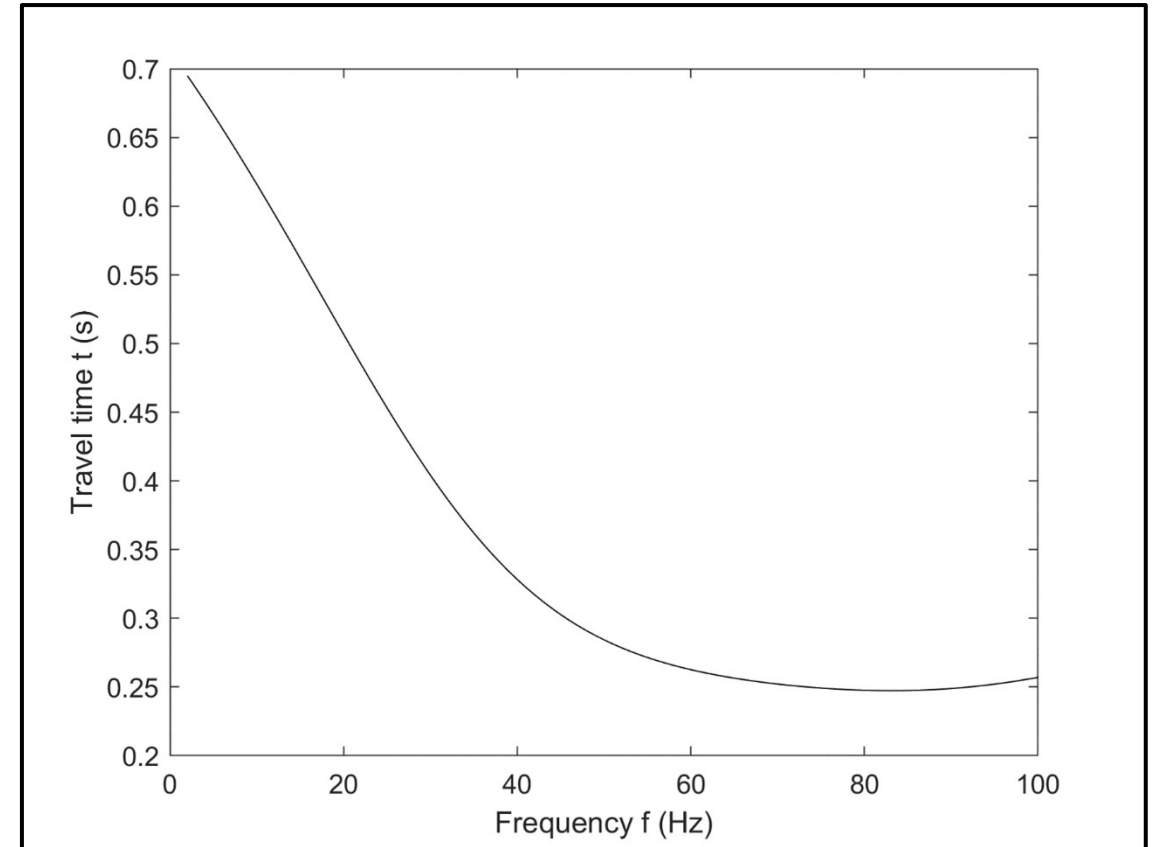




Travel time of the frequency



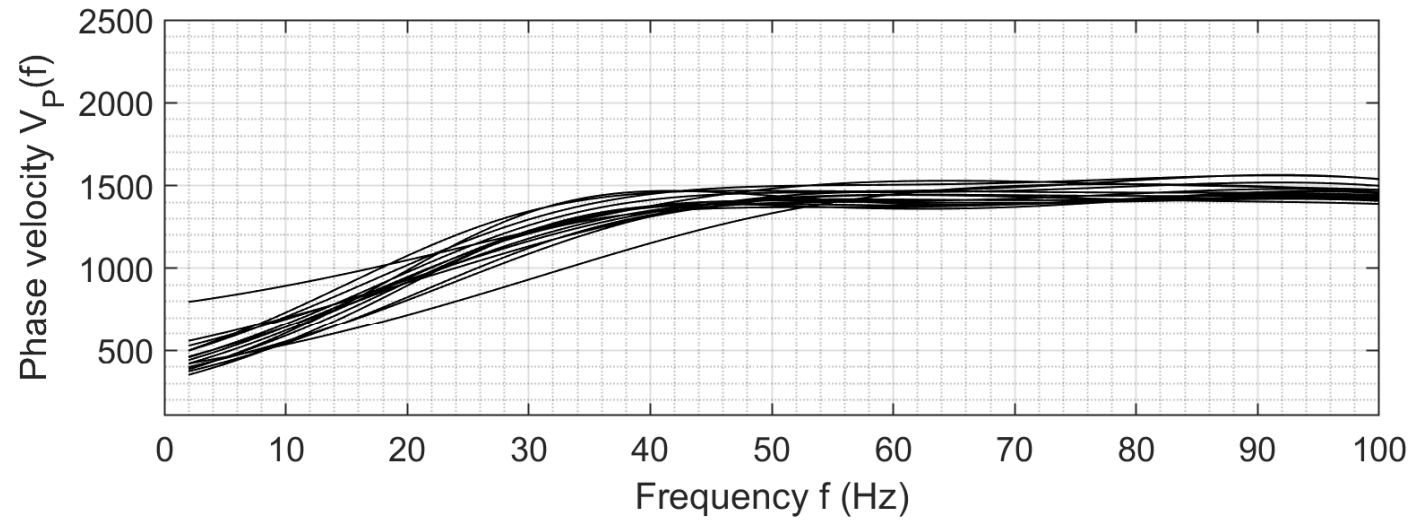
Overlay of arrival and departure time picks



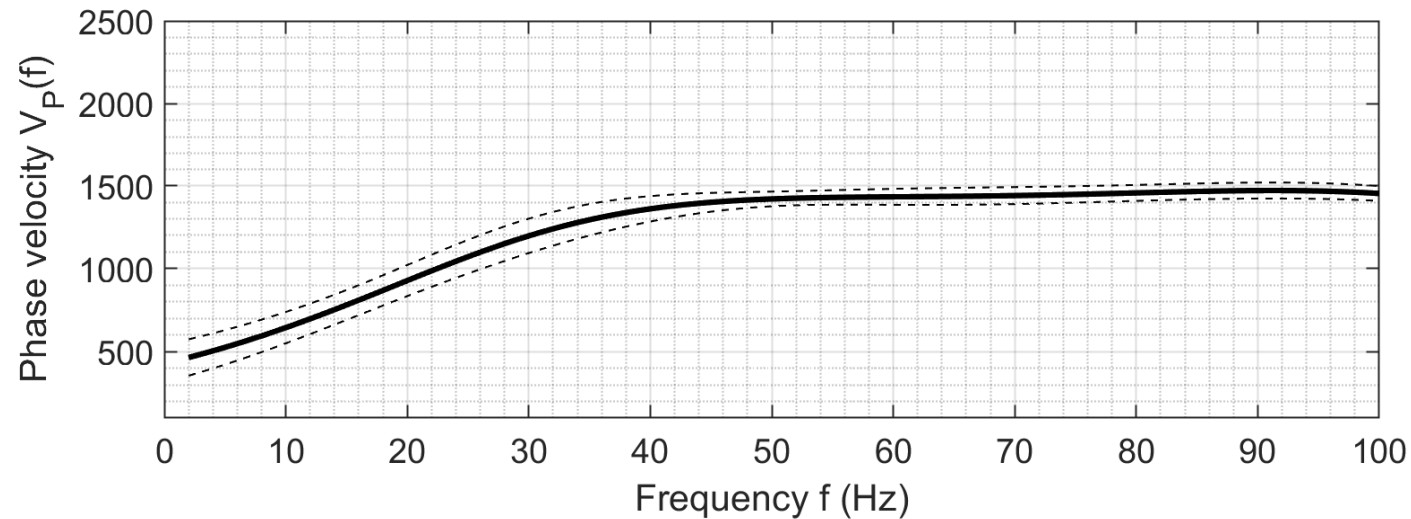
Travel time of the frequencies

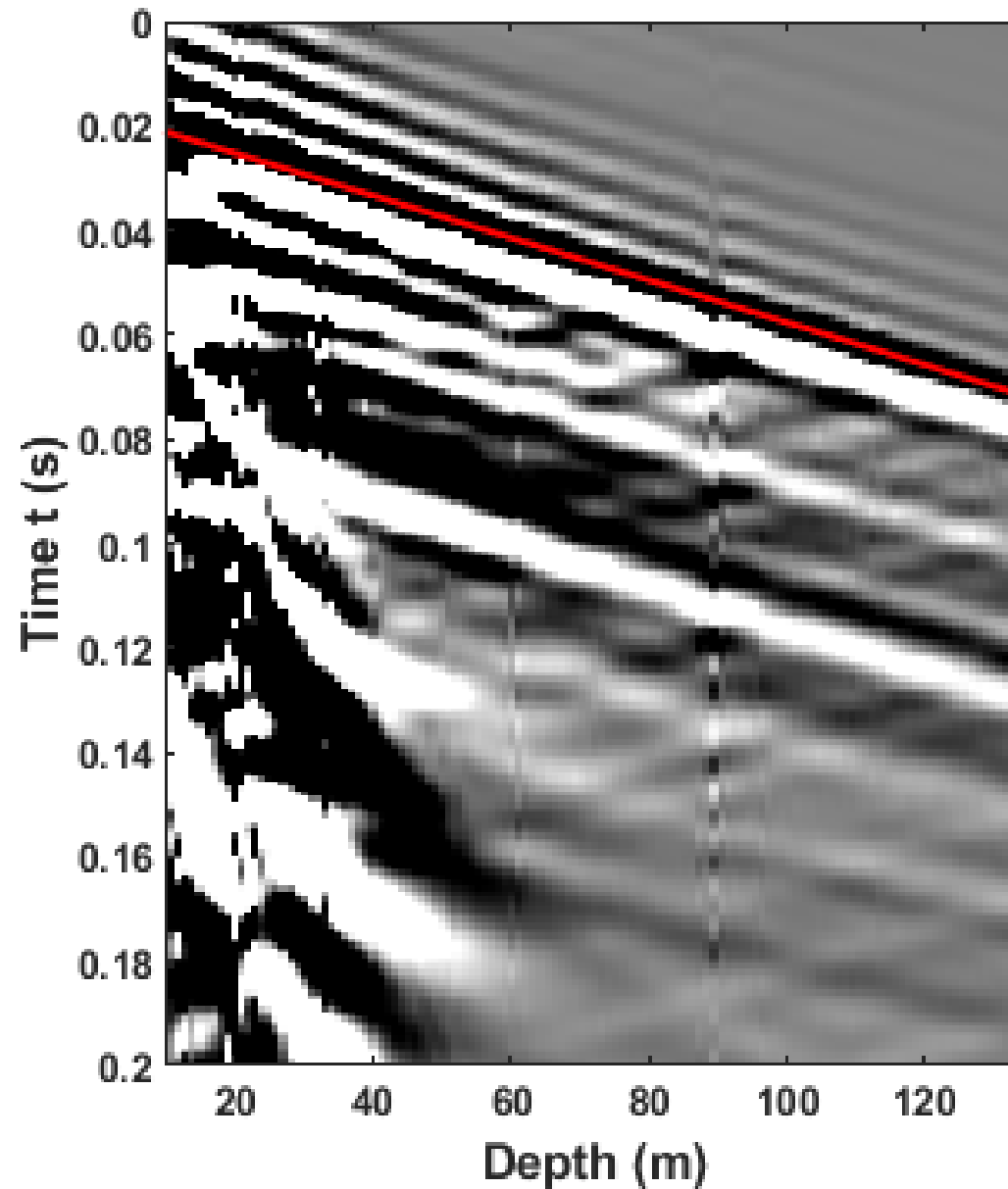


Estimating the uncalibrated(raw) phase velocity



Depth = 1-150m







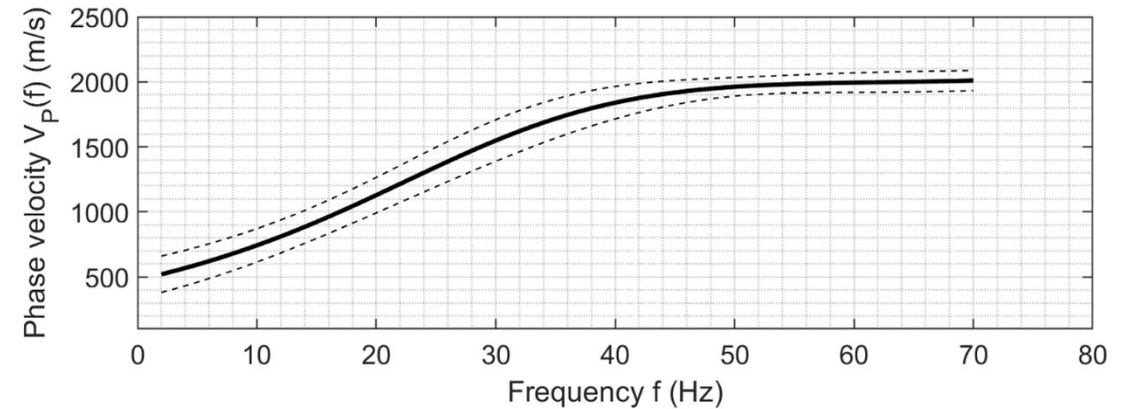
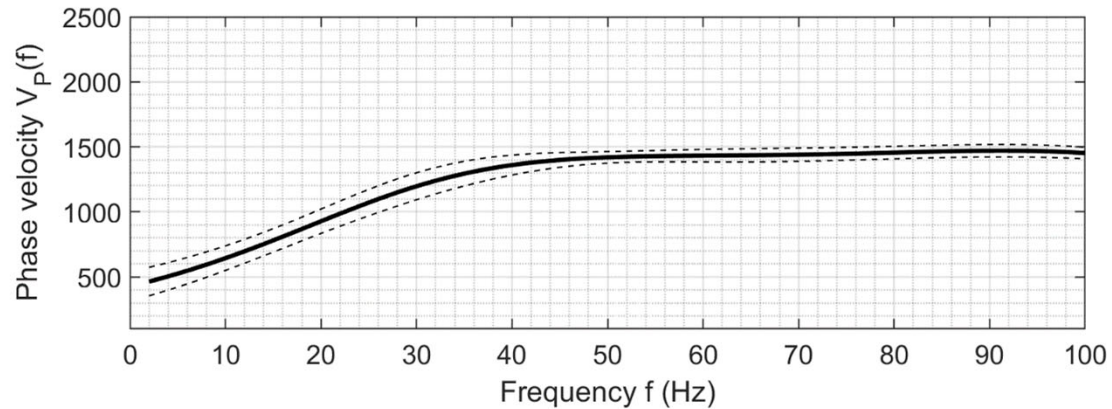
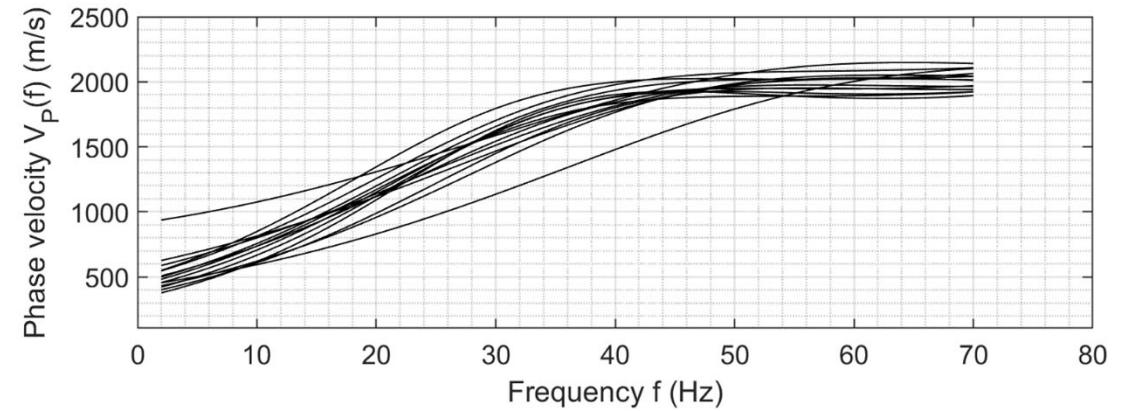
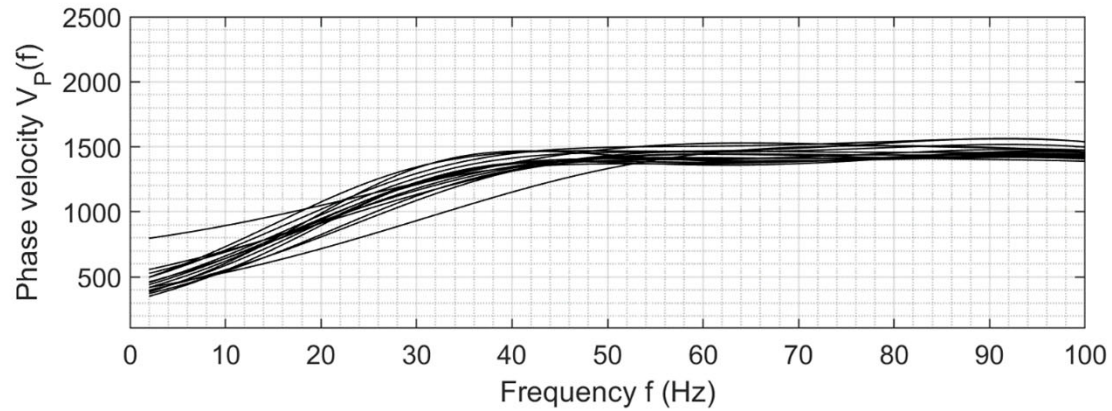
Estimation of calibrated phase velocity

$$V_{p(f)} = \frac{L}{T_{M(f)} - T_{s(f)} - T_{cal}}$$



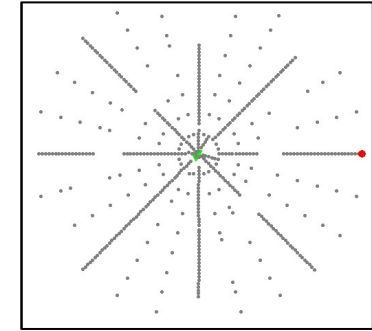
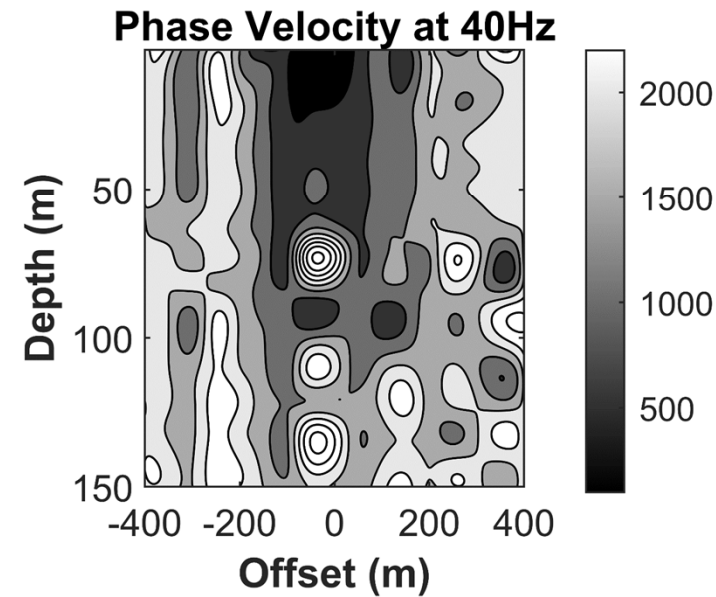
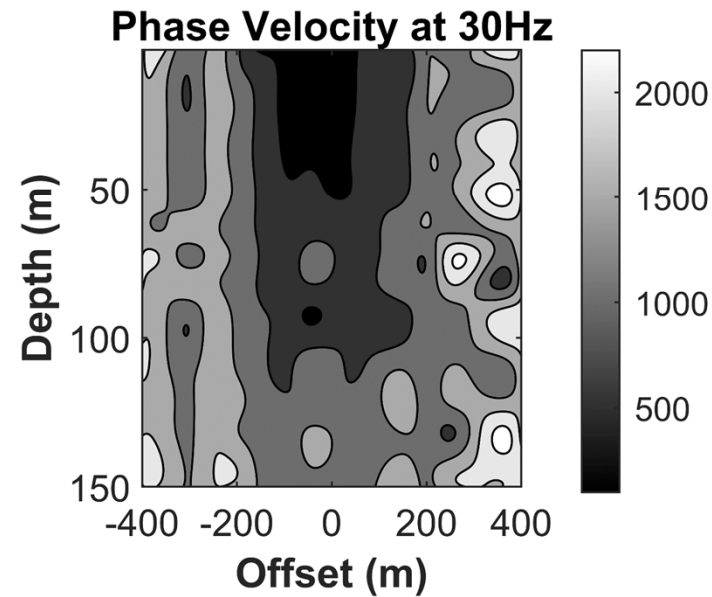
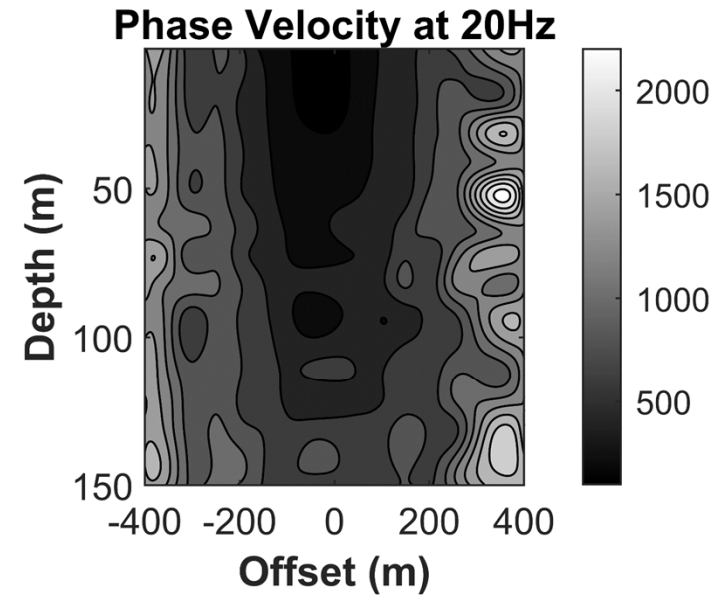
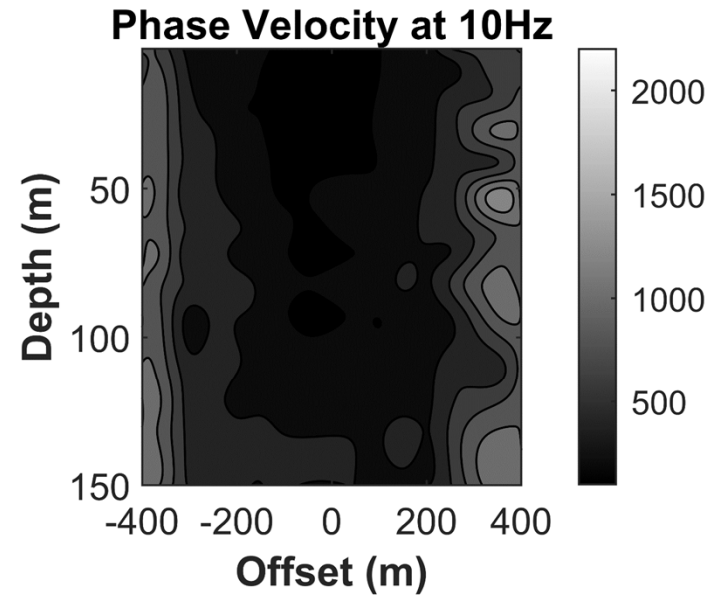
Uncalibrated phase velocity

Calibrated Phase Velocity



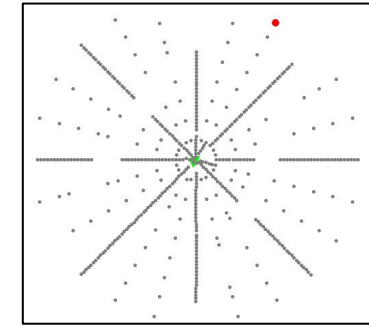
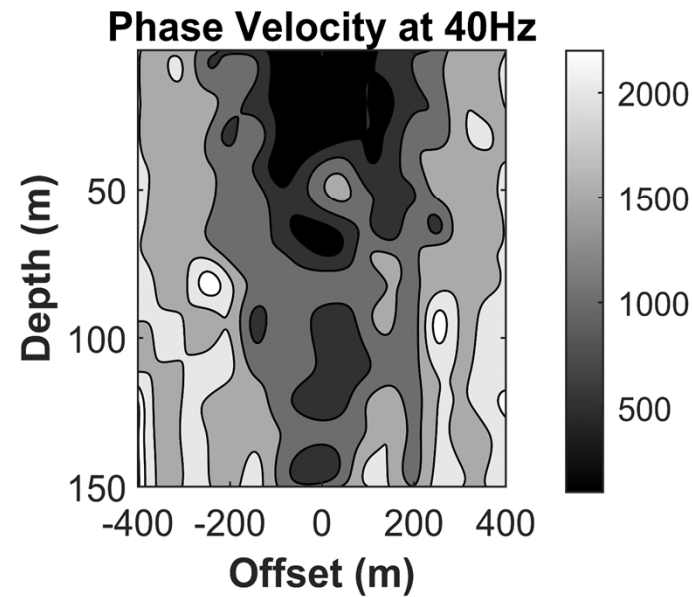
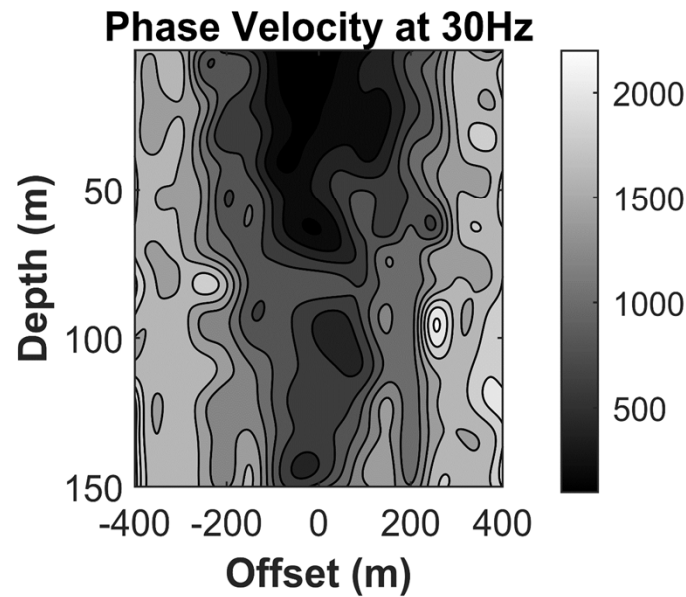
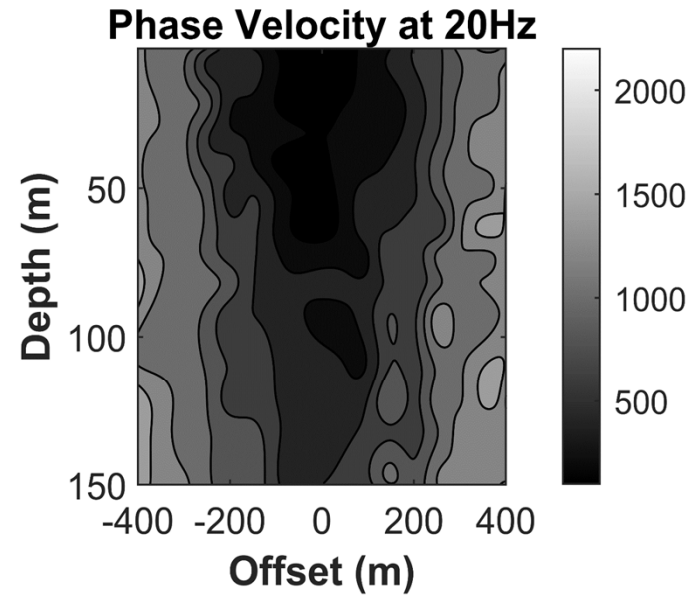
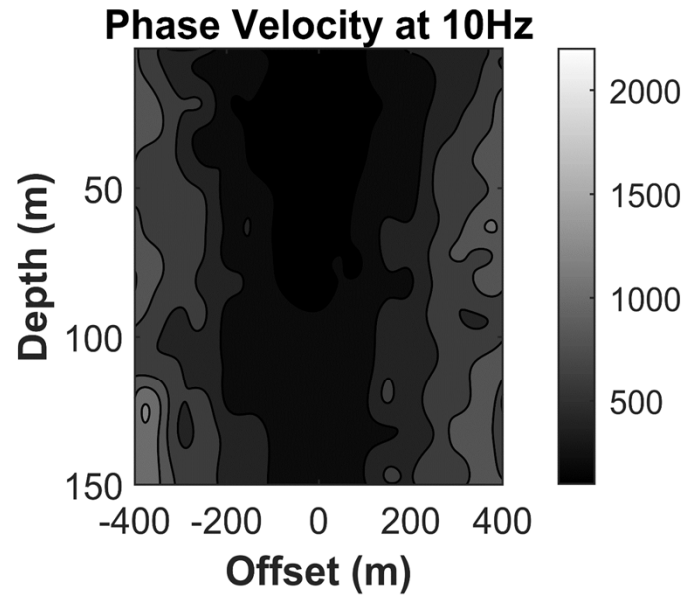


Phase velocity maps



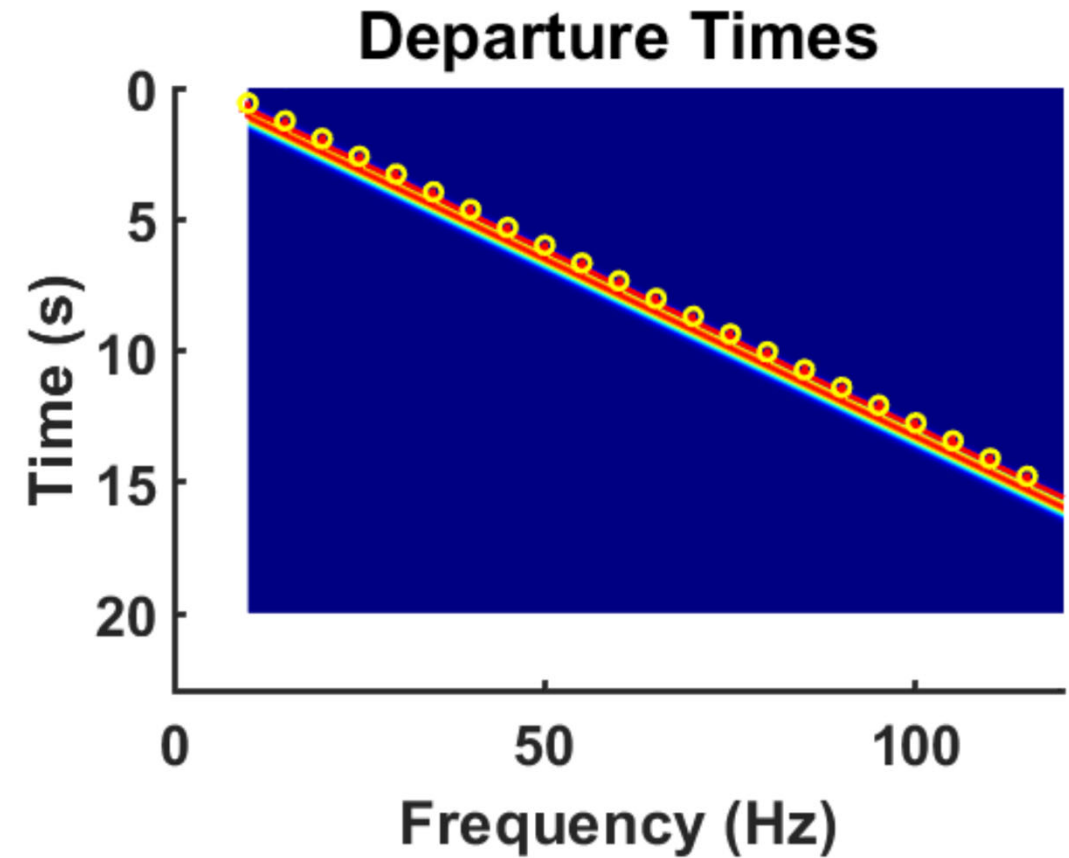
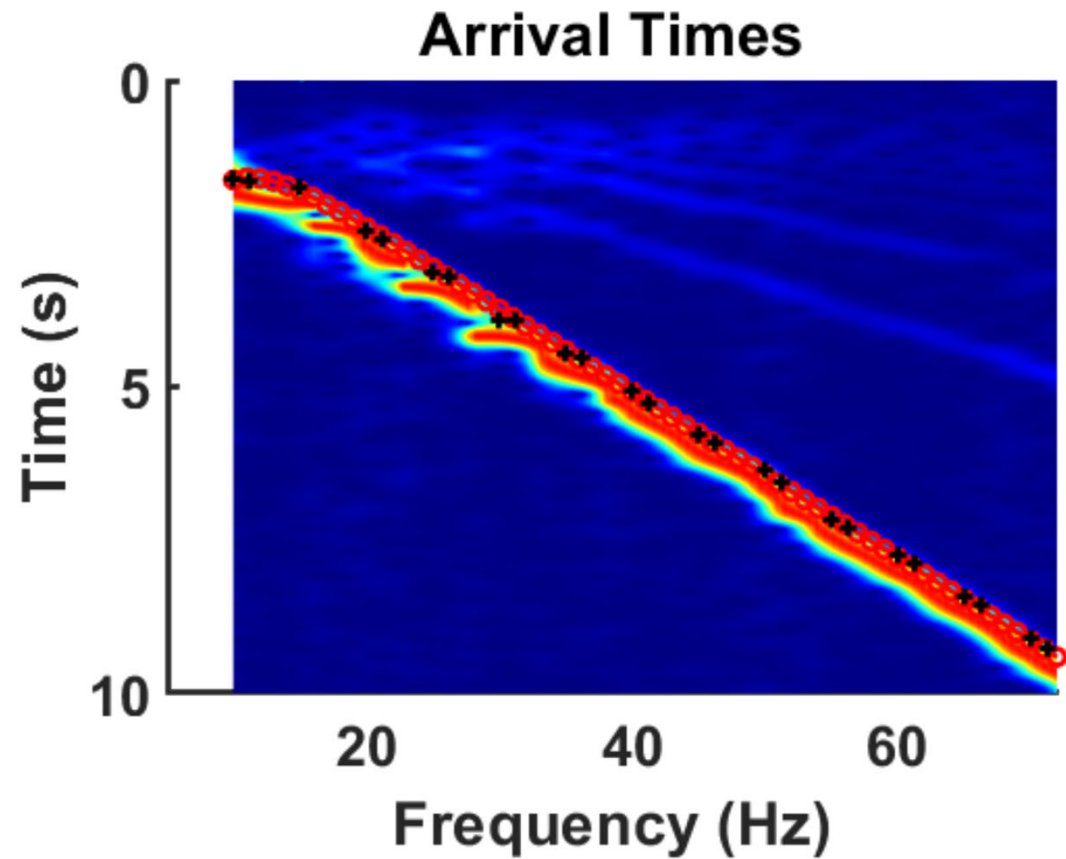


Phase velocity maps





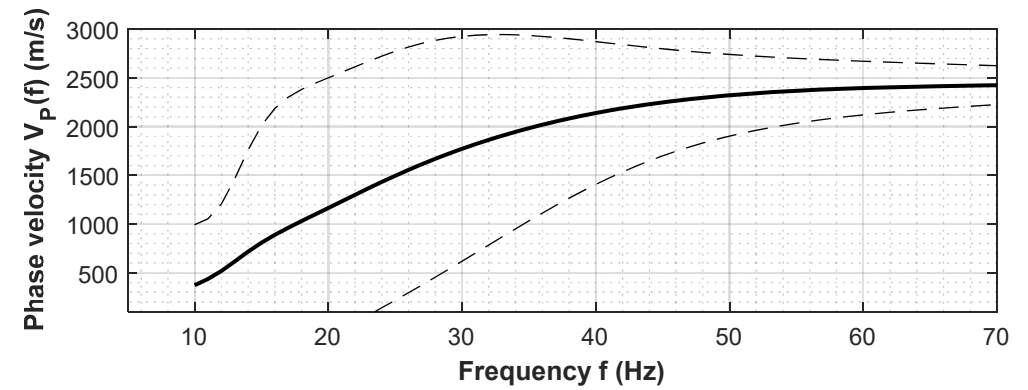
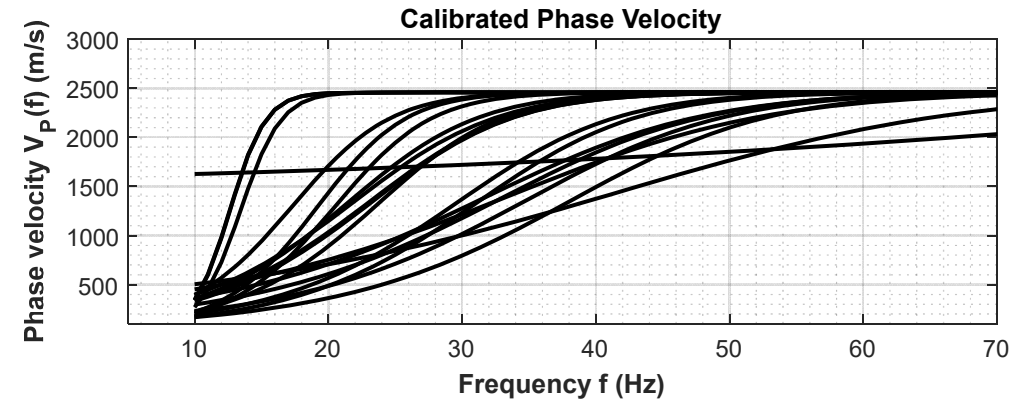
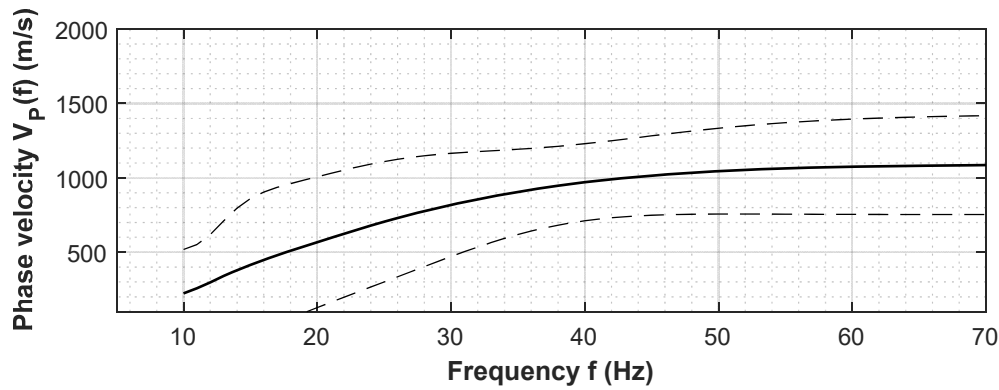
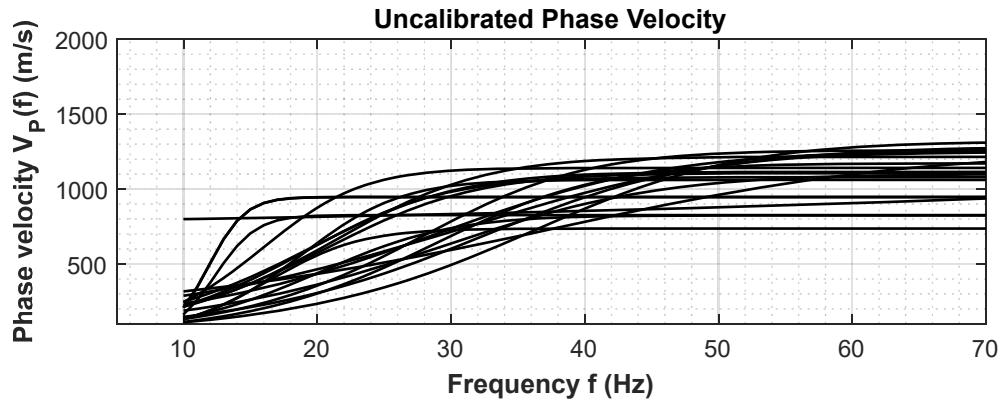
Automatic Picking procedure: Modified Energy Ratio(MER)





Estimating the calibrated phase velocity

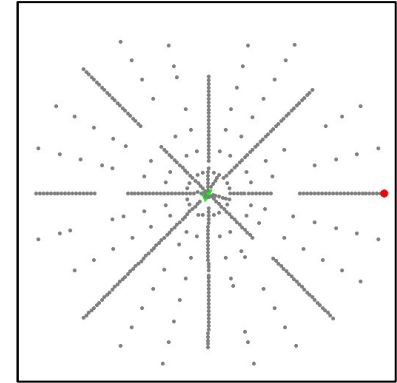
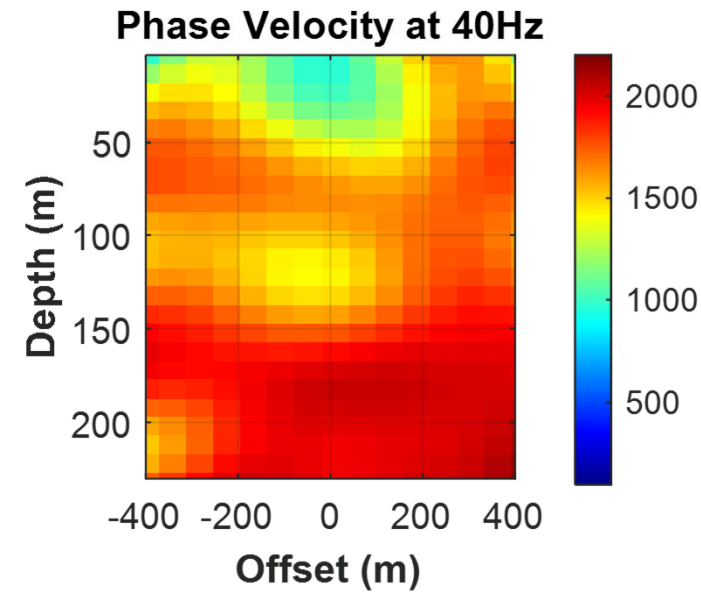
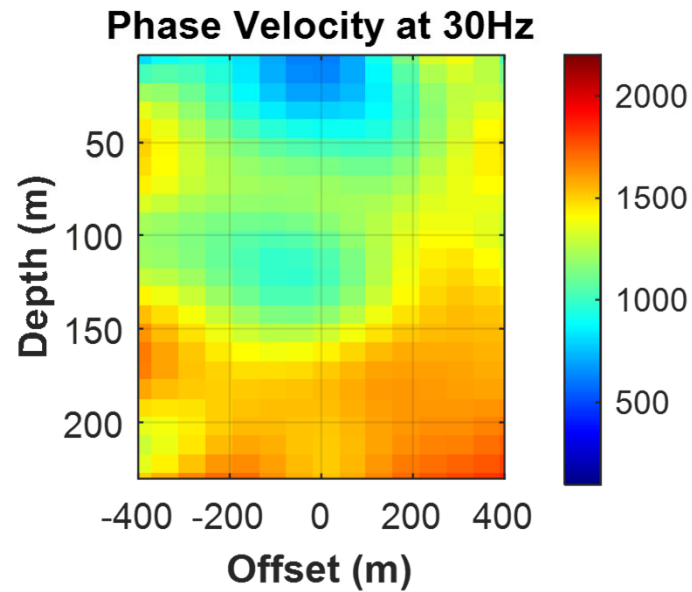
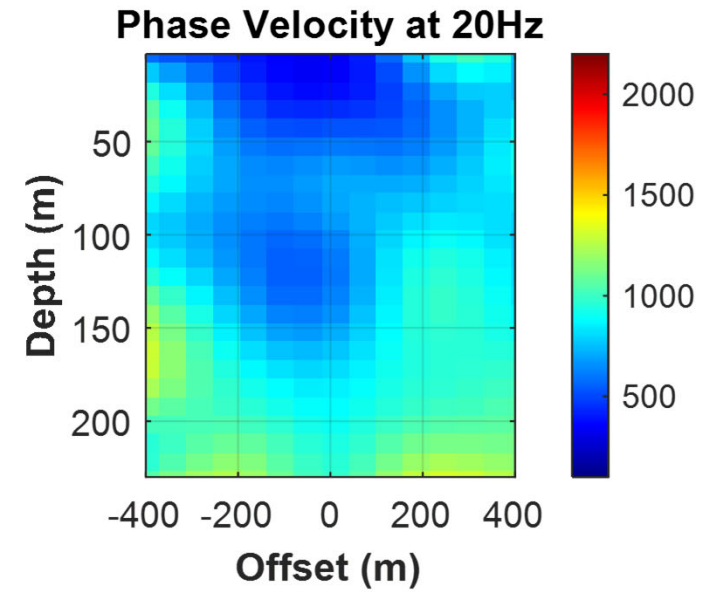
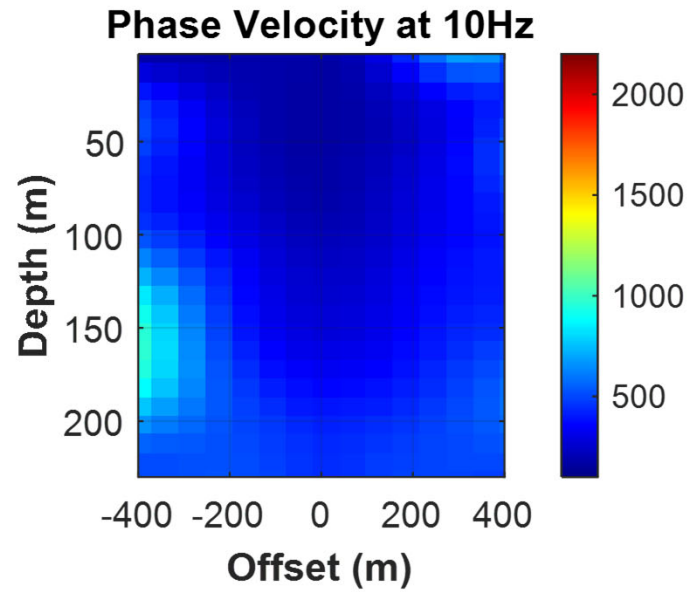
$$V_{p(f)} = \frac{L}{T_{M(f)} - T_{s(f)} - T_{cal}}$$



Depth 1-232m

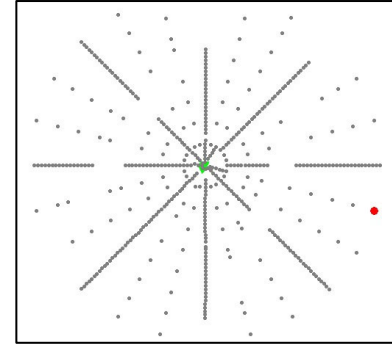
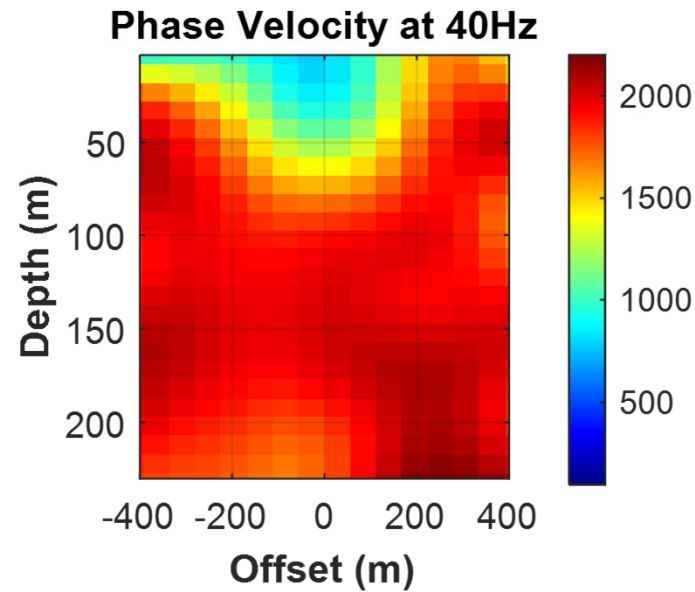
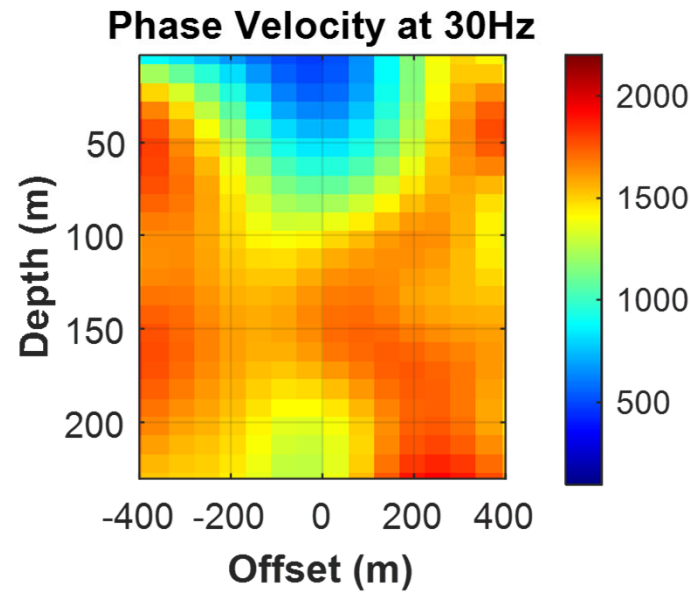
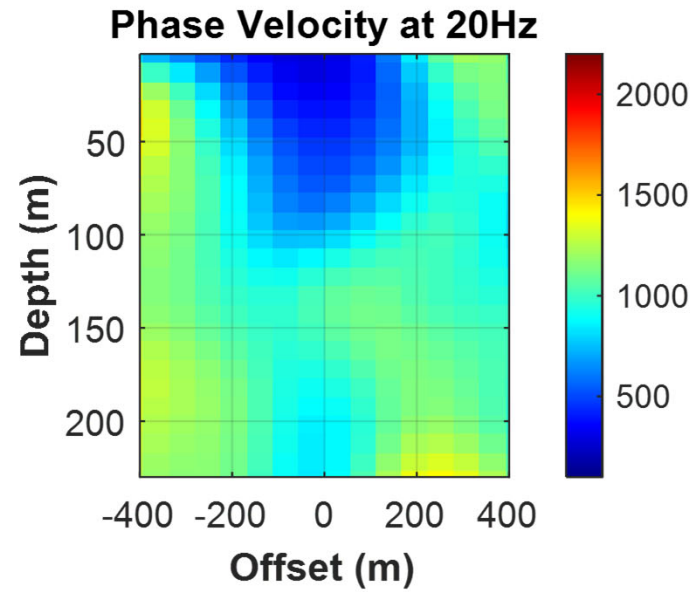
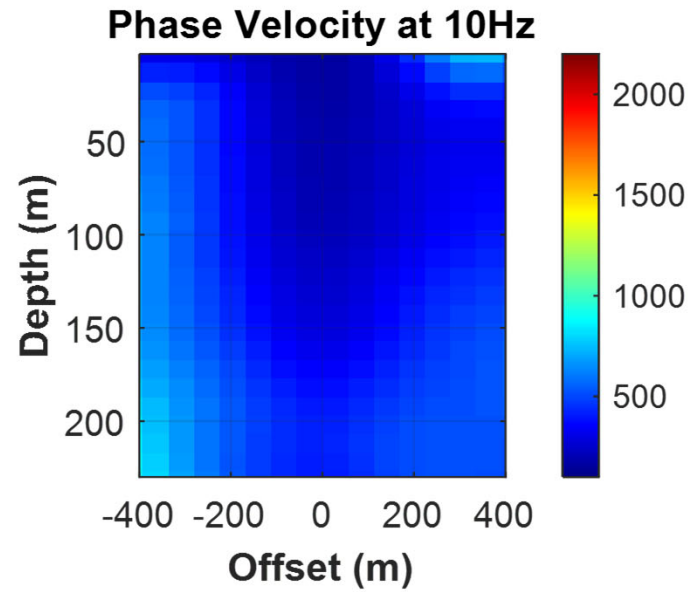


Phase velocity maps





Phase velocity maps





- Analyzing the VSP datasets from the Newell County facility enhances understanding of the complex near-surface.
- Phase velocity changes with VSP line azimuth; spatial heterogeneity in attenuation/dispersion.
- Estimating the frequency-dependent phase velocities in the near-surface supports determining the Q-factor.
- Phase velocity estimates could be used as initial models for FWI
- Next steps: connect with VSP-FWI; analyze S wave data



- CREWES sponsors, faculty, and staff
- NSERC