

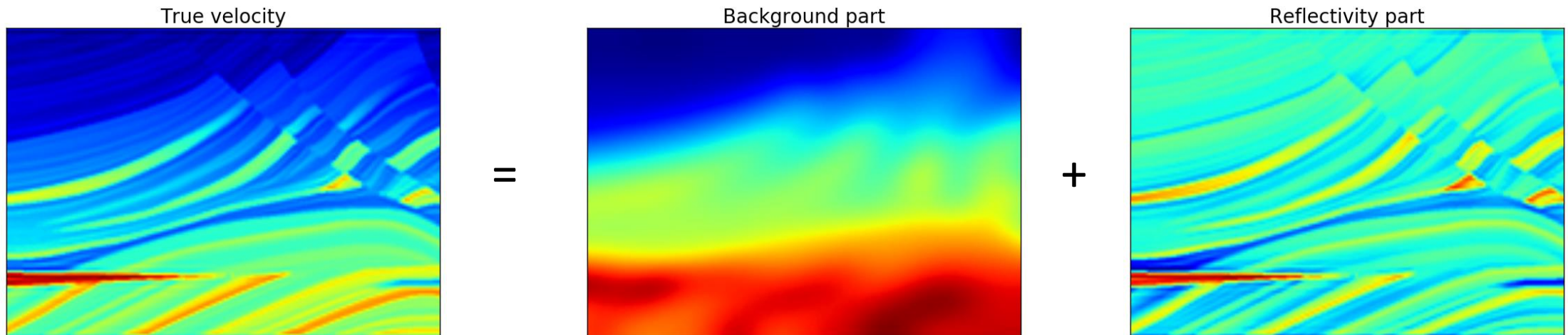
Evaluating the potential of reflection-based waveform inversion (RWI)

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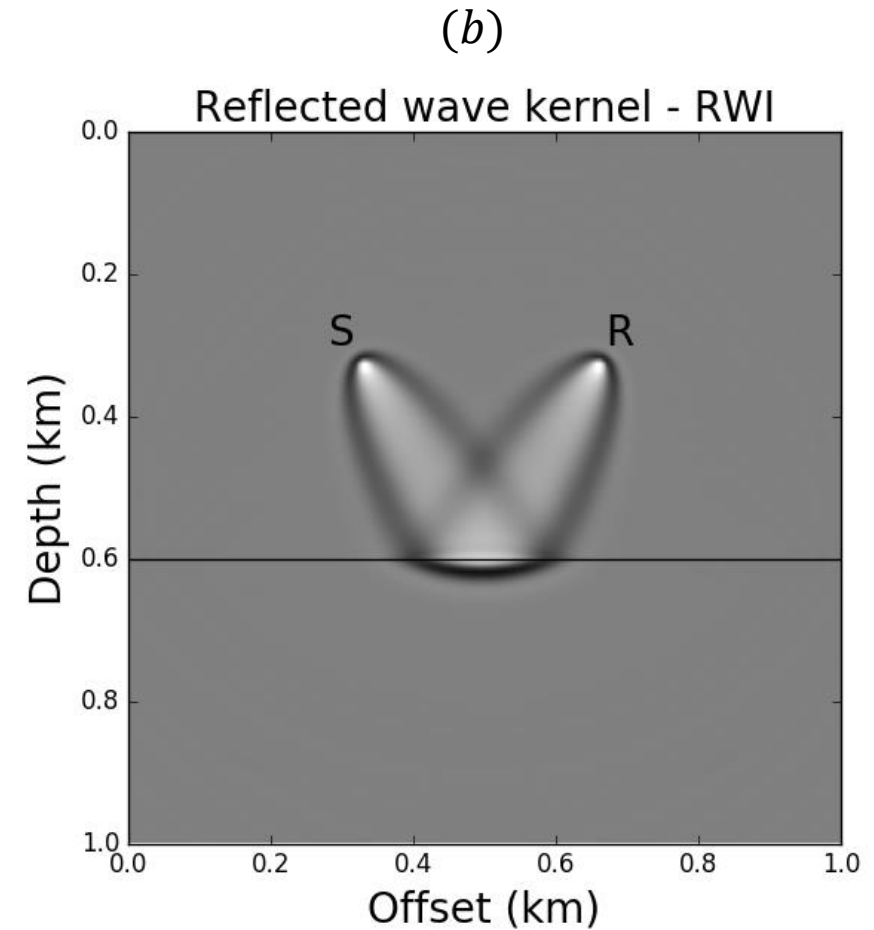
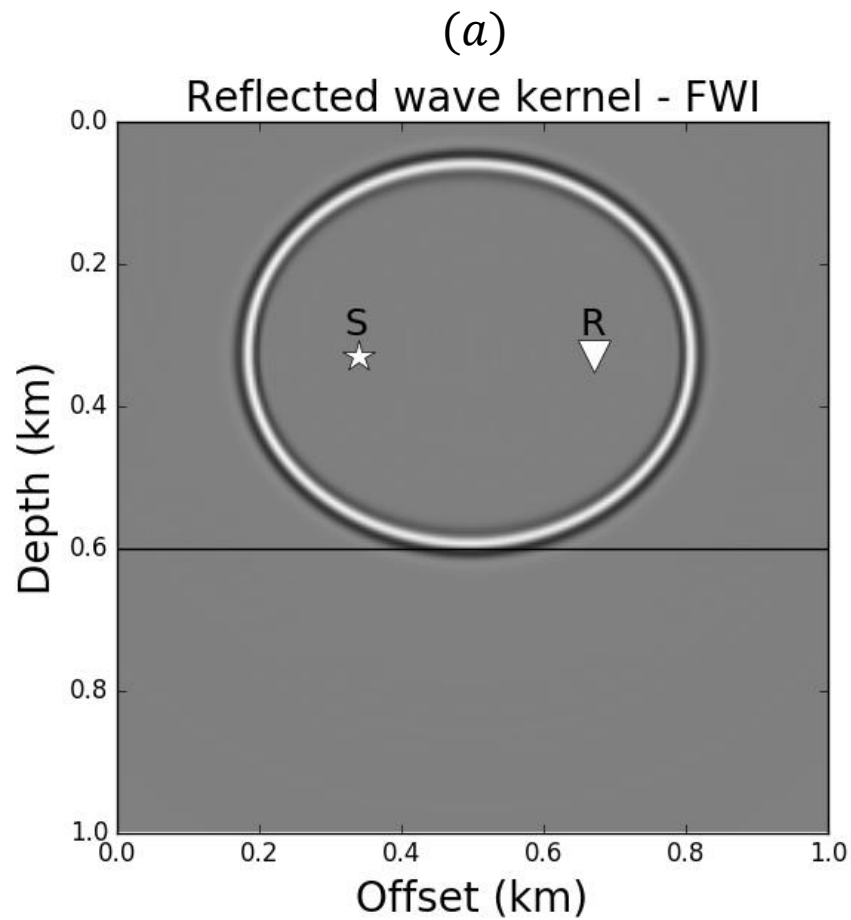
- Motivation
- Modeling by seismic demigration
- FWI workflow
- RWI workflow
- Example
- Issues with RWI
- Conclusion

- The velocity of the earth can be separated into a background velocity model (long-wavelength) and a reflectivity model (short-wavelength).

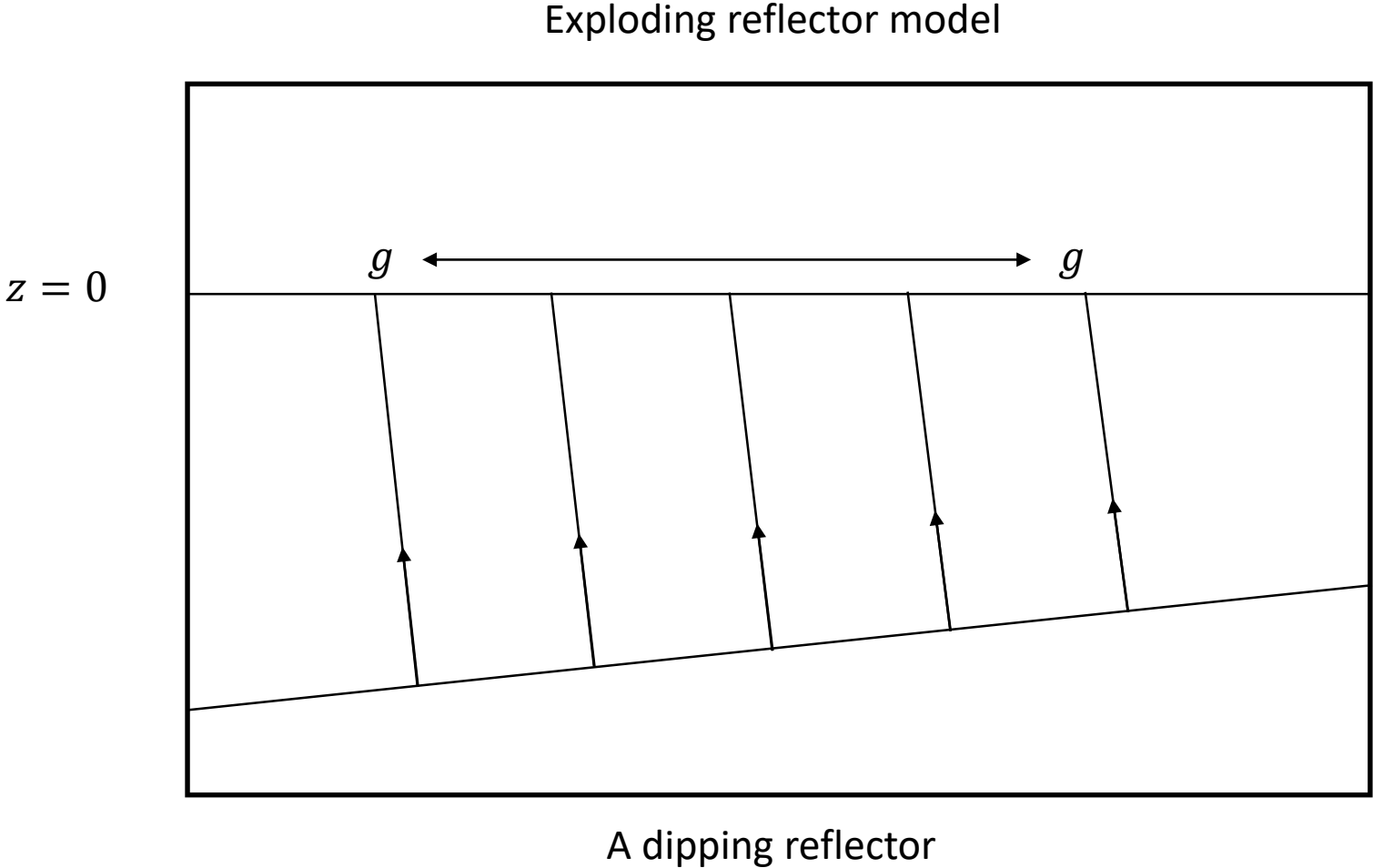


- FWI relies on the presence of low-frequency and long-offset data to recover the long-wavelength components of the velocity model.
- FWI recovers the long-wavelength components of the velocity model for the shallow section but it fails to recover it for the deep section.
- RWI aims at recovering the long-wavelength components of the velocity model in the deep part by using the transmission wavepaths of the reflection data.

Motivation – Reflected waves contribution to model update



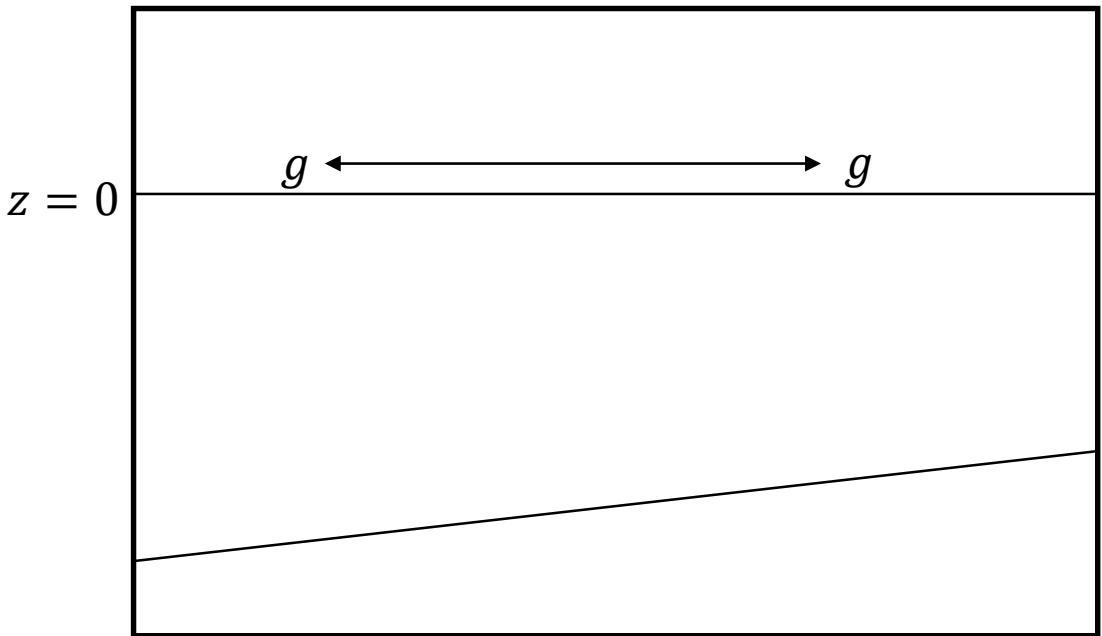
Seismic demigration – The exploding reflector model



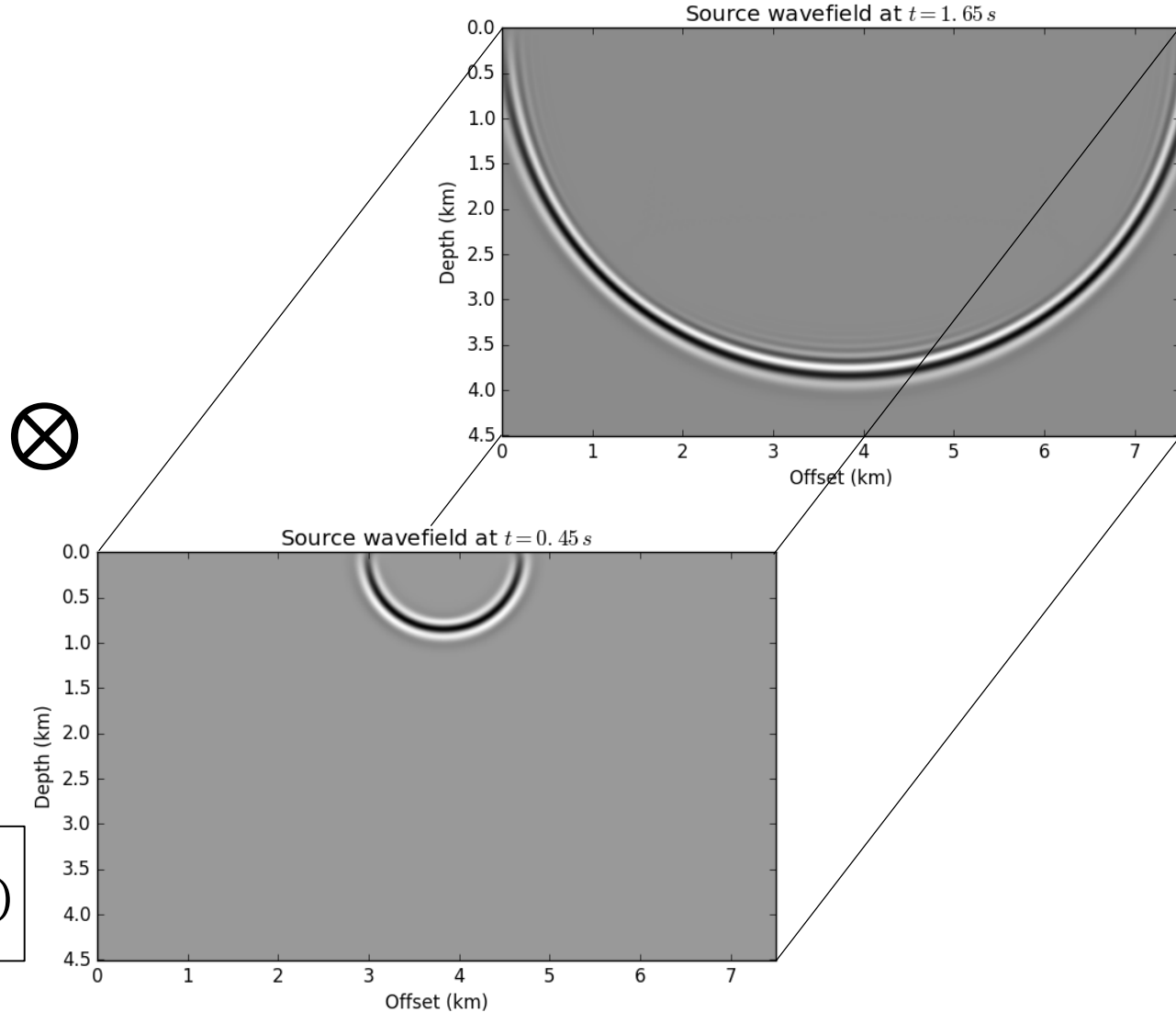
Seismic demigration

$$I(\vec{x})$$

Migrated Image in Depth

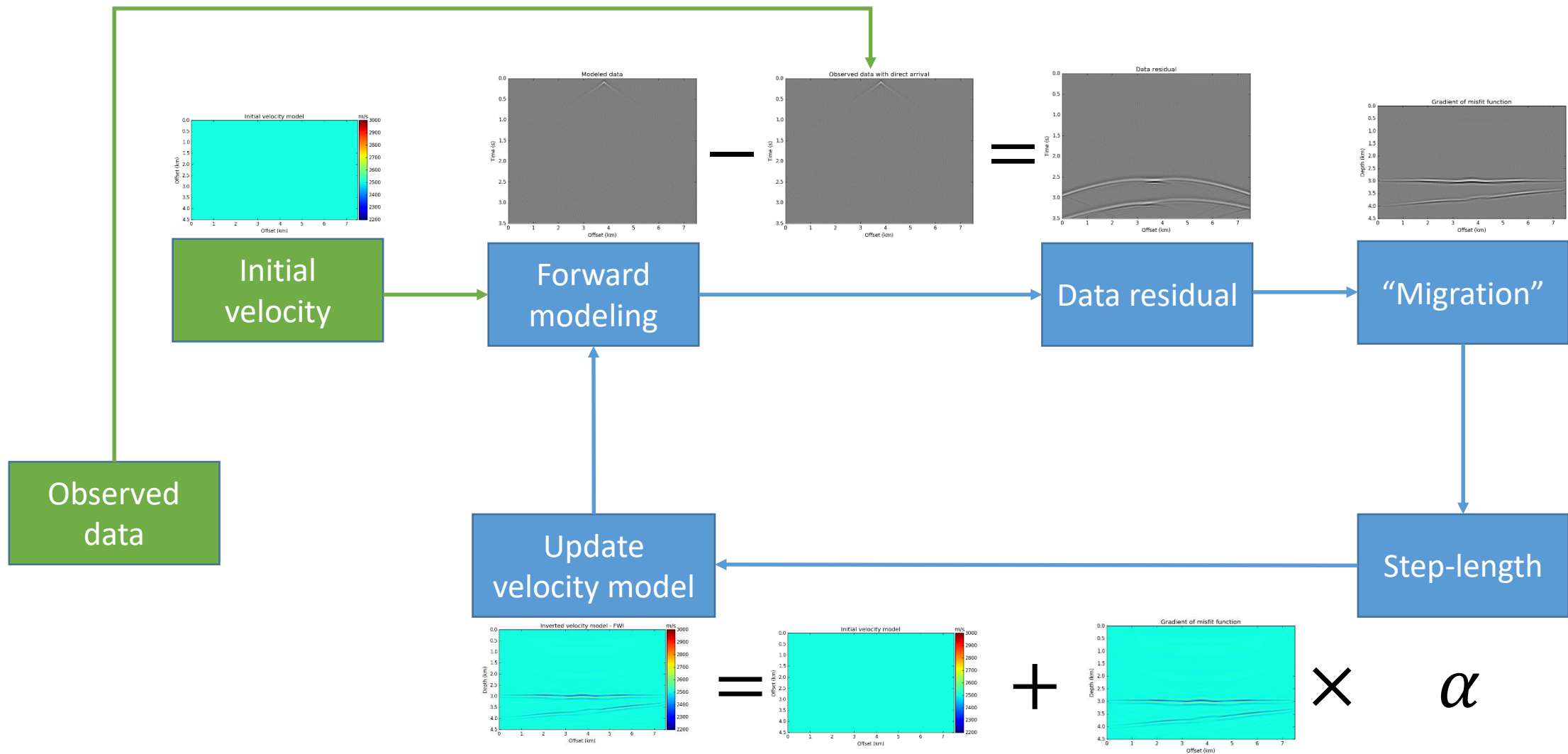


Source wavefield $U_s(\vec{x}, t; \vec{x}_s)$

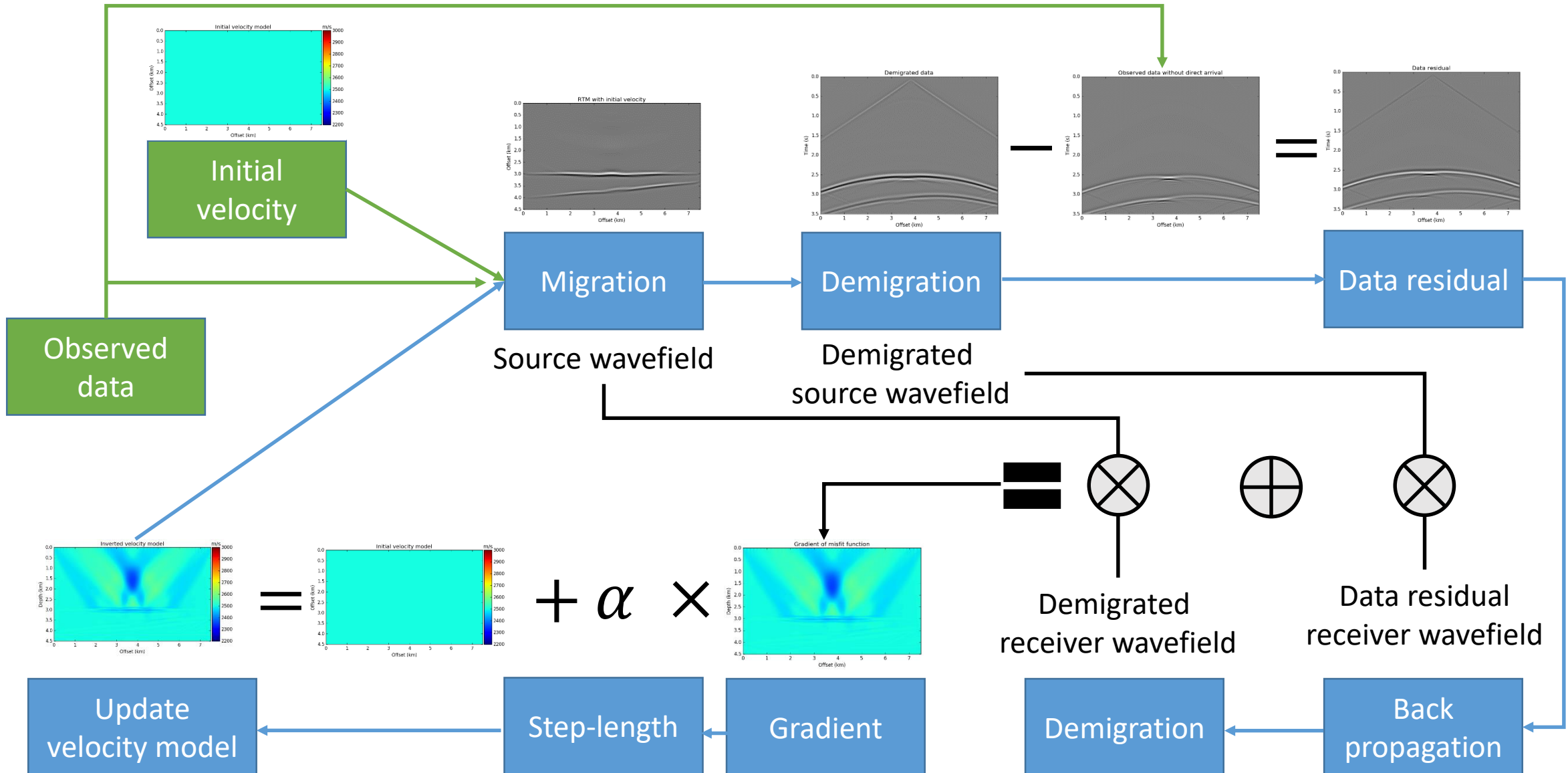


$$\left(\frac{1}{v(x)^2} \frac{\partial^2}{\partial t^2} - \nabla^2 \right) \delta U_s(\vec{x}, t; \vec{x}_s) = I(\vec{x}) \cdot U_s(\vec{x}, t; \vec{x}_s)$$

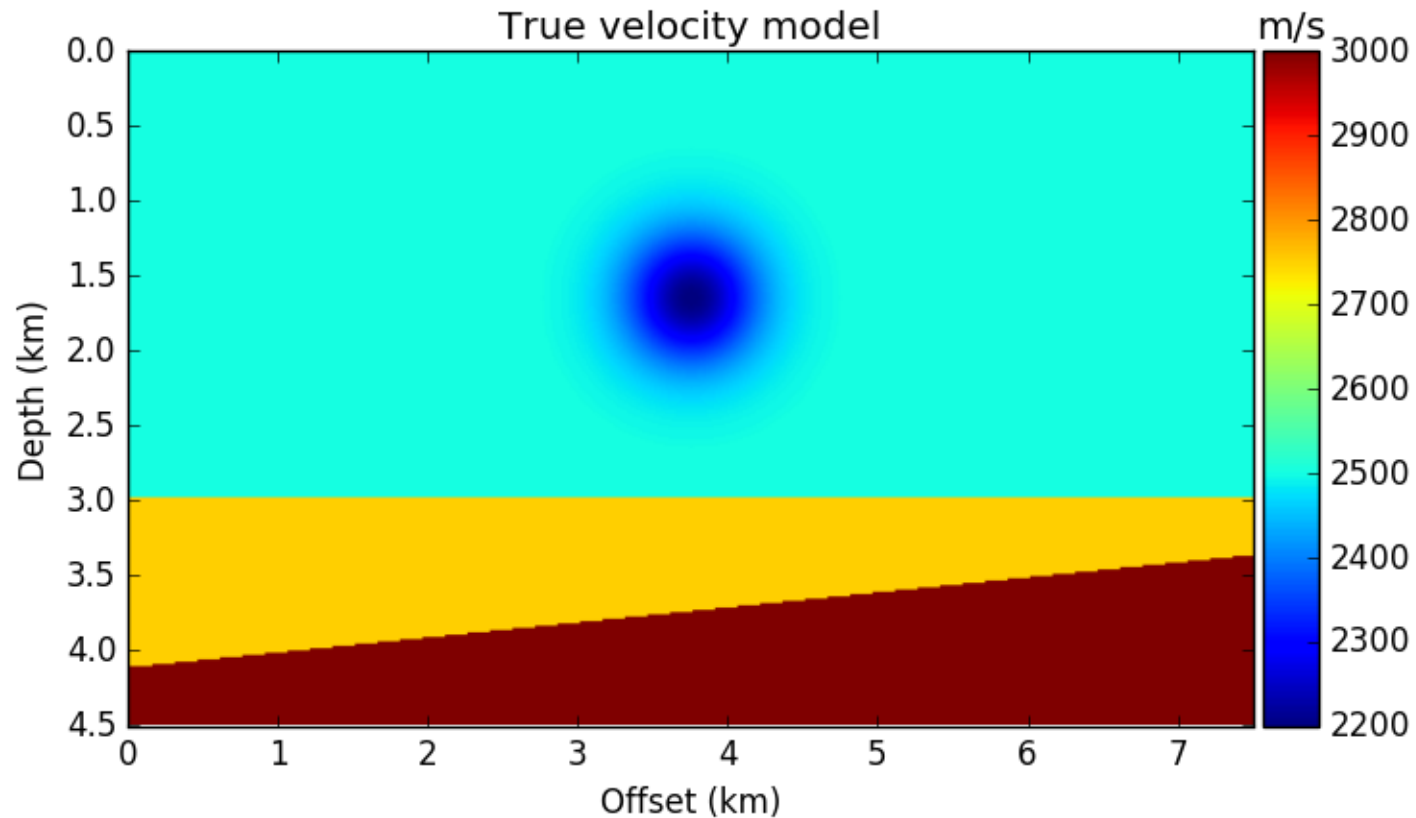
FWI workflow



RWI workflow

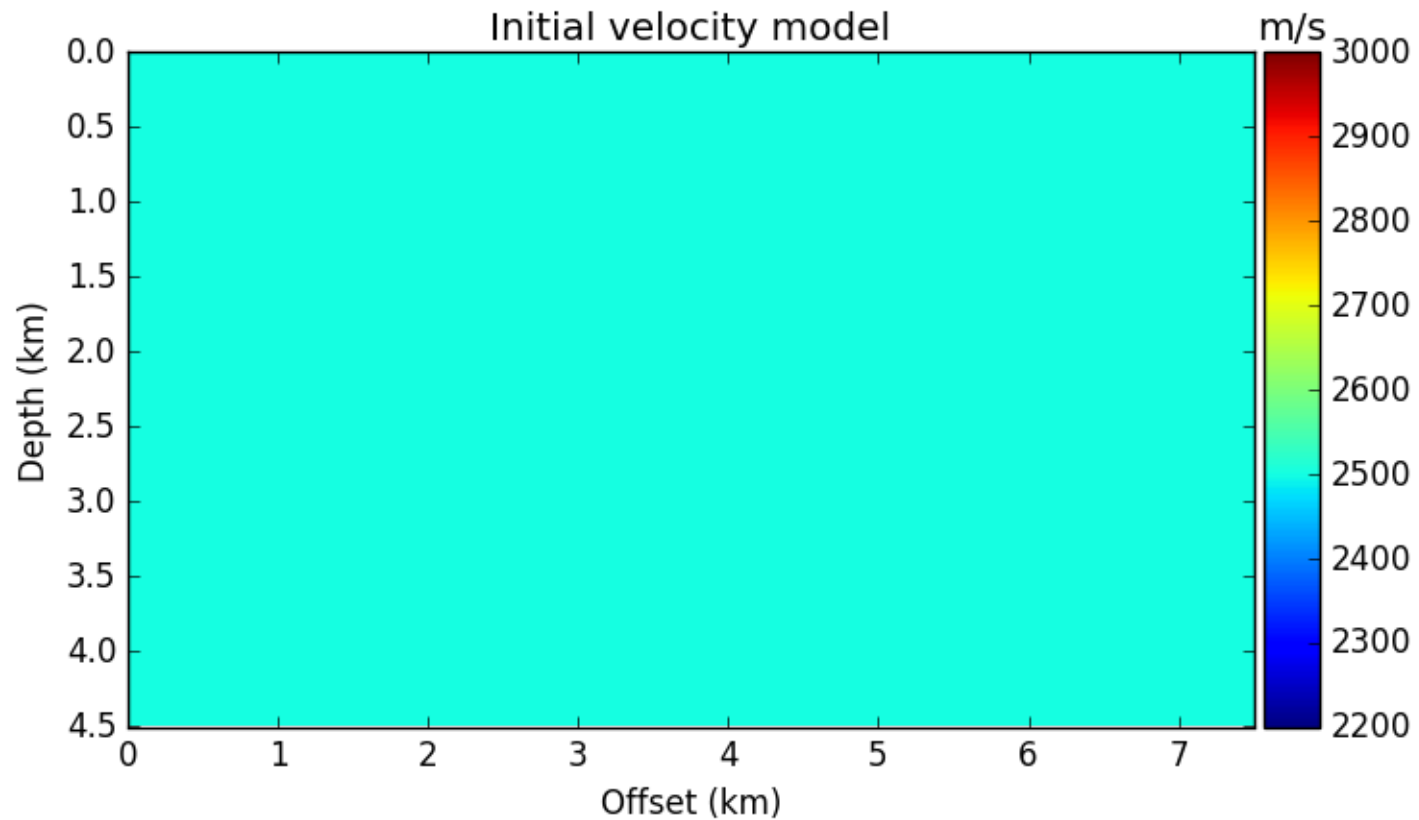


Example: true velocity model

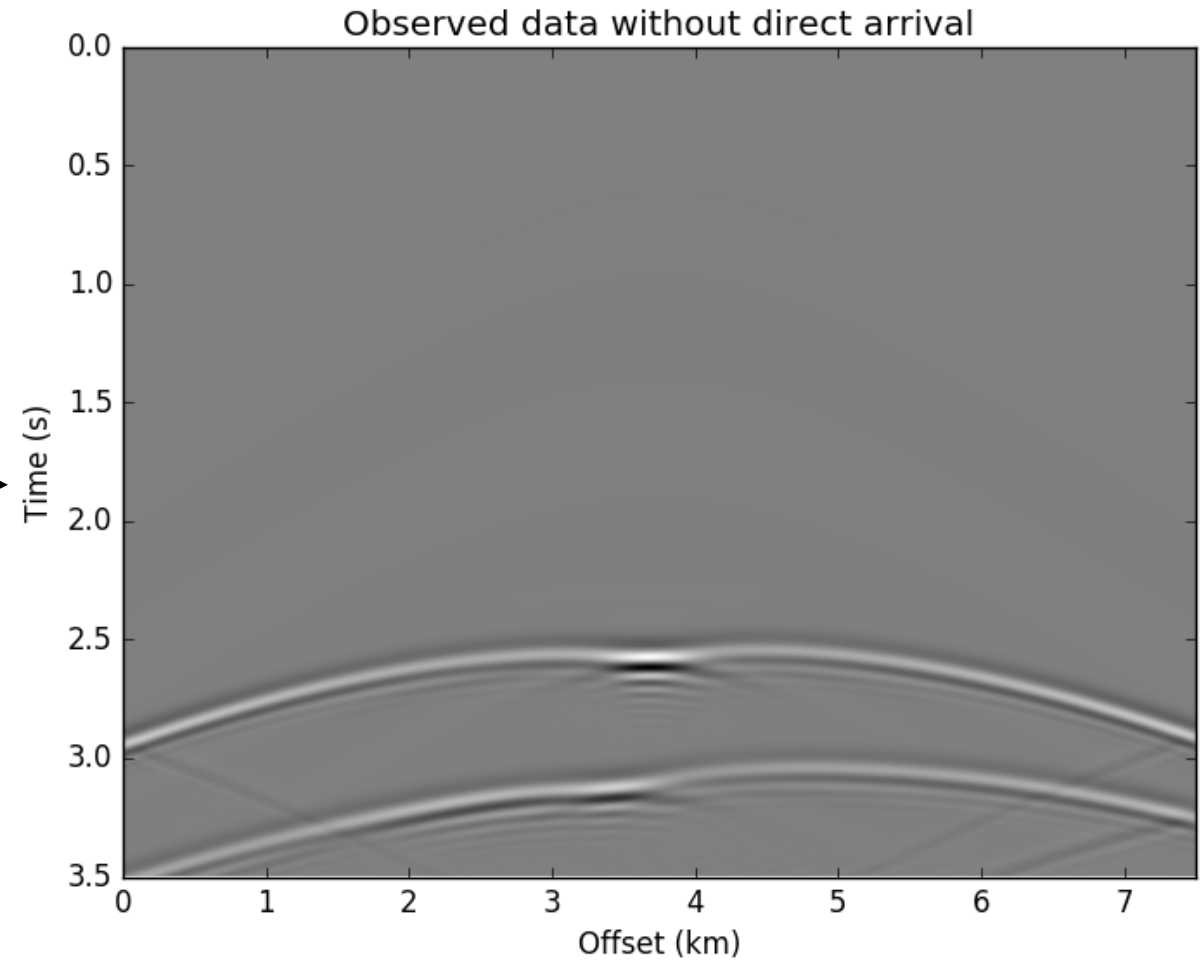
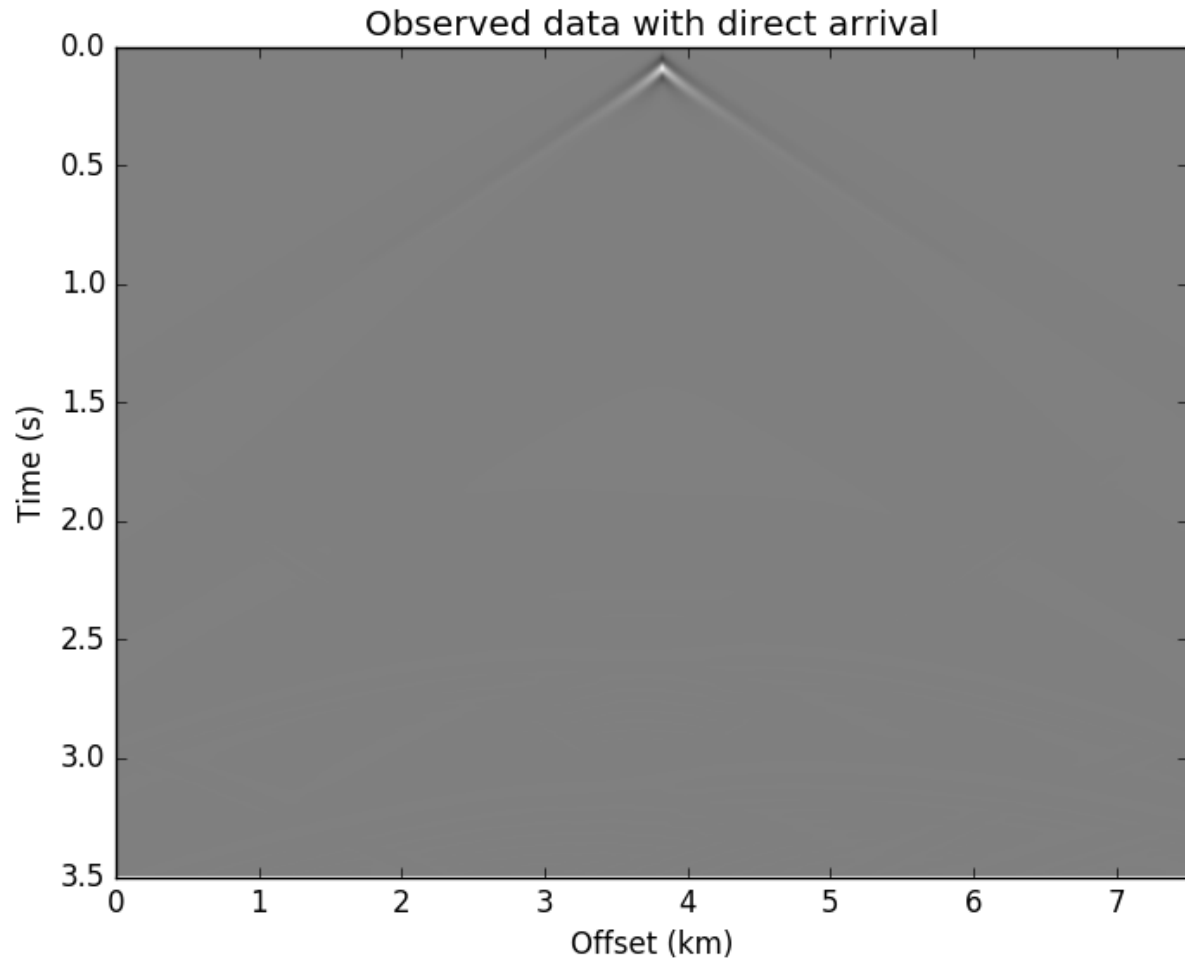


- Background velocity: 2500 m/s
- Center of lens velocity: 2200 m/s
- Source wavelet: Ricker with 10 Hz dominant frequency

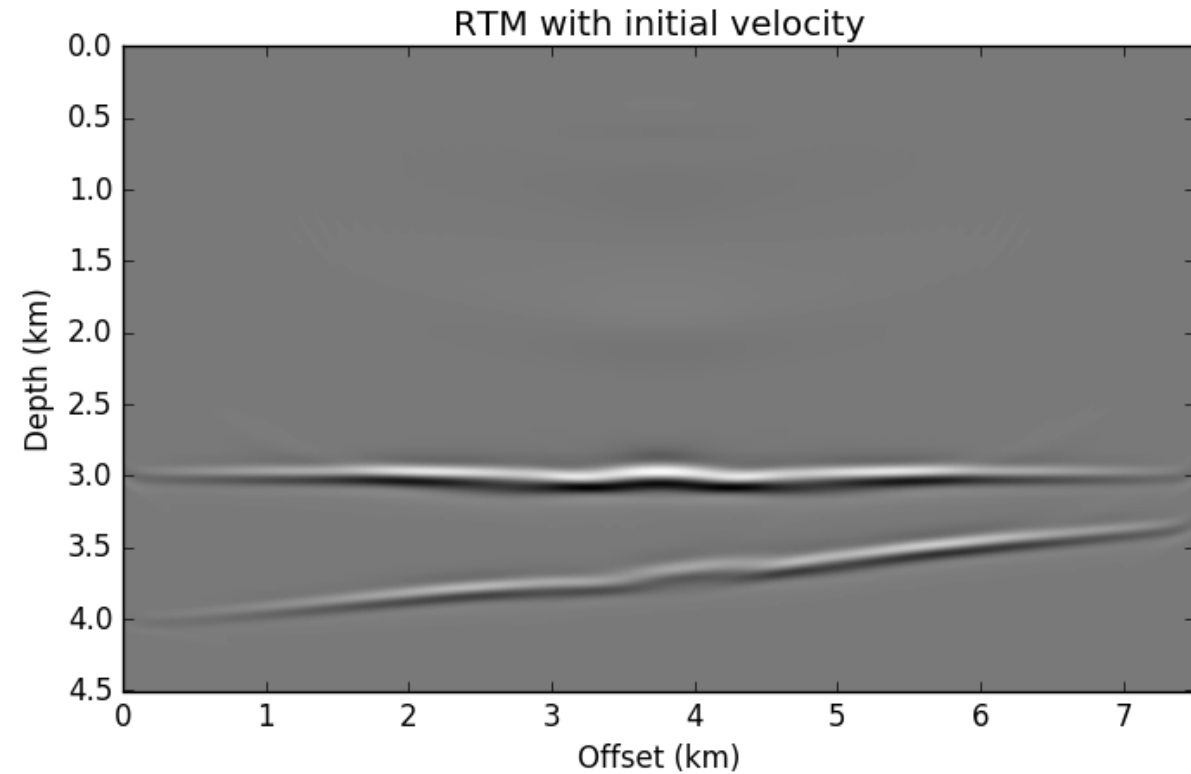
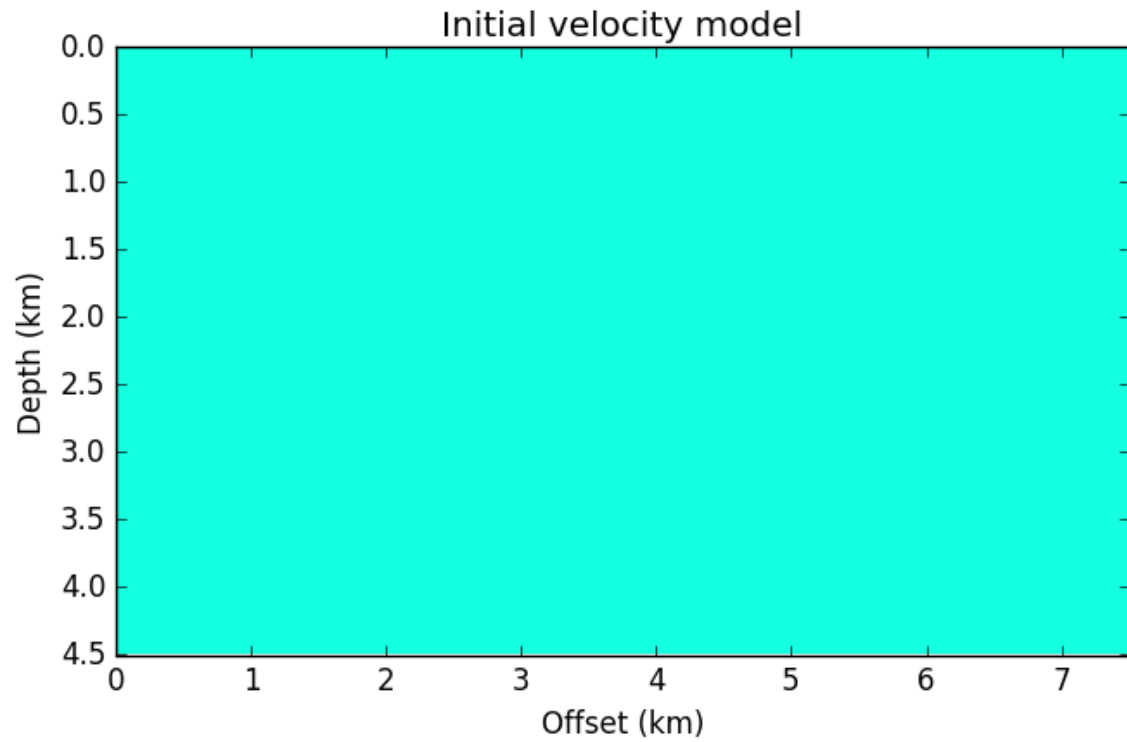
Example: initial velocity model



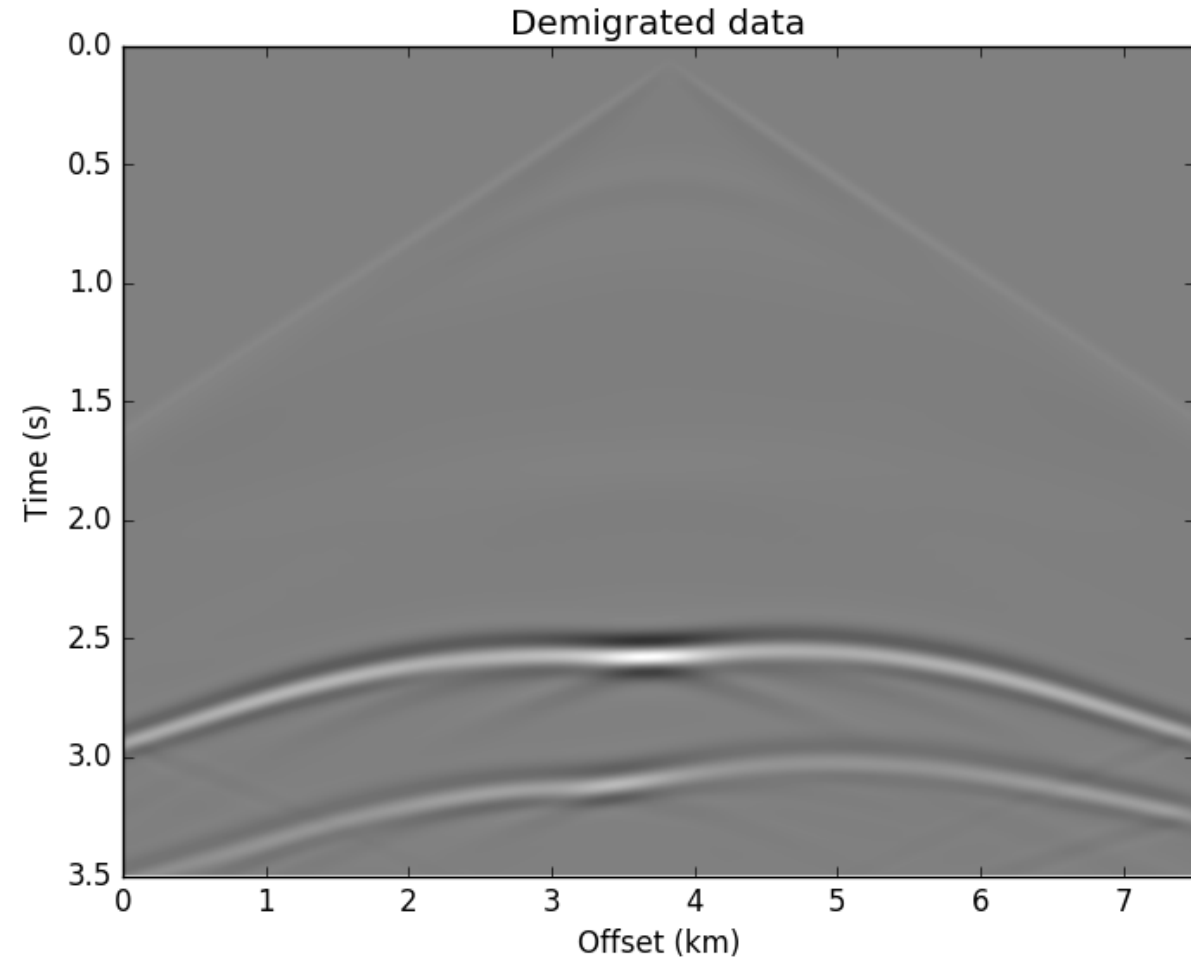
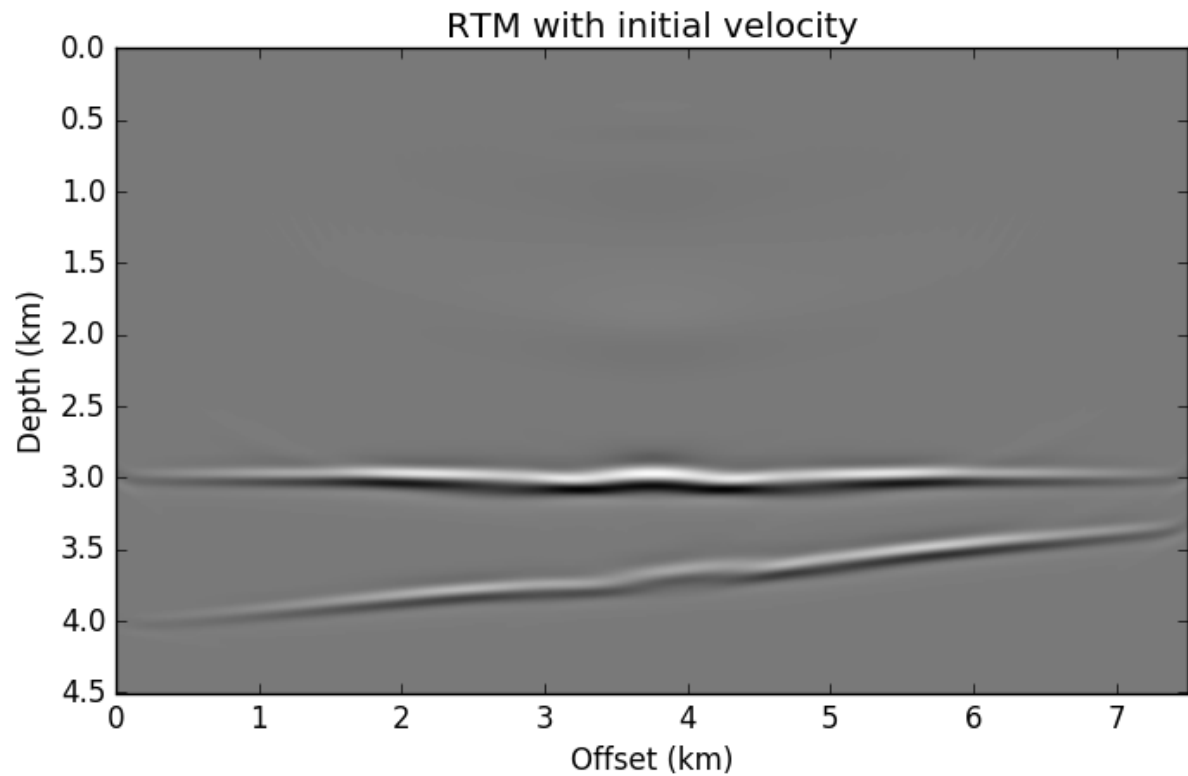
Example: observed data



Example: migration

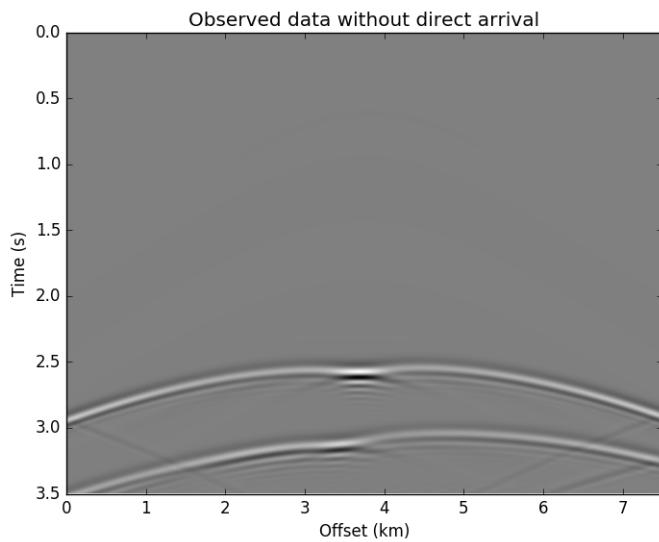


Example: demigration 1

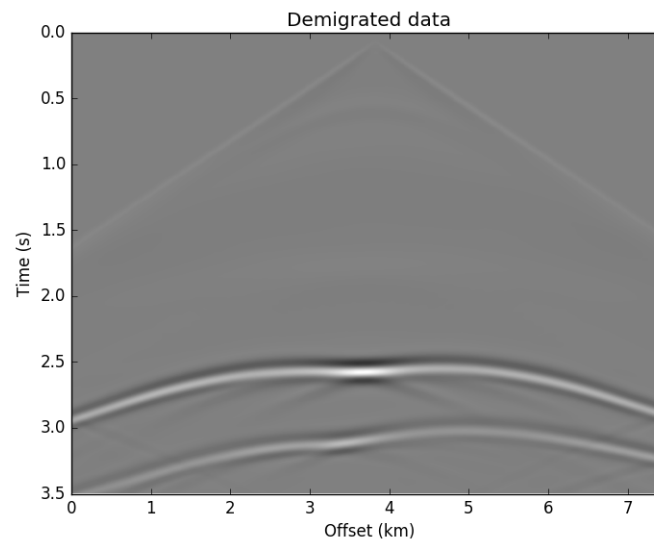


Example: residual, back-propagation, and demigration 2

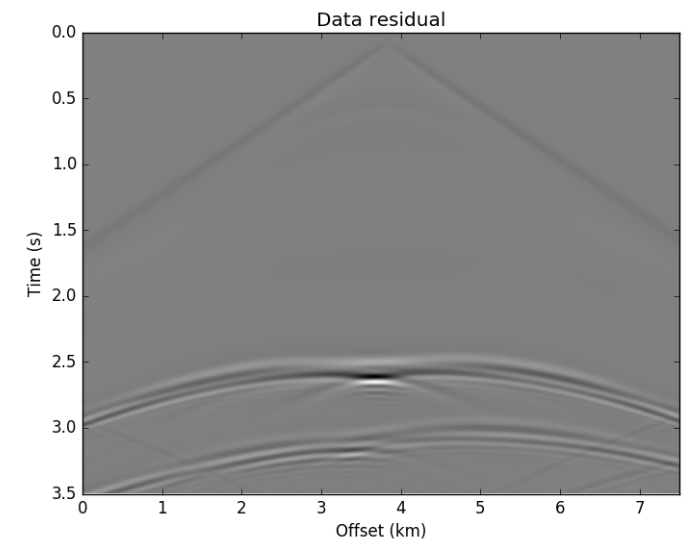
Observed data



Demigrated data

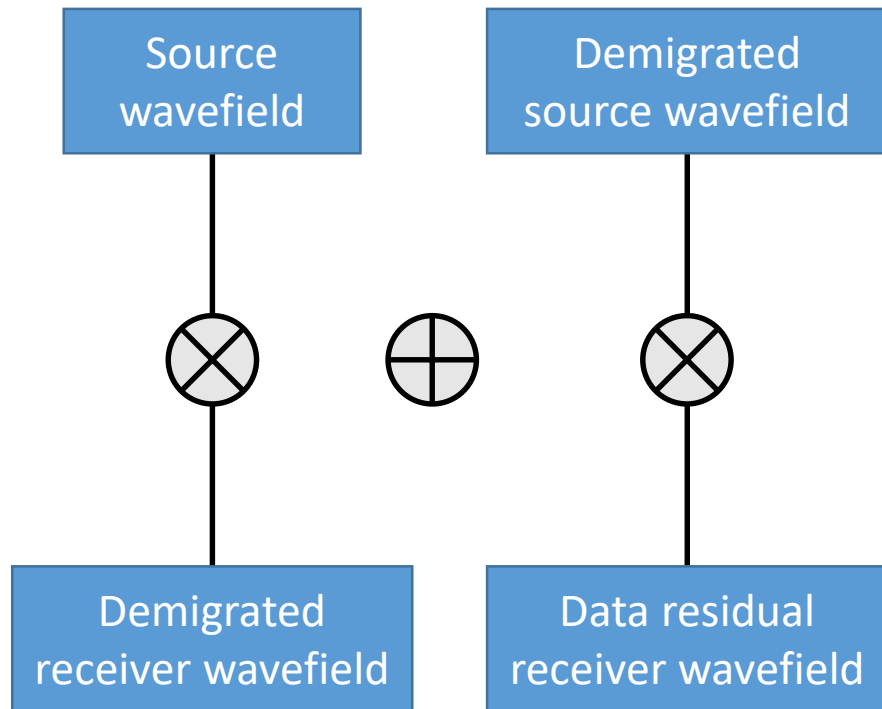


Data residual

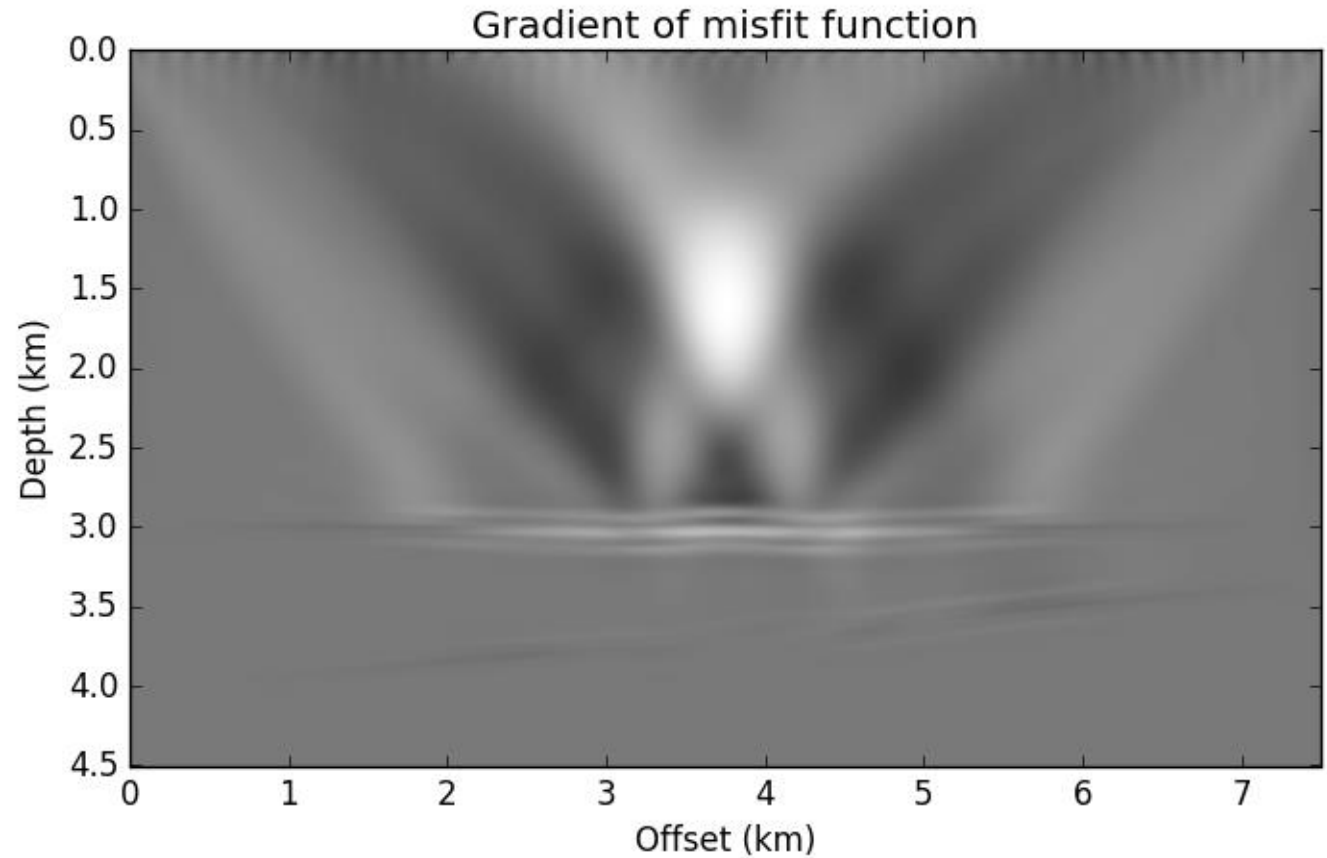


$$d_{obs} - d_{cal} = \delta d$$

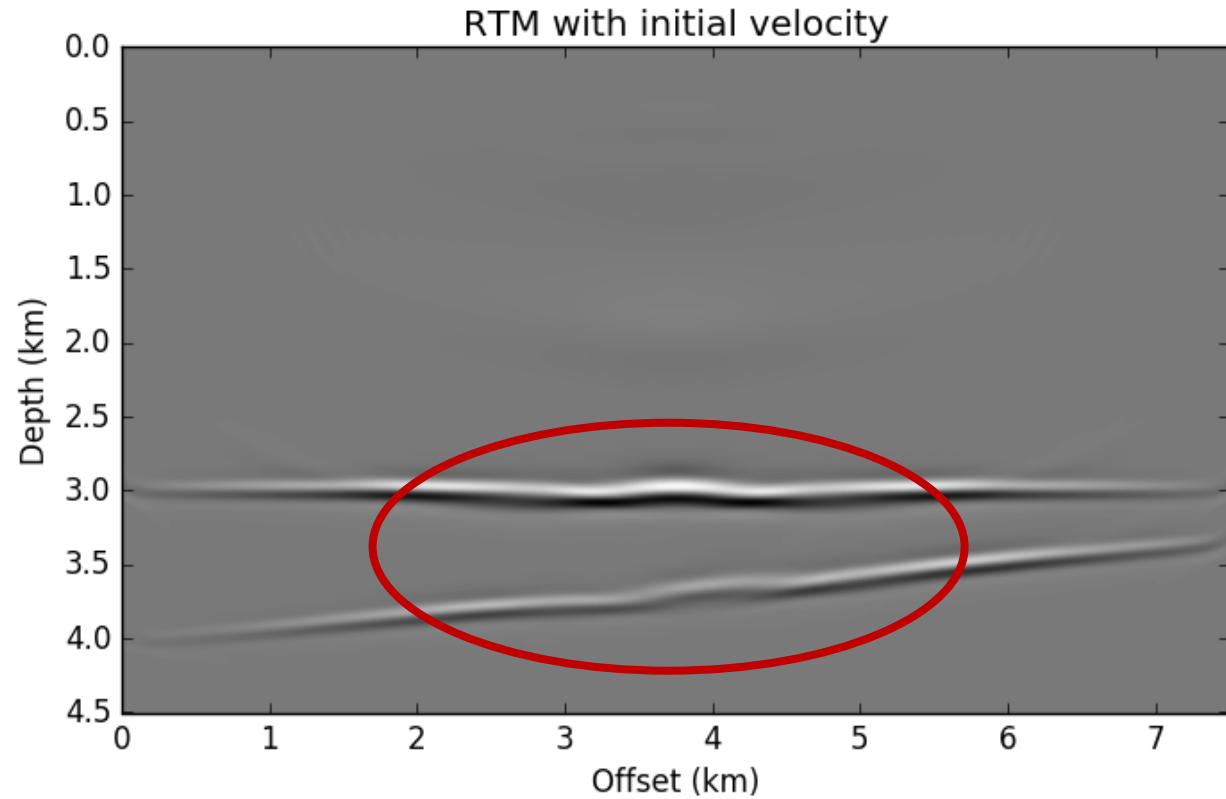
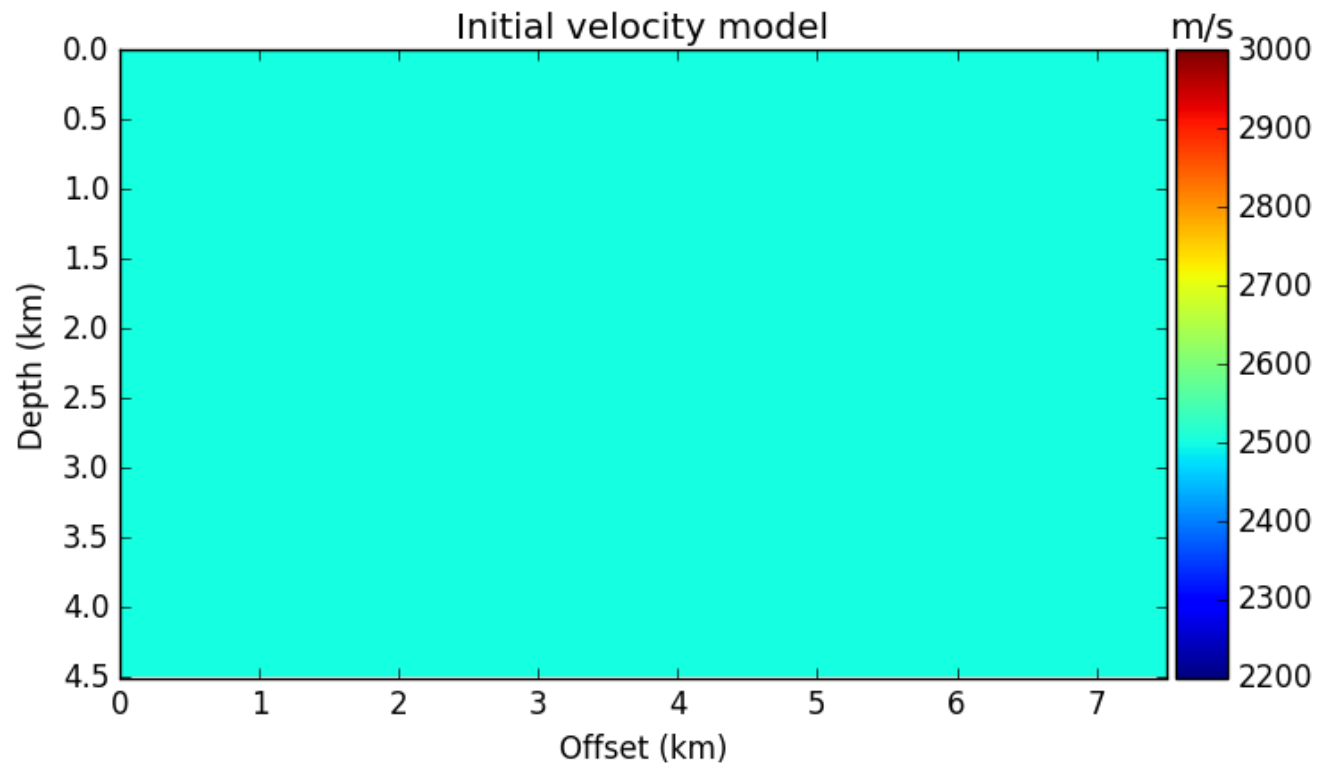
Example: gradient



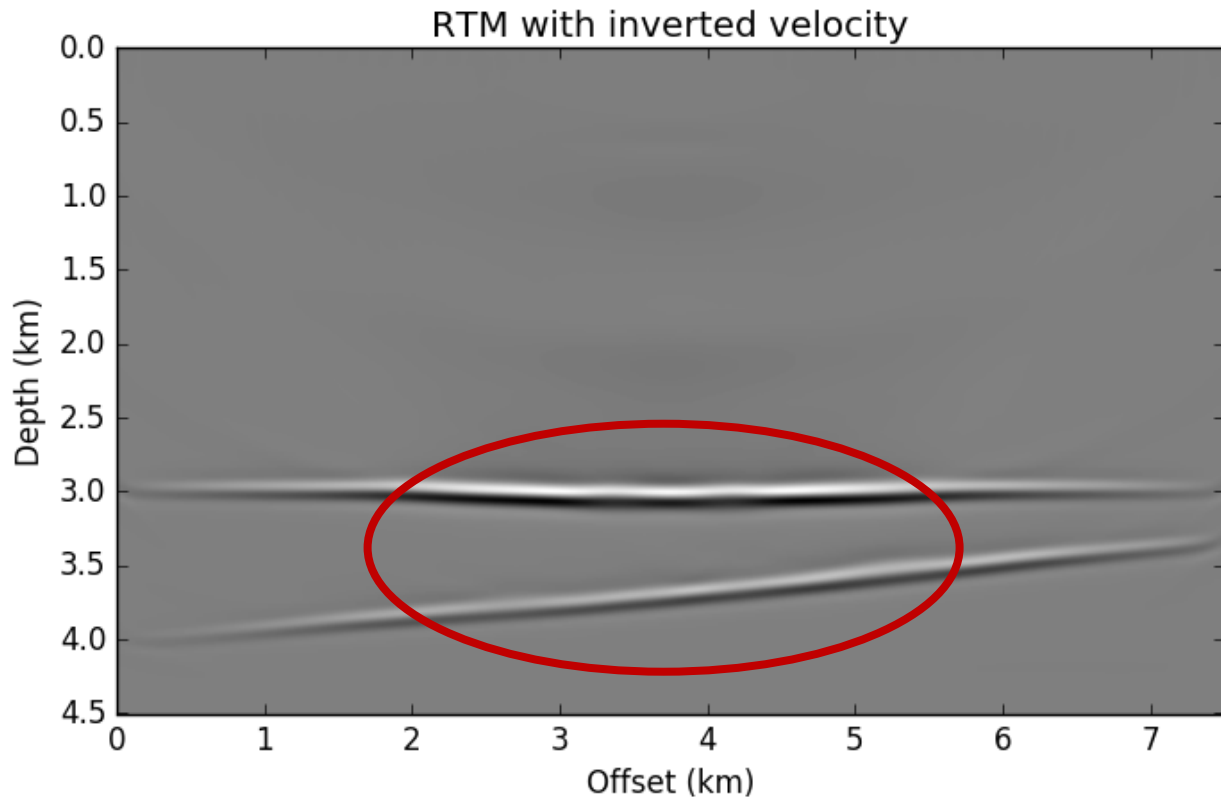
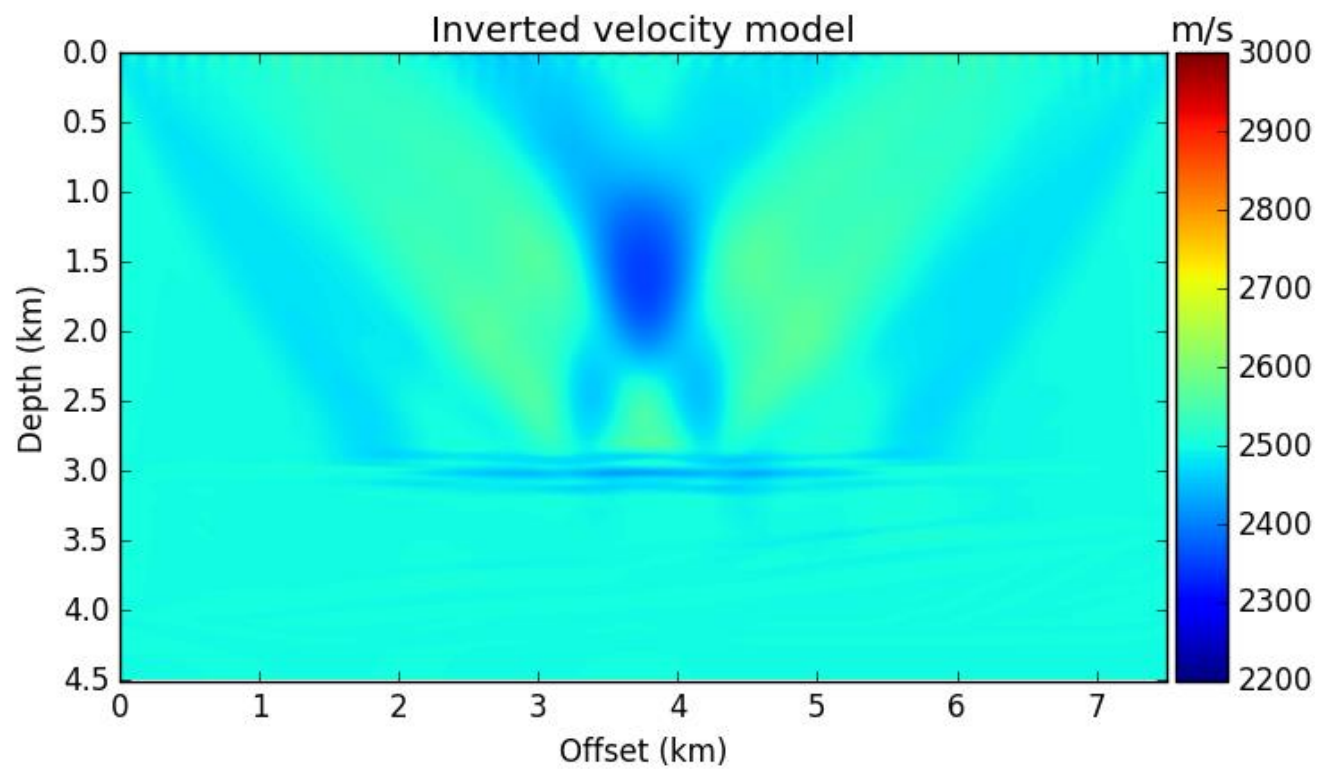
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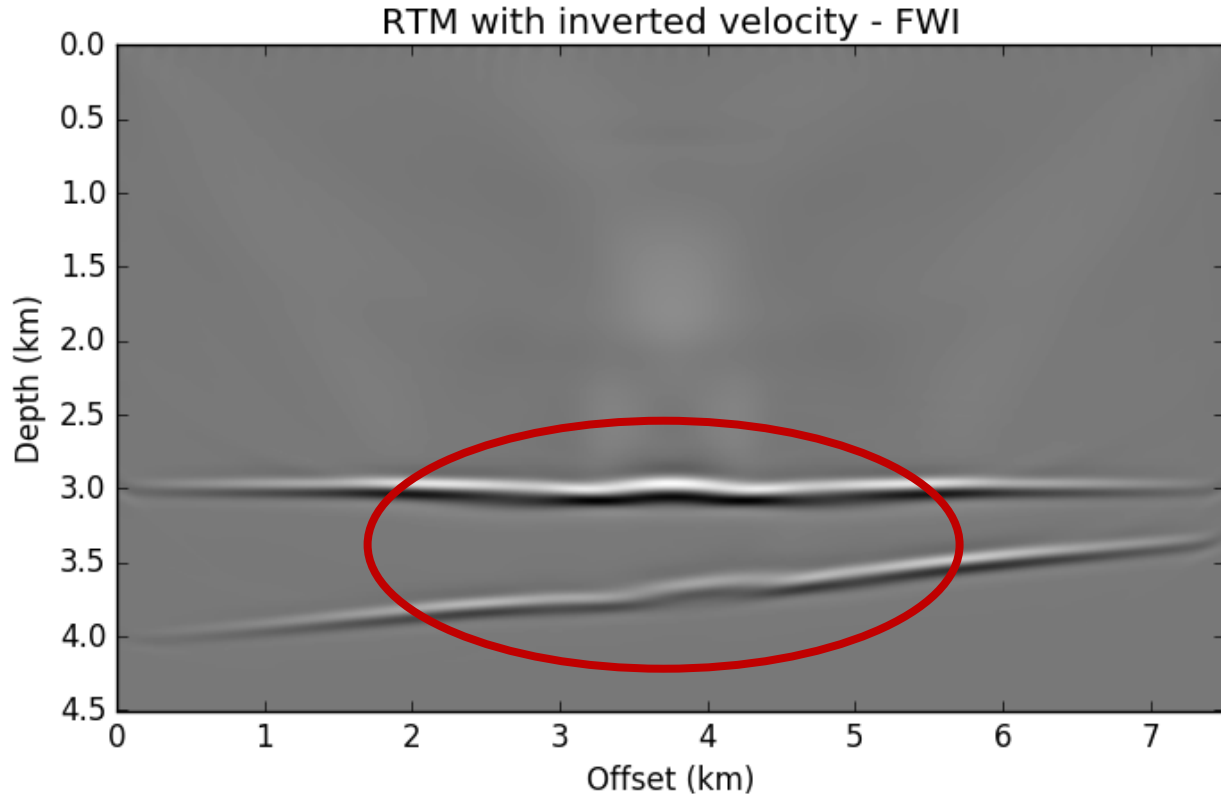
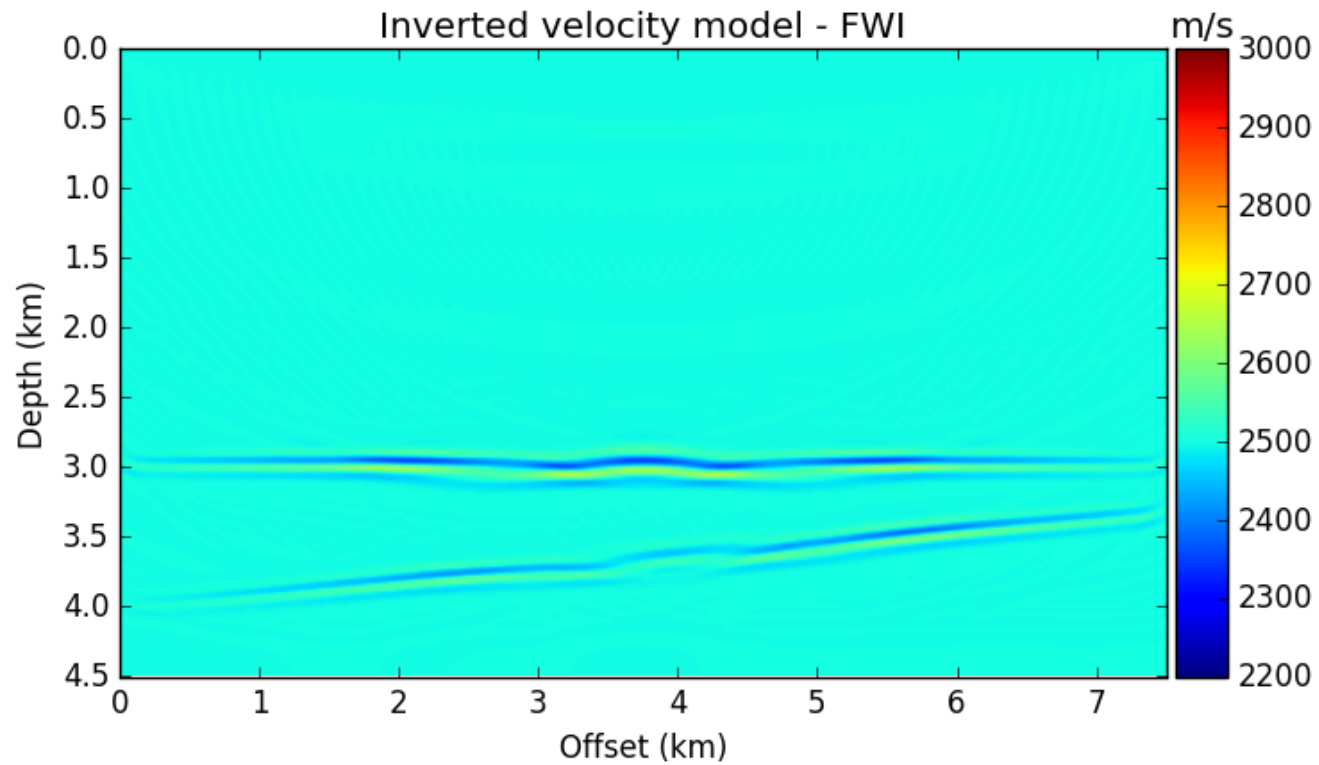
Example: results – initial velocity



Example: results – RWI



Example: results - FWI



- Neglects other waves (e.g. direct waves).
- RTM artifacts, if not handled properly, will contaminate the demigrated data.
- Computationally expensive.

Conclusion

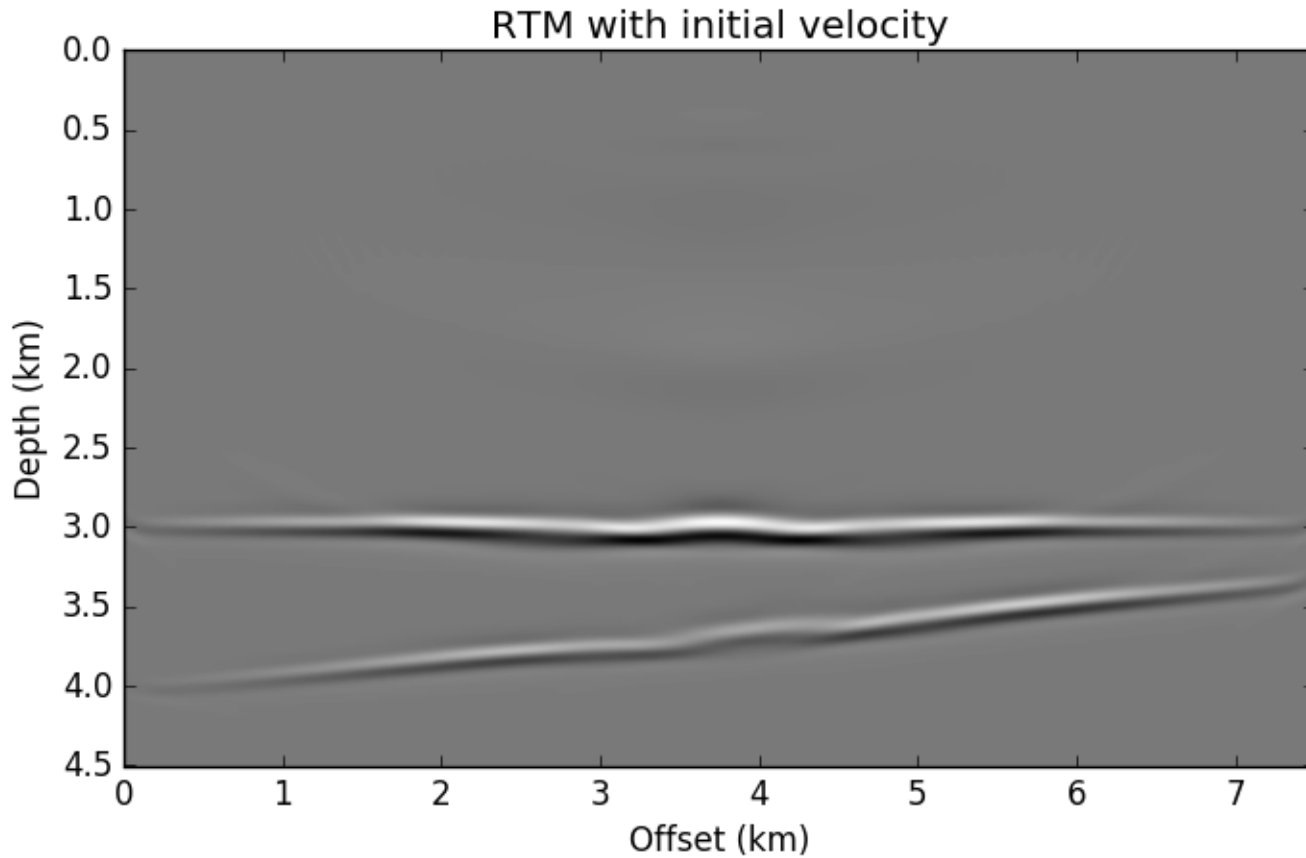
- RWI provides a better environment to retrieve the long-wavelength components in the deep part of the velocity model.
- RWI could provide a good initial velocity model for FWI.
- Future work:
 - Implementing a multi-parameter RWI.
 - Implementing a joint RWI that include direct waves.

Acknowledgements

- CREWES sponsors.
- NSERC (grant 461179-13)
- Scott Keating
- CREWES faculty, staff, and students

Questions?

Appendix: Equations

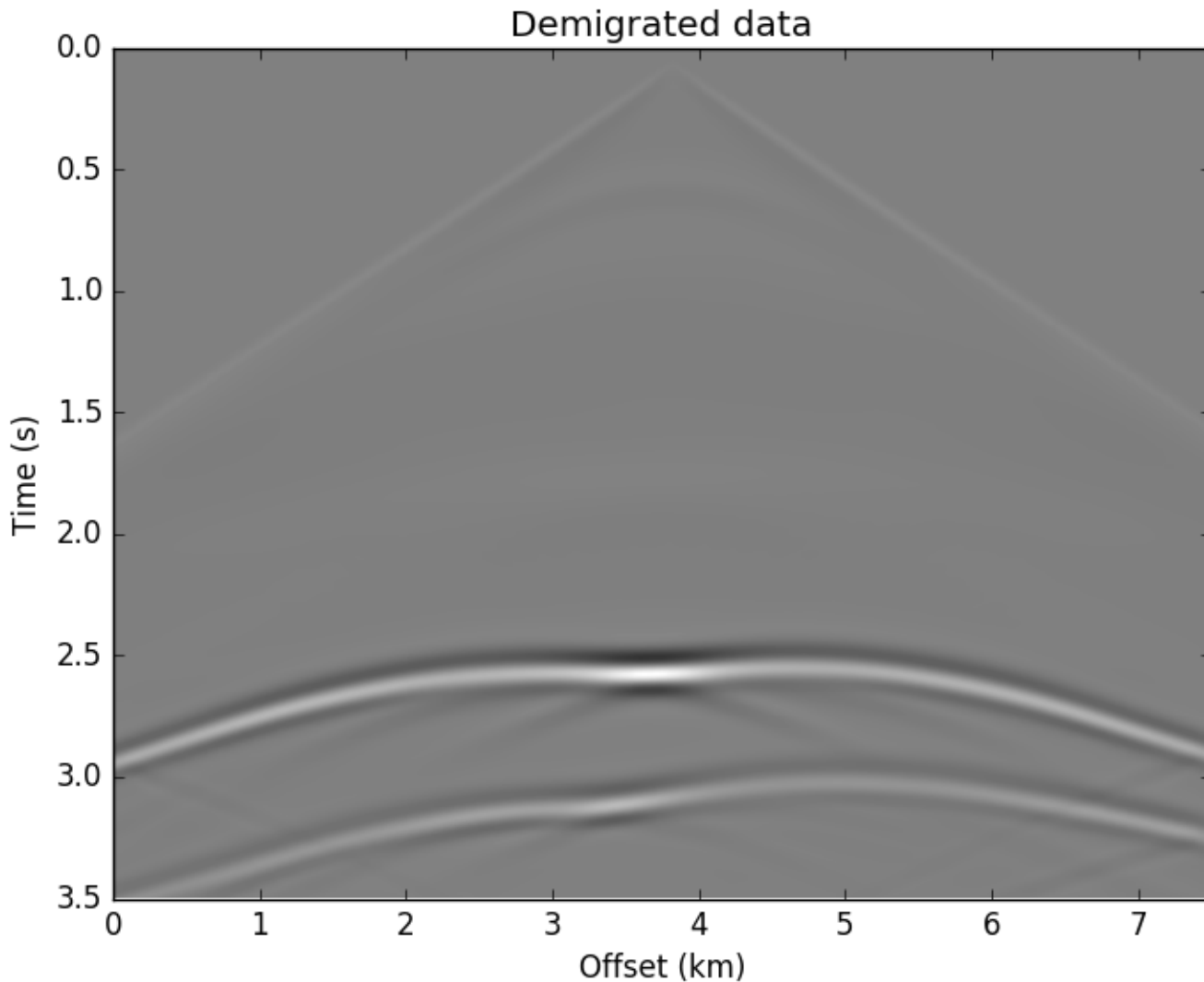


$$\left(\frac{1}{v(x)^2} \frac{\partial^2}{\partial t^2} - \nabla^2 \right) U_s(\vec{x}, t; \vec{x}_s) = f(t) \delta(\vec{x} - \vec{x}_s)$$

$$\left\{ \begin{array}{l} \left(\frac{1}{v(x)^2} \frac{\partial^2}{\partial t^2} - \nabla^2 \right) U_r(\vec{x}, t; \vec{x}_s) = 0 \\ U_r(\vec{x}_r, t; \vec{x}_s) = d_{obs}(\vec{x}_r, t; \vec{x}_s) \end{array} \right\}$$

$$I(\vec{x}) = \sum_{x_s} \int_0^T U_s(\vec{x}, t; \vec{x}_s) \cdot U_r(\vec{x}, T - t; \vec{x}_s) dt$$

Demigration 1

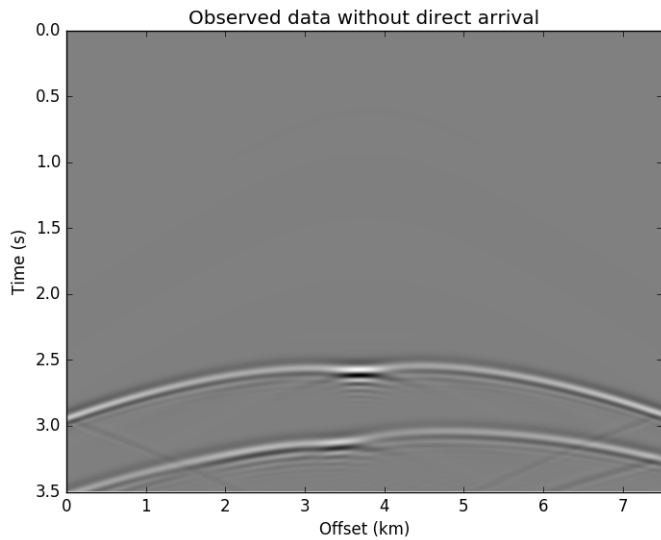


$$\left(\frac{1}{v(x)^2} \frac{\partial^2}{\partial t^2} - \nabla^2 \right) U_s(\vec{x}, t; \vec{x}_s) = f(t) \delta(\vec{x} - \vec{x}_s)$$

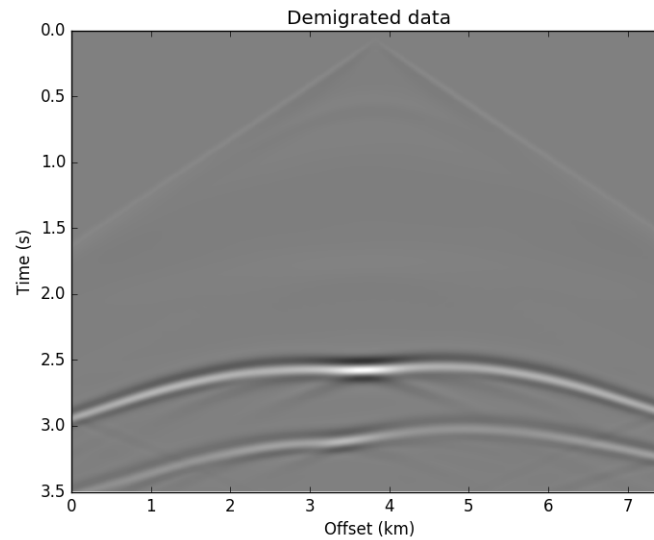
$$\left(\frac{1}{v(x)^2} \frac{\partial^2}{\partial t^2} - \nabla^2 \right) \delta U_s(\vec{x}, t; \vec{x}_s) = I(\vec{x}) \cdot U_s(\vec{x}, t; \vec{x}_s)$$

$$d_{cal} = \delta U_s(\vec{x}_r, t; \vec{x}_s)$$

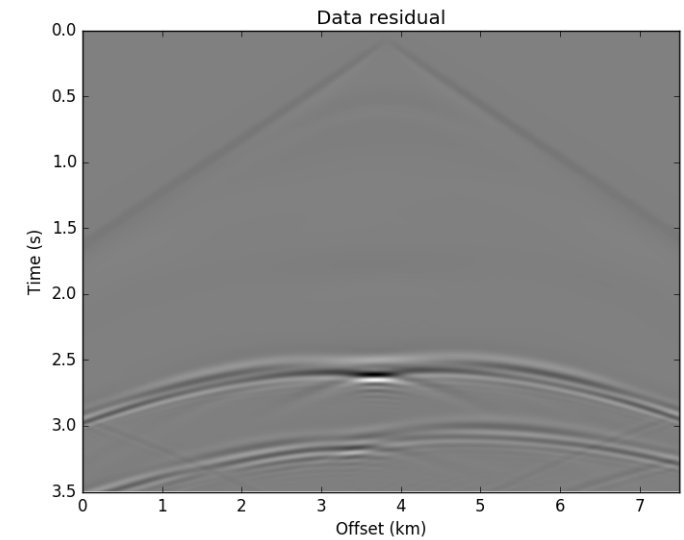
Observed data



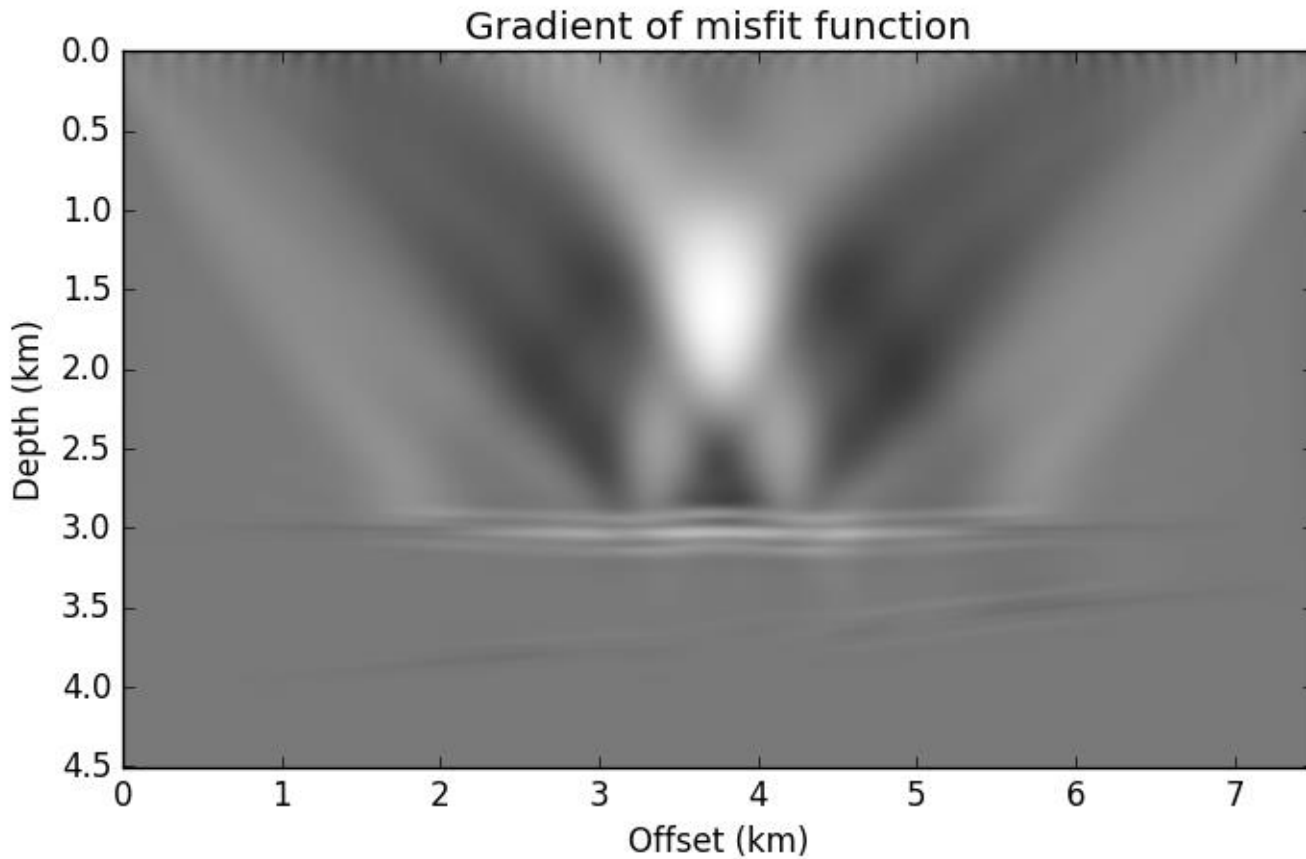
Demigrated data



Data residual



$$d_{obs} - d_{cal} = \delta d$$



$$\left\{ \begin{array}{l} \left(\frac{1}{v(x)^2} \frac{\partial^2}{\partial t^2} - \nabla^2 \right) U_r(\vec{x}, t; \vec{x}_s) = 0 \\ U_r(\vec{x}_r, t; \vec{x}_s) = \delta d(\vec{x}_r, t; \vec{x}_s) \end{array} \right\}$$

$$\left(\frac{1}{v(x)^2} \frac{\partial^2}{\partial t^2} - \nabla^2 \right) \delta U_r(\vec{x}, t; \vec{x}_s) = I(\vec{x}) \cdot U_r(\vec{x}, t; \vec{x}_s)$$

$$k_s = \sum_{x_s} U_r(\vec{x}, t; \vec{x}_s) \cdot \delta U_s(\vec{x}, t; \vec{x}_s)$$

$$k_r = \sum_{x_s} U_s(\vec{x}, t; \vec{x}_s) \cdot \delta U_r(\vec{x}, t; \vec{x}_s)$$

$$G = k_s + k_r$$