STATICS DECONVOLUTION VS INTERFEROMETRY

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Outline

- Motivation—similarity between techniques gives insight for further development
- Statics deconvolution—the basics
- Seismic interferometry—the basics
- Comparing the techniques on field data
- Conclusions and conjectures

Statics deconvolution

- Simple time shift model replaced by more general 'statics distribution function' model
- Statics distribution functions for seismic traces estimated from cross-correlations of raw traces with 'pilot traces'
- Match filter or inverse filter derived for each statics distribution function
- Application of unique filter corrects each trace



Raypath segments beneath surface points not vertical; Sources and receivers can be arrays, with different statics for each point in the array. Multiple raypaths possible between source and receiver location (P_1 and P_2), due to buried velocity anomalies (V_3)

Ideal seismic trace Actual statics distribution function Input seismic

- trace
- Estimated statics distribution function Desired statics distribution function
- **Match filter**
- Match filtered seismic trace
- Ideal seismic trace



The static deconvolution principle

Statics deconvolution

$$\mathbf{T}_{k\beta}(t) = w(t) * s_{k}(t) * R_{k\beta}(t) * r_{\beta}(t)$$

Raw seismic trace

$$\hat{\mathrm{T}}_{\alpha\beta}\left(t\right) = \mathrm{T}_{\alpha\beta}\left(-t\right) * \mathrm{T}_{\alpha\beta}\left(t\right) * \sum_{k=1}^{N} \mathrm{T}_{k\beta}\left(t\right).$$

Corrected seismic trace

Seismic interferometry

- The principles of reciprocity and time reversal are used to create 'virtual traces' from recorded traces
- Time-reversed portion of a raw trace is used as a filter to remove surface-related effects
- Sum of filtered traces over an aperture creates a new trace with a 'virtual' source and no surface-related effects



Raw trace pairs from receiver gathers at A and B (common source k) cross-correlated to cancel source phase $s_k(t)$. Cross-correlations convolved (as match filters) with traces of receiver gather A to cancel receiver phase $r_A(t)$, then summed over shot aperture N to approximately cancel residual source phase $s_k(t)$. This is a 'virtual trace', between A and B for a virtual source (receiver) gather.

Seismic interferometry

$$\mathbf{T}_{k\beta}(t) = w(t) * s_{k}(t) * R_{k\beta}(t) * r_{\beta}(t)$$

Raw seismic trace

$$\hat{\mathbf{T}}_{\alpha\beta}(t) = \sum_{k=1}^{N} \left[\mathbf{T}_{k\alpha}(-t) * \mathbf{T}_{k\alpha}(t) \right] * \mathbf{T}_{k\beta}(t)$$

Virtual seismic trace

How they're related

$$\hat{\mathrm{T}}_{\alpha\beta}\left(t\right) = \mathrm{T}_{\alpha\beta}\left(-t\right) * \mathrm{T}_{\alpha\beta}\left(t\right) * \sum_{k=1}^{N} \mathrm{T}_{k\beta}\left(t\right).$$

Statics deconvolution

$$\hat{\mathbf{T}}_{\alpha\beta}(t) = \sum_{k=1}^{N} \left[\mathbf{T}_{k\alpha}(-t) * \mathbf{T}_{k\alpha}(t) \right] * \mathbf{T}_{k\beta}(t)$$

Seismic interferometry

Hansen Harbour field example

- Receiver spread only 50 stations (15m interval), no appreciable surface functions
- Source spread 200 stations (30m interval), visible variations in coupling, statics
- Raw receiver gathers show source statics and coupling variations, as well as coherent noise



Hansen Harbour 3C seismic line geometry

2 sec



Hansen Harbour receiver gather, bandpass, AGC



Hansen Harbour stack—no filtering, no statics





Hansen Harbour receiver gather after coherent noise attenuation and statics deconvolution



Hansen Harbour stack—coherent noise attenuated, statics deconvolved



Hansen Harbour virtual receiver gather from raw traces, no noise attenuation

2 sec



Hansen Harbour stack—virtual receiver gathers, no noise attenuation

Conclusions and conjectures

- Statics deconvolution and interferometry obviously *similar* in their effects
- Cross-correlation or match-filtering is the key to removing phase effects of surface functions
- Coherent noise attenuation an unexpected benefit of interferometry
- Further exploration of interferometry for correcting surface effects is warranted

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